

The 20th Century Reanalysis (20CR): Capturing 200 years of weather using surface observations

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SYNOPTIC WEATHER MAP
NORTHERN HEMISPHERE
SEA LEVEL 1300 GMT
AUG 16 1915

CIRÉS

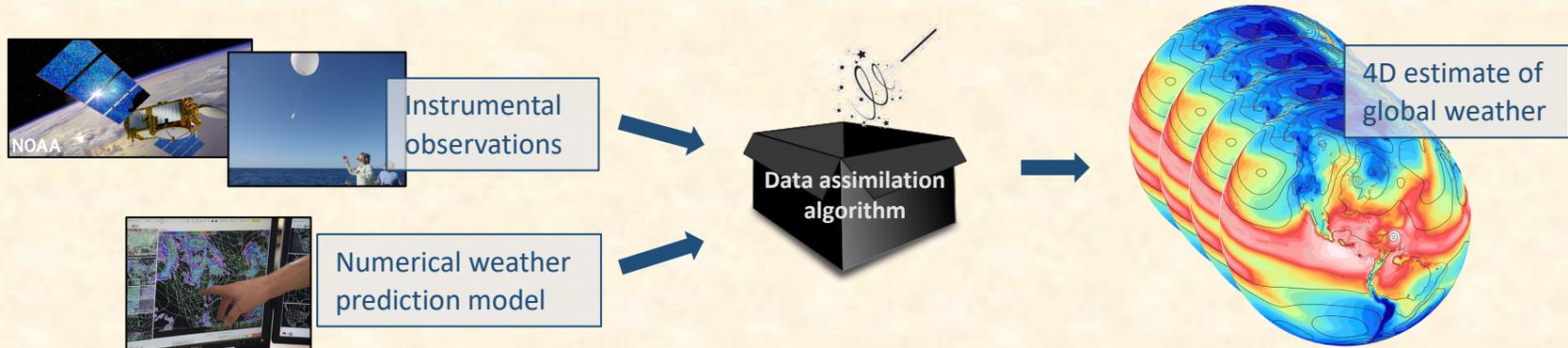


Outline

- Reanalysis fundamentals
- Full input vs sparse input reanalysis
- 20CRv3: methods
- 20CRv3: performance
- Outlook: possibilities for 20CRv4

Reanalysis fundamentals

- Reanalyses provide a consistent, gridded record of weather & climate by assimilating historical observations into a modern weather forecast model (“re-analyzing” the data.)
- To achieve consistency, a single forecast model and single data assimilation method are fixed.
 - To some degree, the observing network may also be fixed.



Full input vs sparse input reanalysis

- **Full input**

- ERA-interim, ERA5, MERRA, MERRA2
- Assimilate most observations that are available (*in-situ, satellite, upper-air, aircraft*)
- Cover latter half of 20th century to avoid spurious trends and signals arising from significant changes in the observing system
- ...Can still be impacted by instruments coming online

- **Sparse input**

- 20th Century Reanalysis, CERA-20C
- Assimilate only surface observations (surface pressure, optionally marine winds)
- Extend 100+ years into the past
- Less impact from changes in observing network

Full input vs sparse input reanalysis

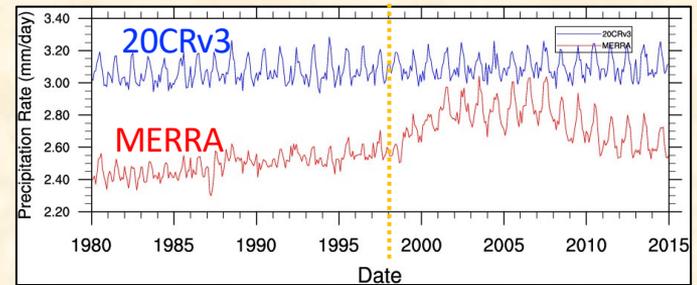
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- **Sparse input**

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Monthly mean global precipitation



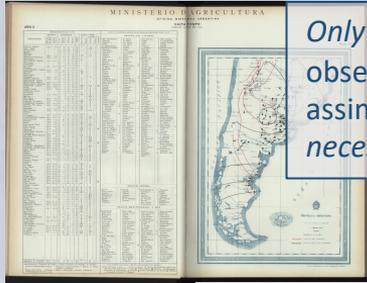
Example: spurious trend in global

precipitation from observing network change

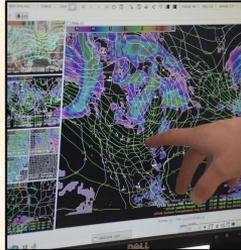
- MERRA: model may have dry bias that was corrected by water vapor sensitive radiances from AMSU, which came online in 1998
- 20CRv3 does not assimilate any radiances, therefore does not exhibit this discontinuity

20th Century Reanalysis (20CR)

200 years of weather (1806-2015):
20th Century Reanalysis Version 3

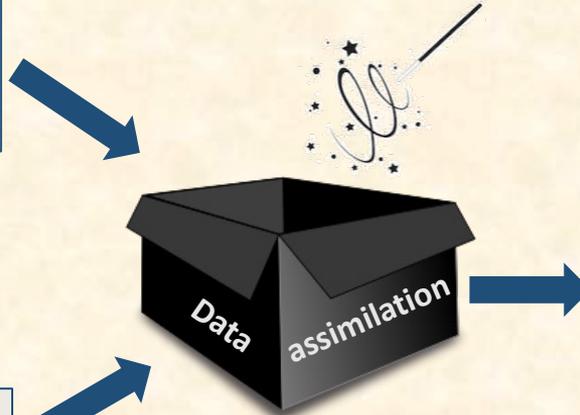


Only surface pressure observations are assimilated (*other obs necessary for SSTs, ice*)

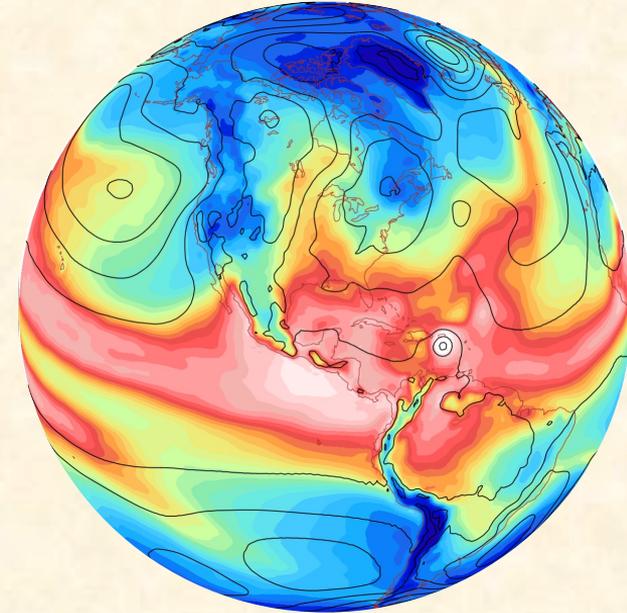


Numerical Weather Prediction model:

- GFSv14, ~0.75deg
- Prescribed sea surface temperature (SST), ice coverage, and radiative forcings



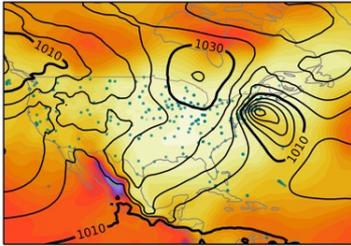
Ensemble Kalman Filter with 80 ensemble members.
➤ EnKF can dynamically move sparse information around via covariances.



