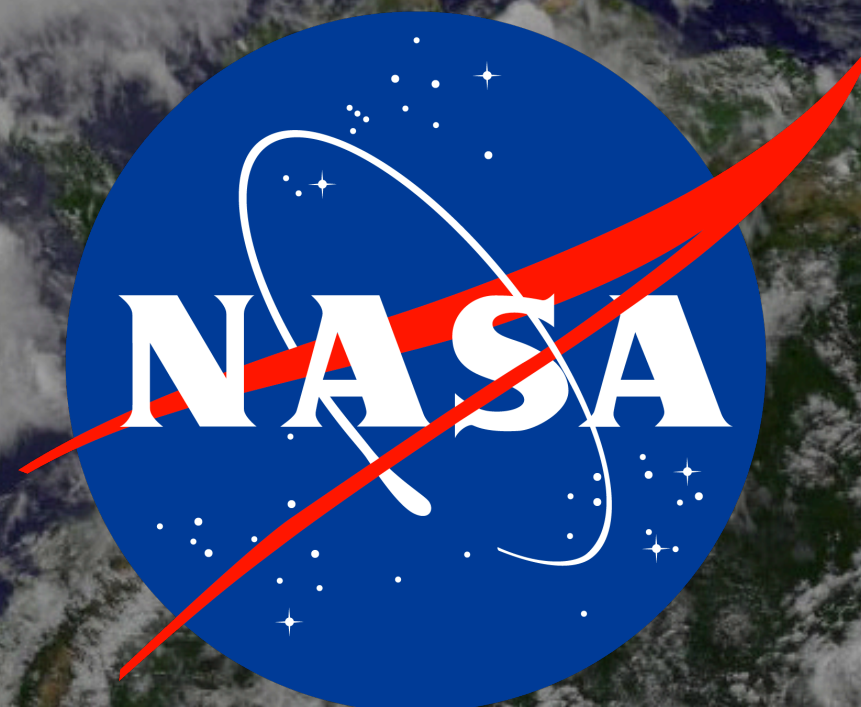


Identifying dominant atmospheric drivers of ocean variability using ECCO and the MITgcm adjoint: Implications for reducing model bias

Dan Amrhein, Dafydd Stephenson
National Center for Atmospheric Research

LuAnne Thompson, Noah Rosenberg
University of Washington



What are the **dominant patterns and pathways** by which the atmosphere drives **ocean variability**?

Outline

The ocean as an integrator (and source) of random variation

What dominates ocean variability? Views from the atmosphere and ocean

Reconciliation: A dynamics-weighted principal components analysis

Dominant atmospheric drivers of decadal AMOC variability

Impacts of dominant drivers over the observational period