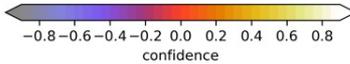
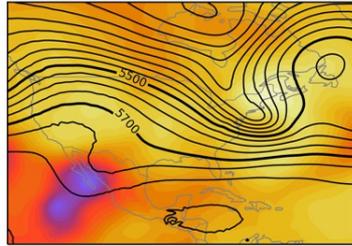
The 20th Century Reanalysis (20CR) provides a global, 200-year history of sub-daily weather by assimilating *only* surface pressure observations into a modern weather model

NOAA-CIRES-DOE 20CRv3, 13 Mar 1888 (0Z)

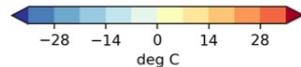
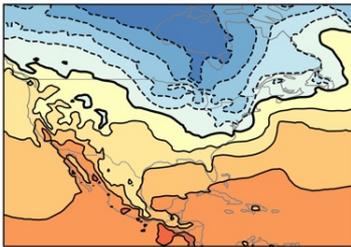
(a) Ens. mean SLP, obs., & confidence



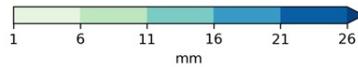
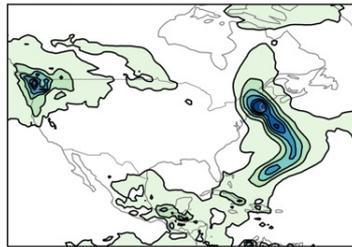
(b) Ens. mean Z500 and confidence



(c) Ens mean. 2m air temperature



(d) Ens. mean precip (6h accum.)



NOAA-CIRES-DOE 20th Century Reanalysis Version 3

- Estimates temperature, wind, precipitation, pressure, humidity, & other variables, from the ground to the top of the atmosphere
- Prescribed sea surface temperature (*quasi-weakly coupled*), sea ice concentration, and radiative forcing
- Global 75km grid
- 3-hourly resolution
- Spans 1836-2015 [1806-1835 experimental]
- Data assimilation: Ensemble Kalman Filter with 80 ensemble members to quantify uncertainty

■ Publicly available: <https://go.usa.gov/XTd>

20th Century Reanalysis version 3 (20CRv3)

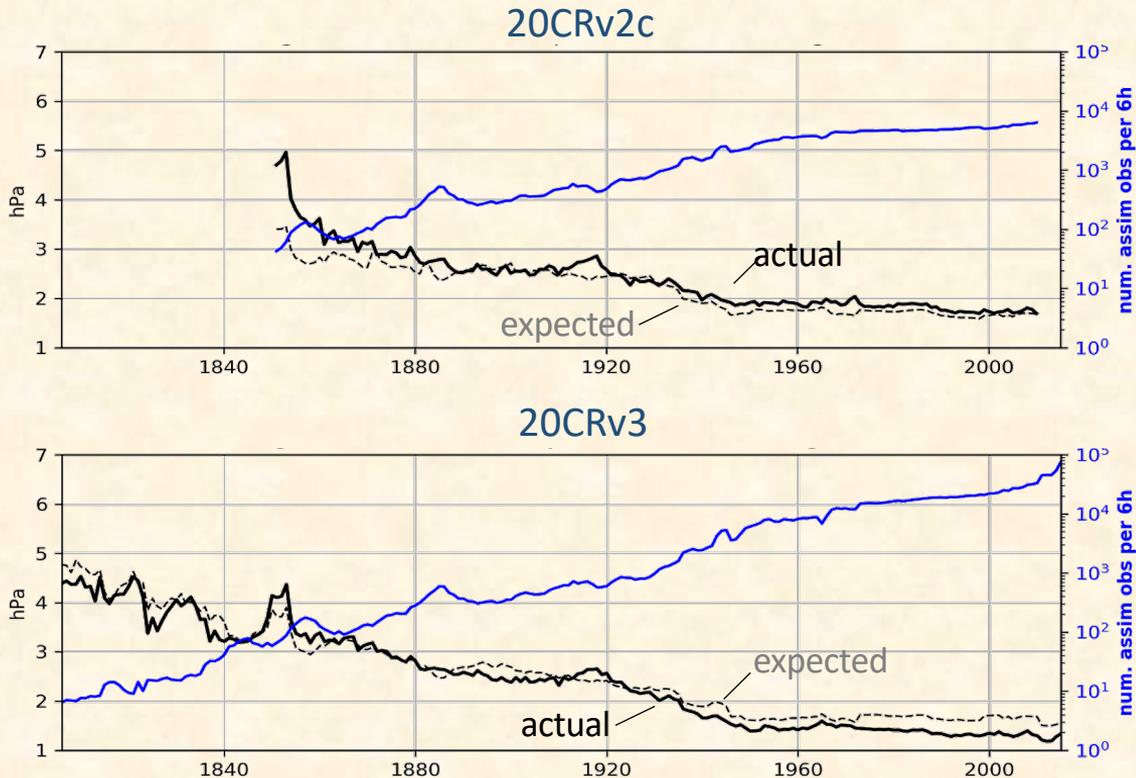
System improvements over 20CRv2c

- Newer, higher-resolution forecast model
- Larger available set of observations to be assimilated
- Improved observation quality control methods
- Improved DA methods (4D incremental analysis update, adaptive localization, adaptive inflation)
 - Improved confidence estimates

Slivinski, L.C., et. al. (2019) Towards a more reliable historical reanalysis: Improvements for version 3 of the Twentieth Century Reanalysis system. *Quarterly Journal of the Royal Meteorological Society*, **145**: 2876– 2908.
<https://doi.org/10.1002/qj.3598>

20CRv3 outperforms 20CRv2c: fit to surface pressure obs

Global annual first-guess root mean squared errors in surface pressure



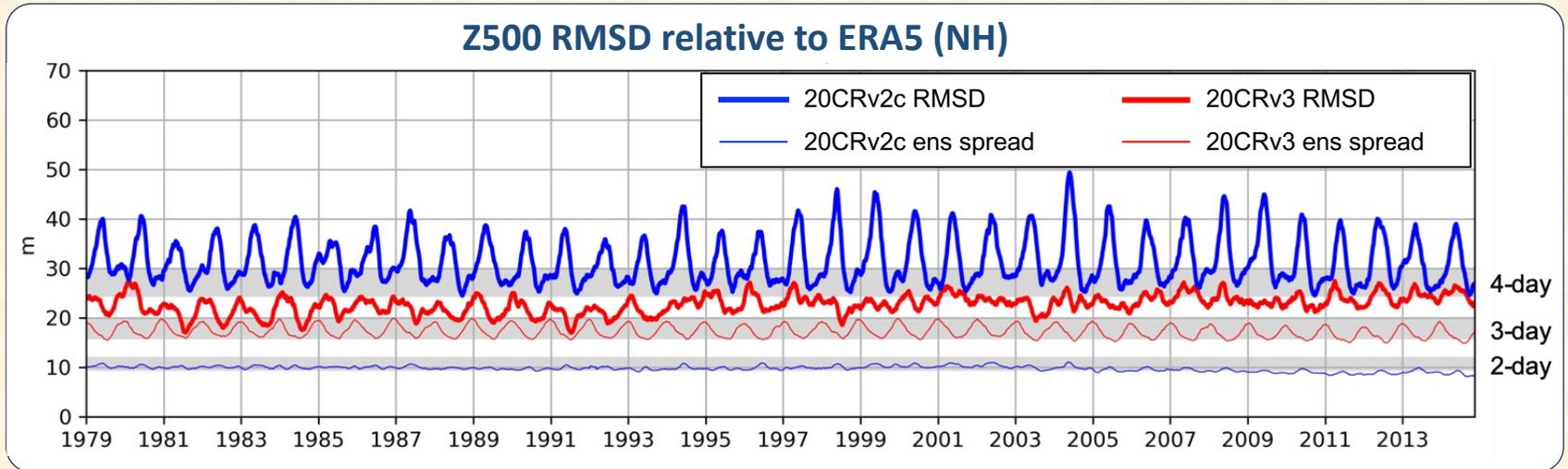
Actual error is the difference between 6-hour forecasts and not-yet-assimilated observations:
 $\langle (ob - fg)^2 \rangle^{1/2}$

Expected error is the root-mean of the sum of ob error variance and background ensemble covariance at ob time/location:

$$\langle (\sigma_{ob}^2 + \sigma_{fg}^2)^2 \rangle^{1/2}$$

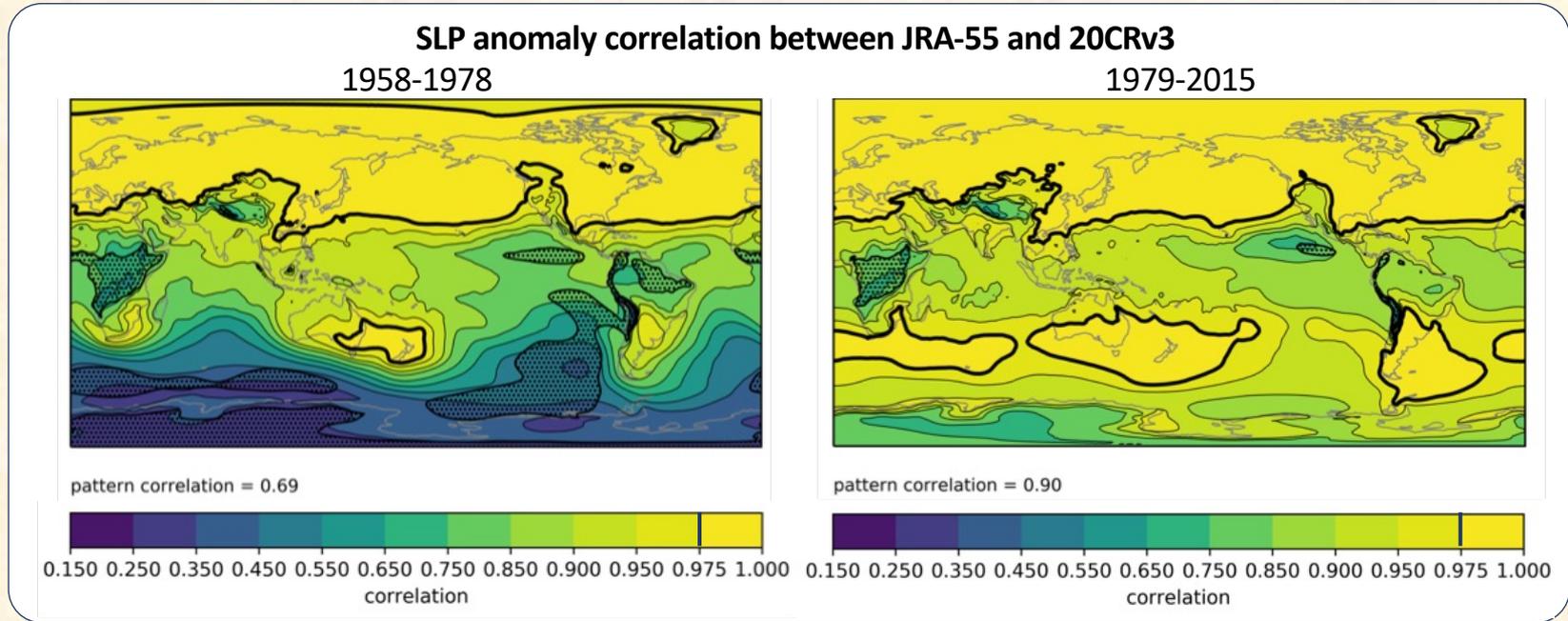
- **20CRv3 errors that are lower and more consistent with ensemble spread than 20CRv2c.**

20CRv3 outperforms 20CRv2c: 500hPa geopot. height agreement with ERA5



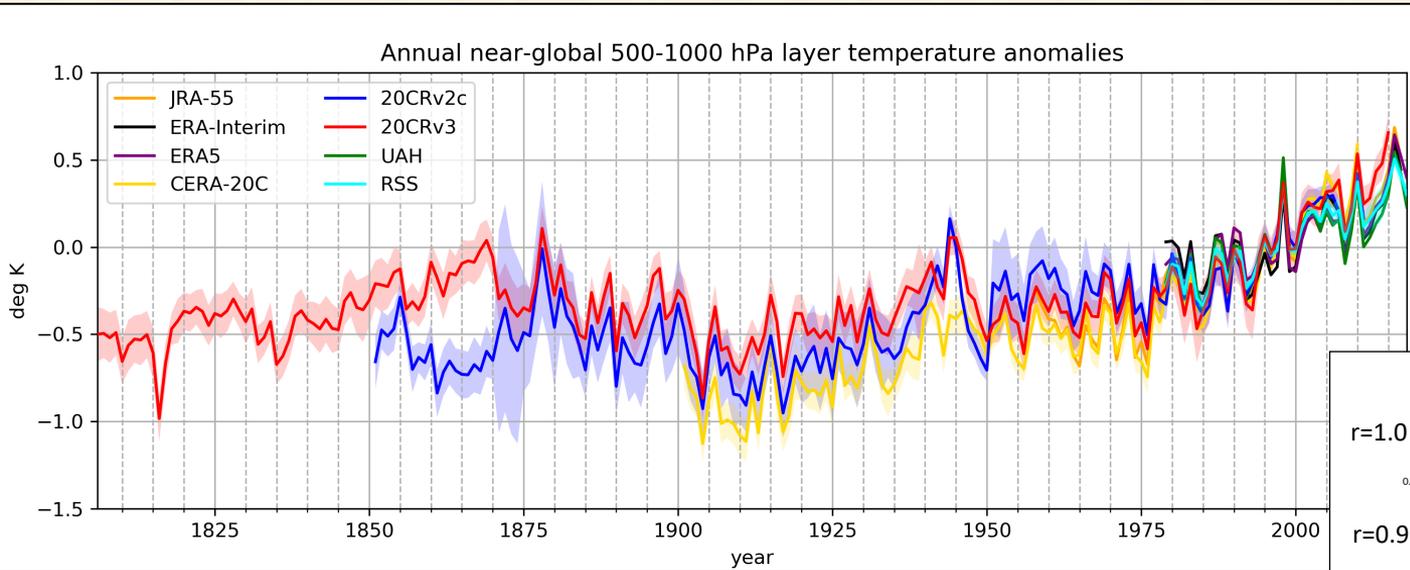
- **20CRv3 Z500 agrees with ERA5 better than 20CRv2c, and has more consistent ensemble spread**
- Gray shading shows range of ECMWF and NCEP oper. forecast errors in 2019 for 2-, 3-, 4-day leads
- **20CRv3 Z500 errors are comparable to modern 3-4 day operational forecast skill in the NH** (consistent with earlier predictions by Compo et al, 2006)

20CRv3 correlates well with other reanalyses, and can “predict” that correlation

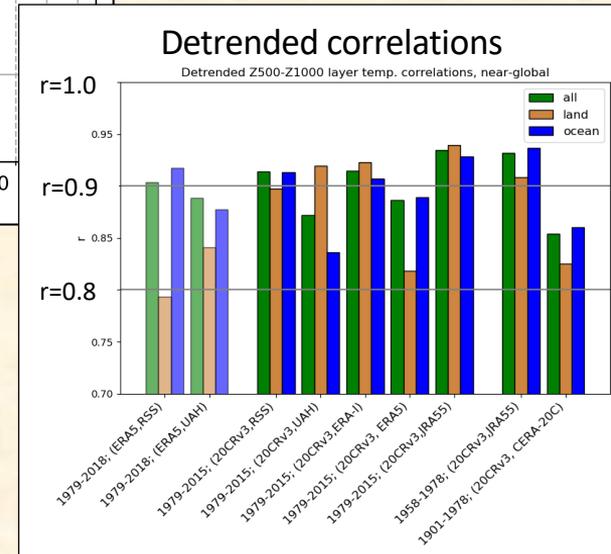


- Stippling indicates regions of low confidence (large ensemble spread) in 20CRv3
- Pattern correlation is given between confidence field and correlation field
- **20CRv3 uncertainty estimates are a good predictor of skill relative to JRA-55**

20CRv3 can capture trends & variability in temperature



UAH & **RSS**: two satellite-based temperature reconstructions



Future of 20CR – Possibilities

✓ Larger set of available observations
*(smaller errors, greater confidence,
maybe extend further back in time)*

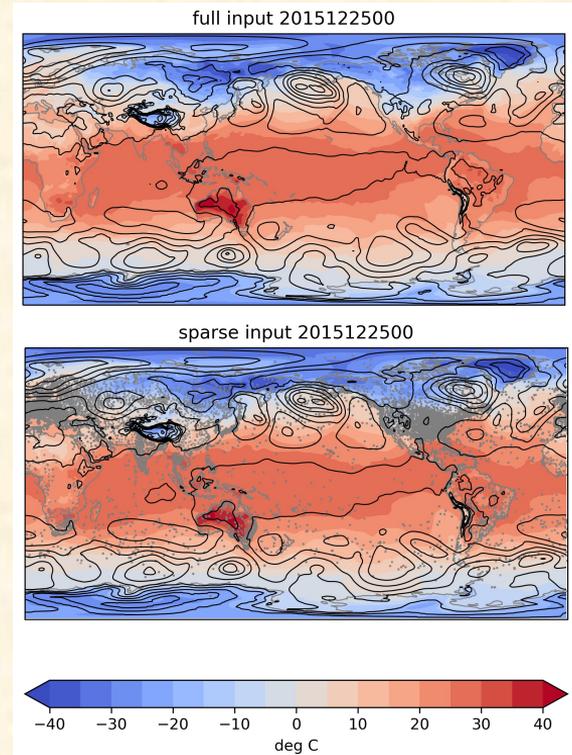
✓ Up-to-date forecast model

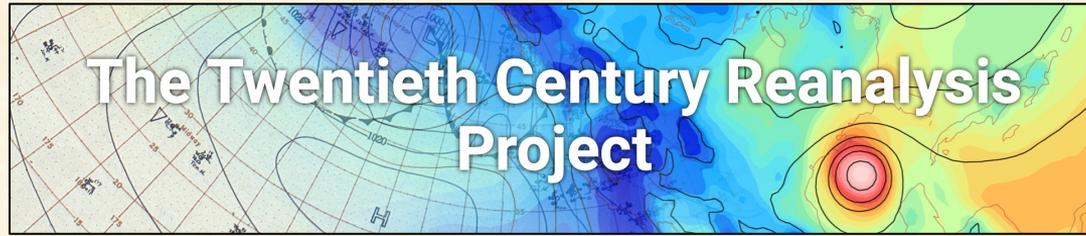
Coupled ocean-atmosphere

Additional observation types *(SST, wind
direction)*

Data-driven models *(incorporate Linear
Inverse Model [LIM] for ocean)*

No-DA counterpart simulation





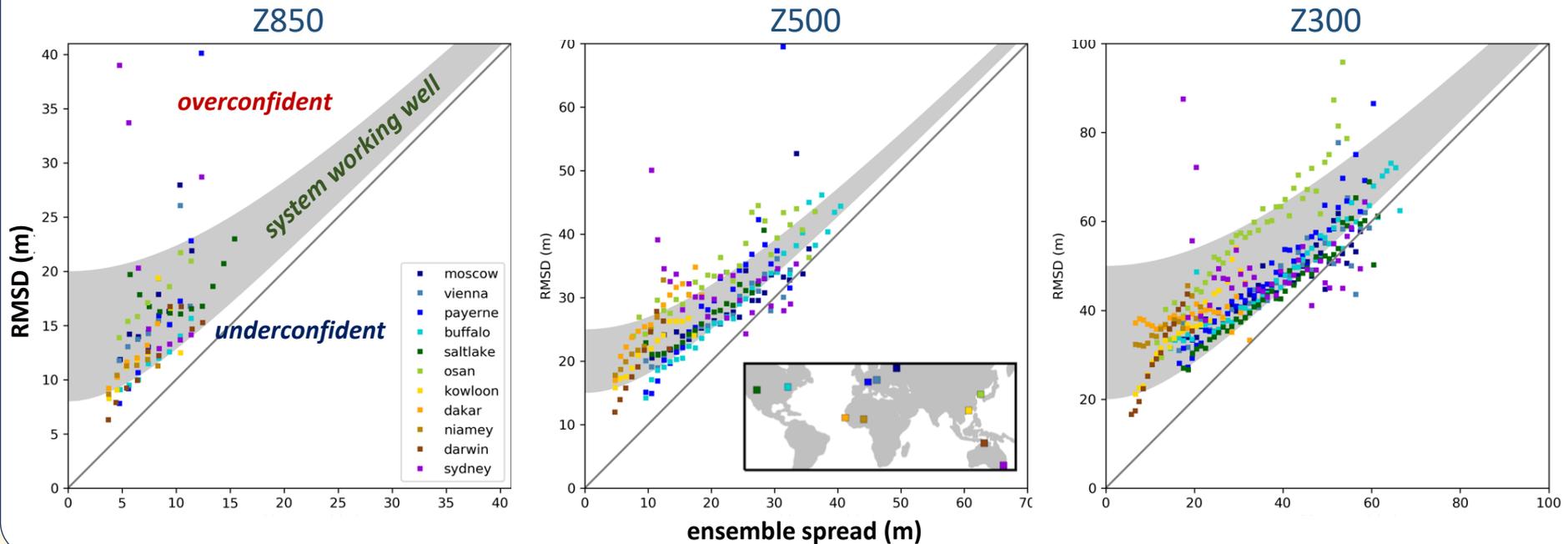
- The 20th Century Reanalysis version 3 (20CRv3) *only* assimilates surface pressure observations in order to consistently extend 200 years into the past
- It agrees well with other reanalyses, independent upper-air observations, and station/satellite-based datasets
- Confidence estimates are reliable and 20CRv3 can often predict its own skill
- More information:
 - Compo, G.P., et. al. (2011) The Twentieth Century Reanalysis Project. *Q.J.R. Meteorol. Soc.*, 137: 1-28. <https://doi.org/10.1002/qj.776>
 - Slivinski, L.C., et. al. (2019) Towards a more reliable historical reanalysis: Improvements for version 3 of the Twentieth Century Reanalysis system. *Quarterly Journal of the Royal Meteorological Society*, 145: 2876–2908. <https://doi.org/10.1002/qj.3598>
 - Slivinski, L.C., et. al. (2021) An Evaluation of the Performance of the Twentieth Century Reanalysis Version 3. *Journal of Climate*, 34(4): 1417-1438. <https://doi.org/10.1175/JCLI-D-20-0505.1>

For data access, visualization tools, and references, please visit <https://go.usa.gov/XTd>

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20CRv3 performs well relative to indep. upper air obs

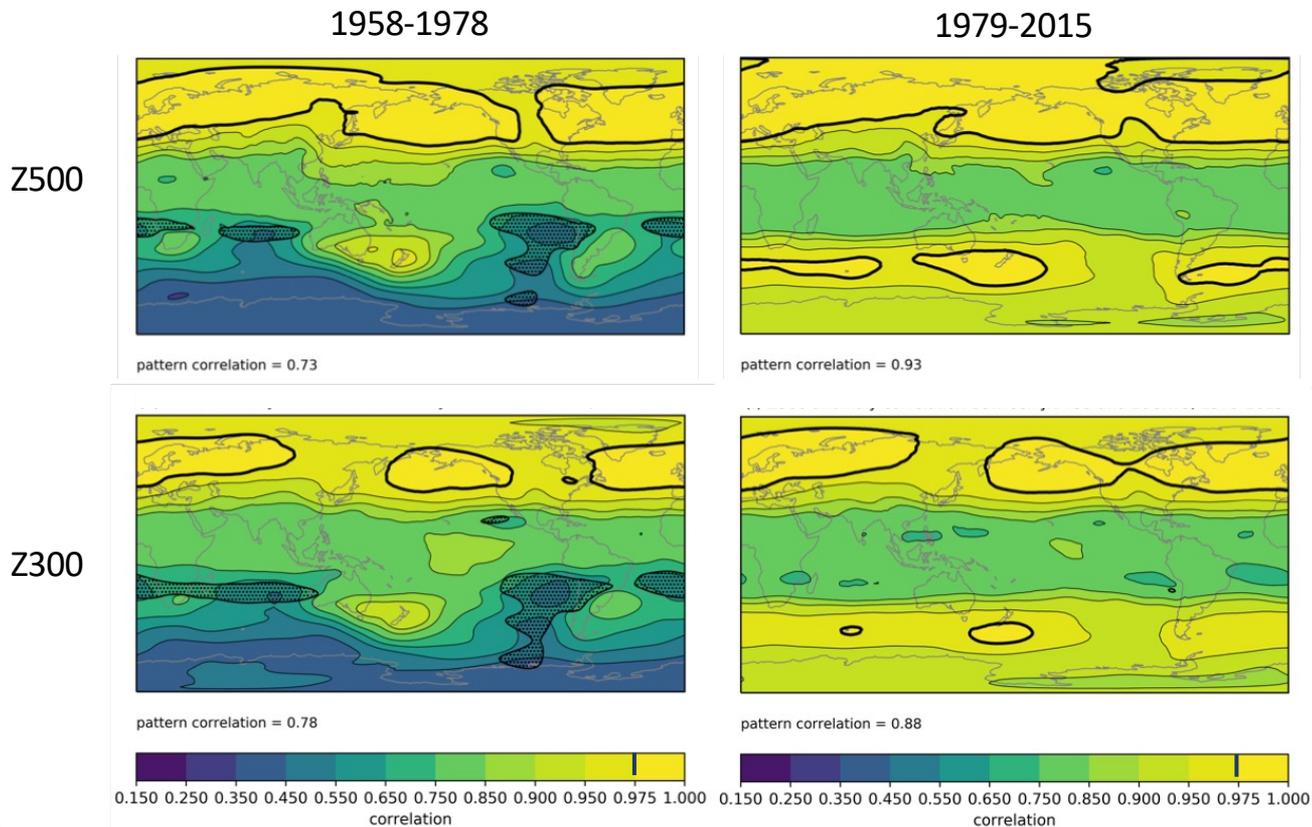
Difference between observed and analyzed values, as function of 20CRv3 ensemble spread; 1943-2015



- If obs were perfect (zero error), then RMSDs should fall on diagonal.
- If ob error range estimated accurately and system works well, RMSDs ideally fall in gray swath.
- Above swath: 20CRv3 is overconfident. Below swath: underconfident.
- **20CRv3 geopot. height analysis performs well globally at several vertical levels**

20CRv3 correlates well with other reanalyses, and can “predict” that correlation

Geopot. height anomaly correlation between JRA-55 and 20CRv3



(Similar results for 20CRv3/ERA5)

Like any EnKF, this system will suffer from filter divergence and/or spurious long-distance correlations without inflation and localization.

- **Inflation:** prevents “ensemble collapse” by artificially spreading out ensemble members.
 - Simple example: multiplying the ensemble covariance by a predefined (often tuned) factor larger than 1
- **Localization:** prevents an observation from incrementing the state at unreasonably long distances.
 - Simple example: Gaspari-Cohn localization, a function applied to the background covariance matrix which smoothly tapers long-distance correlations to zero

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Inflation

- Previous version of reanalysis used **pre-defined multiplicative inflation** factors based on year and location
- Needed larger inflation factors for densely-observed times and places; smaller inflation for sparsely-observed times and places

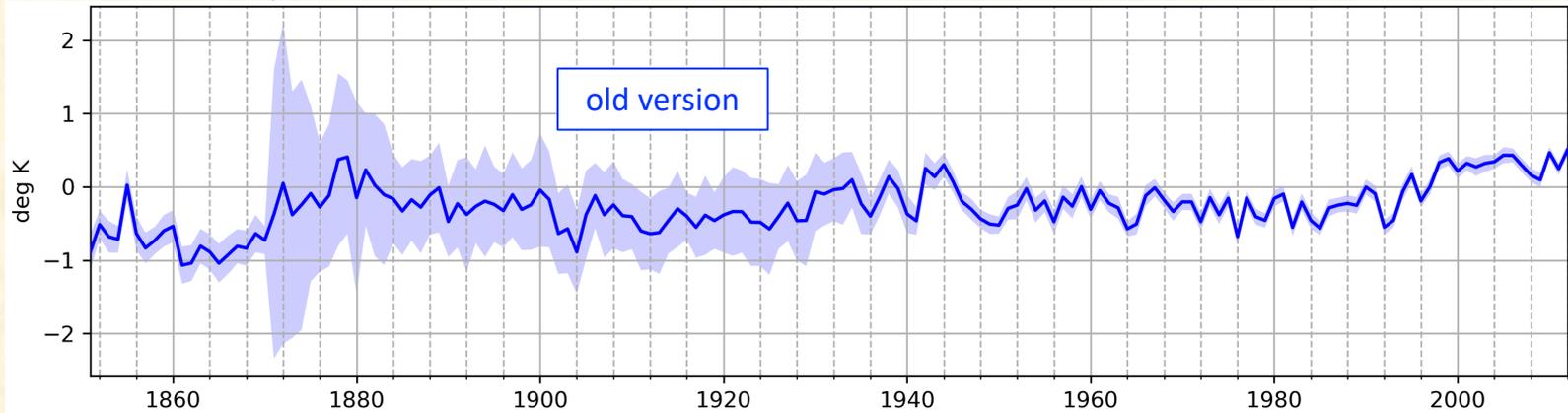
	Northern Hemisphere	Tropics	Southern Hemisphere
1851 – 1870	1.01	1.01	1.01
1871 – 1890	1.05	1.01	1.01
1891 – 1920	1.09	1.02	1.01
1921 – 1950	1.12	1.03	1.02
1951 – 2012	1.12	1.07	1.07

Simple adaptive inflation

	Northern Hemisphere	Tropics	Southern Hemisphere
1851 – 1870	1.01	1.01	1.01
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1951 – 2012	1.12	1.07	1.07

- Unrealistic signals in uncertainty
- Inhibits accurate studies of significance of long-term trends

Atmospheric layer temperature anomalies, Northern Hemisphere



New inflation: relaxation-to-prior-spread

- Adaptive inflation: larger inflation when observations are dense, smaller inflation when observations are sparse
- Inflation parameter λ_{inf} is defined as function of individual gridpoints (x,y) and timesteps (t) :

$$\lambda_{inf}(x, y, t) = p_{relax} \left(\frac{\sigma_b(x, y, t) - \sigma_a(x, y, t)}{\sigma_a(x, y, t)} \right) + 1,$$

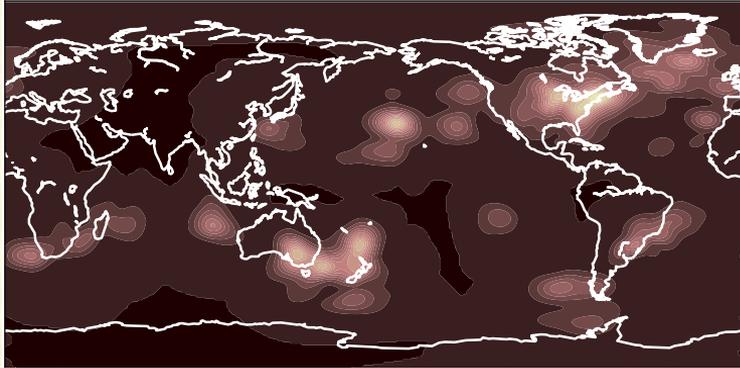
σ_b is background ensemble standard deviation,

σ_a is analysis ensemble standard deviation, and

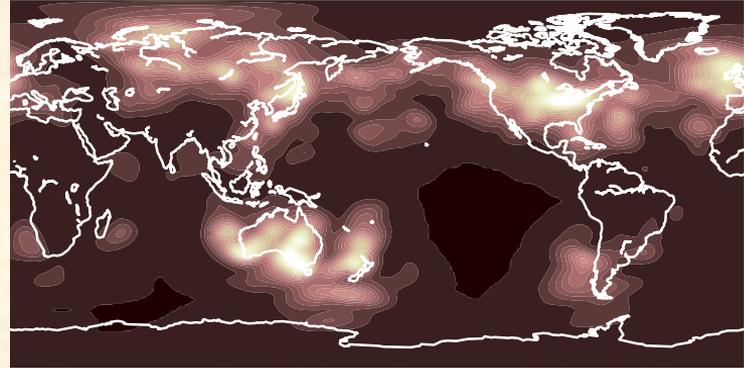
p_{relax} is a parameter varying from 0 (no inflation) to 1 (inflate fully to prior spread)

New inflation: relaxation-to-prior-spread

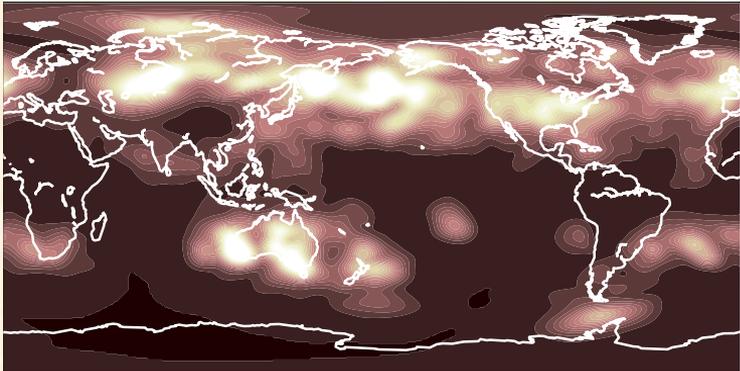
1854



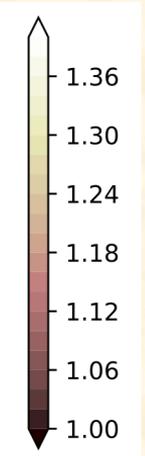
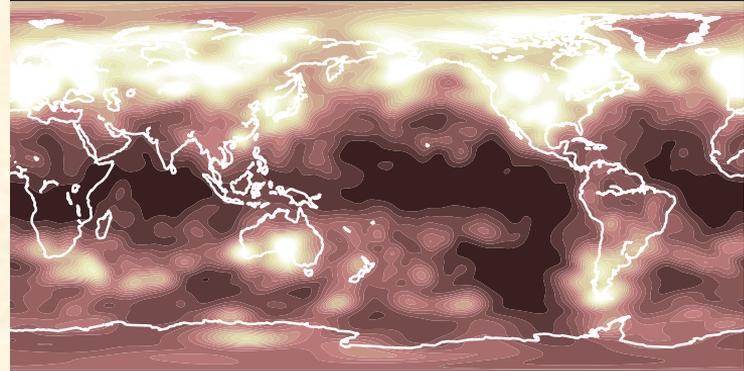
1915



1935

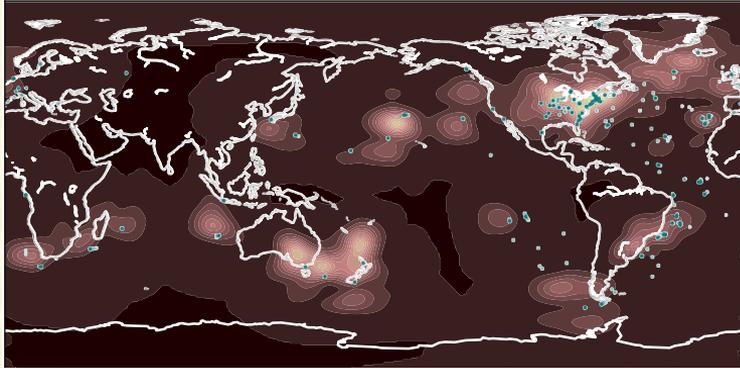


2000

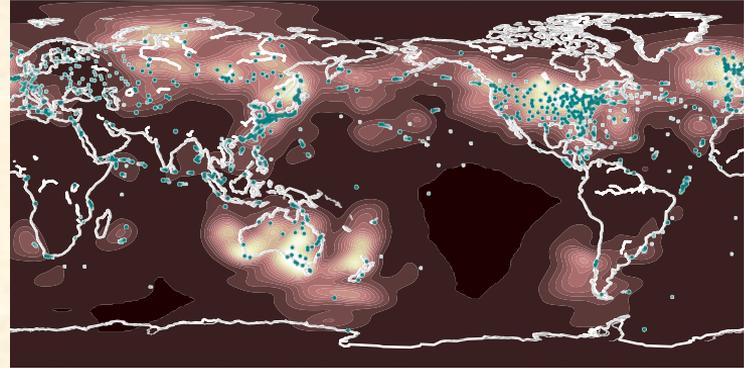


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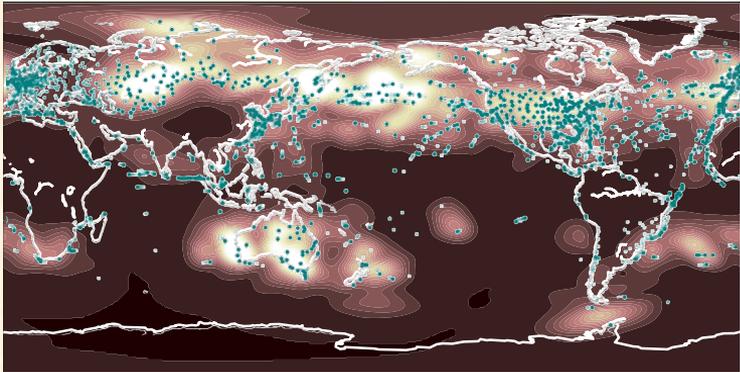
1854



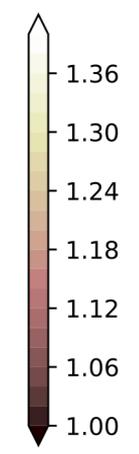
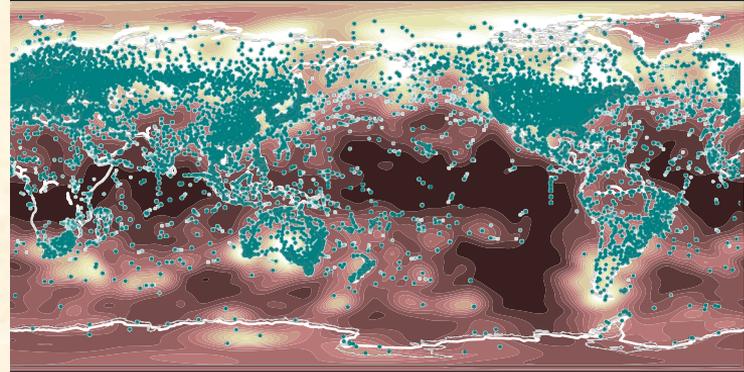
1915



1935



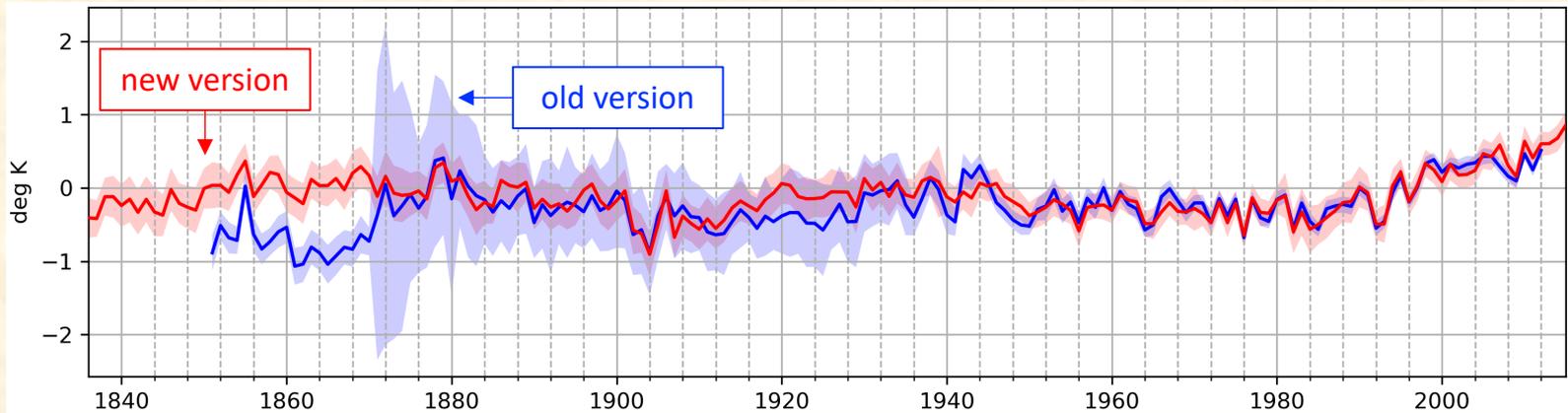
2000



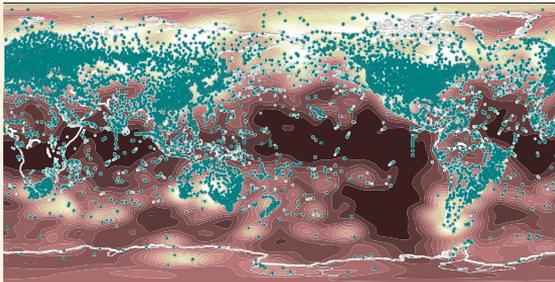
New inflation: relaxation-to-prior-spread

- More accurate, consistent estimates of uncertainty
- Can make stronger statements about trends

Atmospheric layer temperature anomalies, Northern Hemisphere



- **Inflation:** prevents “ensemble collapse” by artificially spreading out ensemble members.
 - Simple example: multiplying the ensemble covariance by a predefined (often tuned) factor larger than 1
- **Localization:** prevents an observation from incrementing the state at unrealistically long distances.
 - Simple example: Gaspari-Cohn localization, a function applied to the background covariance matrix which smoothly tapers long-distance correlations to zero
 - “Long-distance” defined in terms of a localization radius (can be fixed)



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 - “Long-distance” defined in terms of a localization radius (can be fixed)
 - **Adaptive localization: allow the localization radius to change for each observation, based on that observation’s expected influence**

Adaptive localization

- **Hypothesis:** for a given single observation, the optimal localization length scale is proportional to the reduction of ensemble variance in observation space
 - In other words, the more a given observation would reduce the analysis covariance, the longer localization length it is given.
- Define ρ as the reduction of ensemble covariance in observation space:

$$\rho = \mathbf{HP}^a\mathbf{H}^T / \mathbf{HP}^b\mathbf{H}^T = \mathbf{R} / (\mathbf{HP}^b\mathbf{H}^T + \mathbf{R})$$

where \mathbf{H} is the linearized observation operator, \mathbf{P}^a is the analysis ensemble covariance, \mathbf{P}^b is the background ensemble covariance, and \mathbf{R} is the observation error covariance

- Small $\rho \leftrightarrow$ large reduction in variance \leftrightarrow useful observation \Rightarrow larger optimal localization radius
- Conversely, ρ close to 1 implies a smaller optimal localization radius

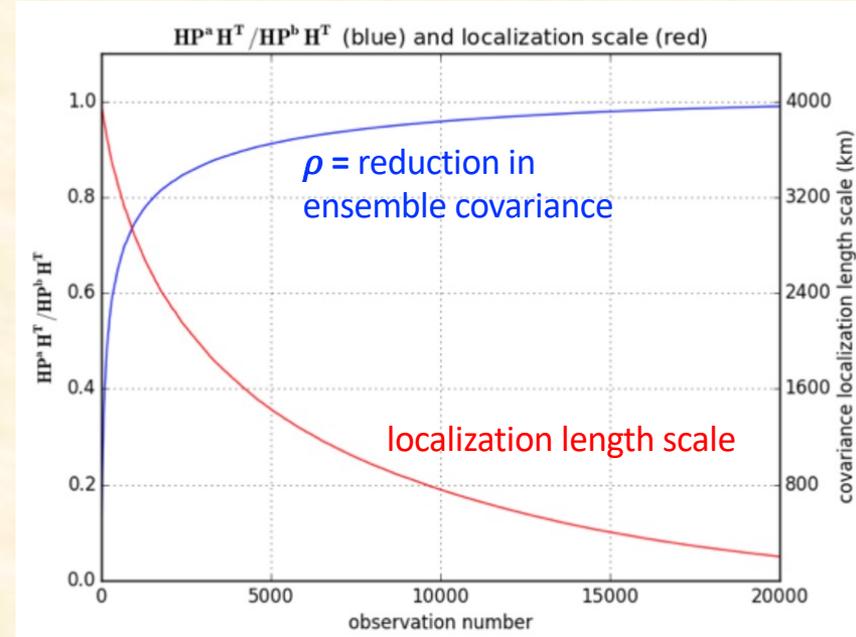
Adaptive localization

Empirically define

$$L = L_0(1 - e^{-(1-\rho)/r})$$

where

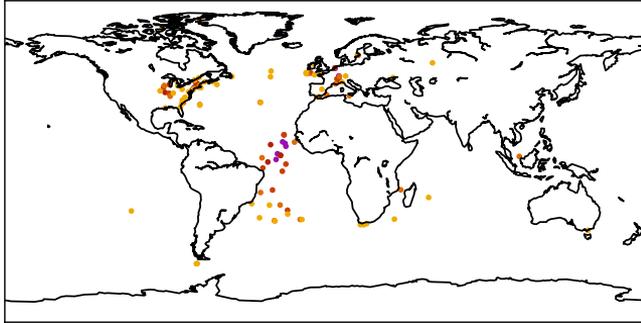
- L is the localization length scale
- L_0 is the maximum allowed localization length scale
- ρ is the reduction of ensemble variance in ob space (prev. slide); $\rho \in (0,1]$
- r is a parameter governing how tight the relationship between ρ and L is



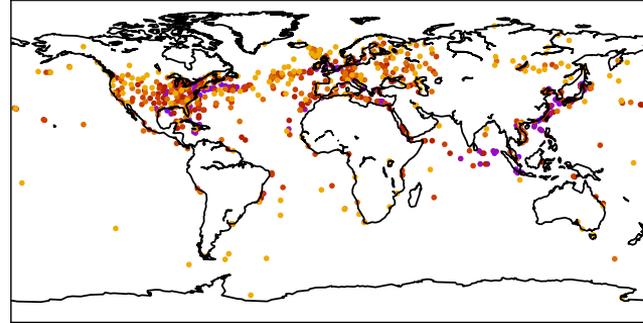
Observations are assimilated serially in order of increasing ρ , and ρ is recomputed after each ob is assimilated.

Adaptive localization

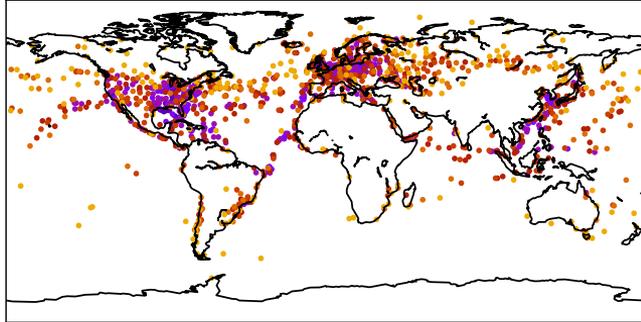
(a) 1854



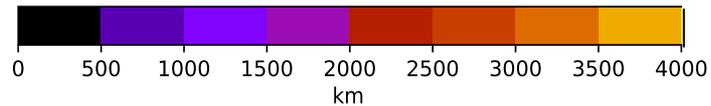
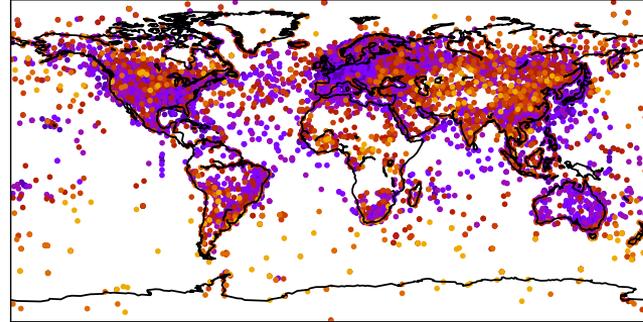
(b) 1915



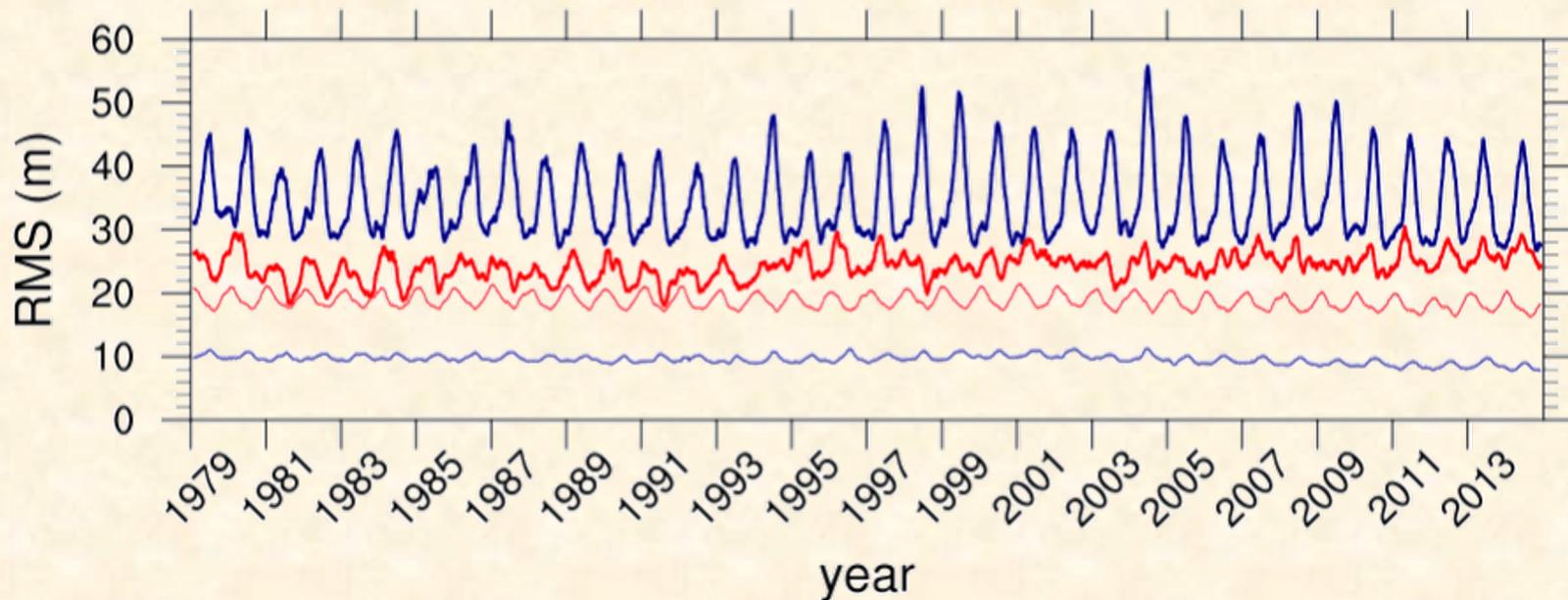
(c) 1935



(d) 2000

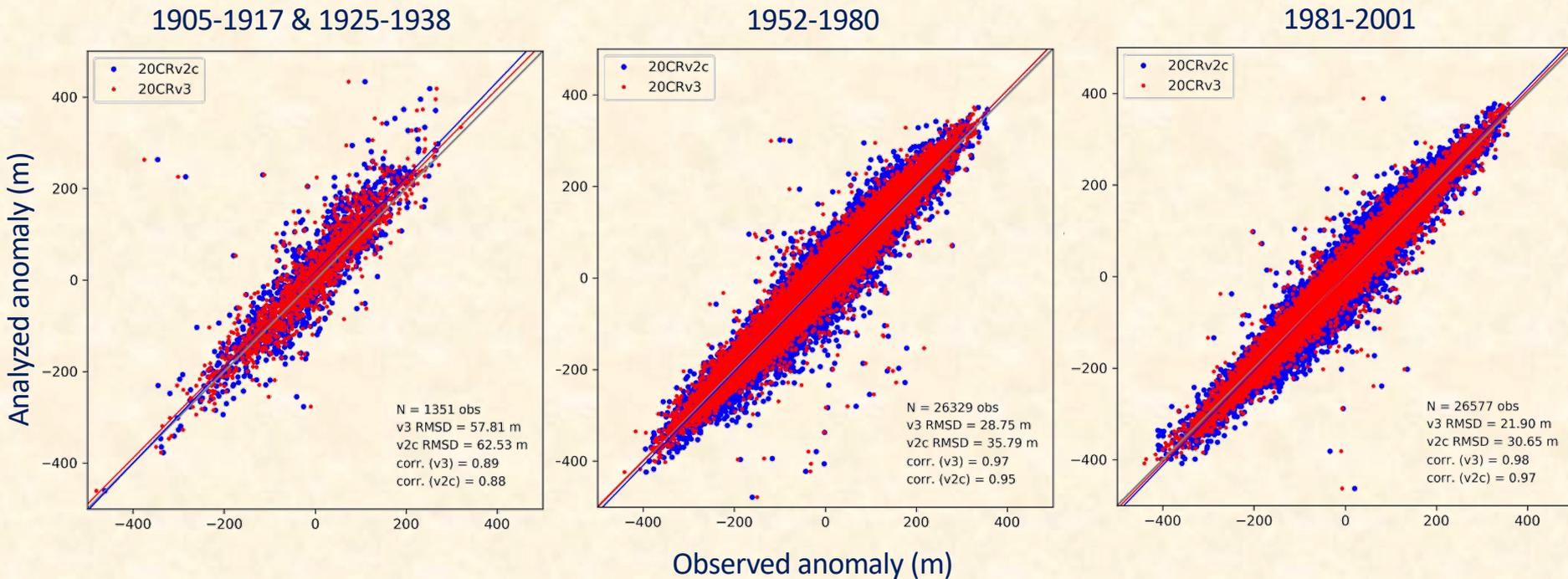


500hPa Geopotential Height RMS relative to ERA5, 30N90N



20CRv2c avg RMS = 34.4424
20CRv2c avg spread = 9.56379
20CRv3 avg RMS = 24.2104
20CRv3 avg spread = 18.7837

500hPa geopotential height analyzed anomalies versus observed anomalies from upper-air measurements at Lindenberg, Germany



Summary

- The 20th Century Reanalysis version 3 reconstructs nearly 200 years of sub-daily, global weather history by assimilating *only* surface pressure observations
- Reliable confidence and uncertainty estimates provided by 80 ensemble members in EnKF
- Updated inflation algorithm (relaxation-to-prior-spread) yields more consistent estimates of uncertainty
- Adaptively-varying localization makes observation-thinning unnecessary; more observations can be assimilated
- Overall, the 20CRv3 system includes many improvements over the previous 20CRv2c, leading to improved performance of the dataset (Slivinski et al 2019)

References

- 20th Century Reanalysis:
 - Homepage: <https://go.usa.gov/XTd>
 - Slivinski, L.C., G.P. Compo, J.S. Whitaker, P.D. Sardeshmukh, et. al. (2019) Towards a more reliable historical reanalysis: Improvements for version 3 of the Twentieth Century Reanalysis system. *Quarterly Journal of the Royal Meteorological Society*, **145**: 2876– 2908
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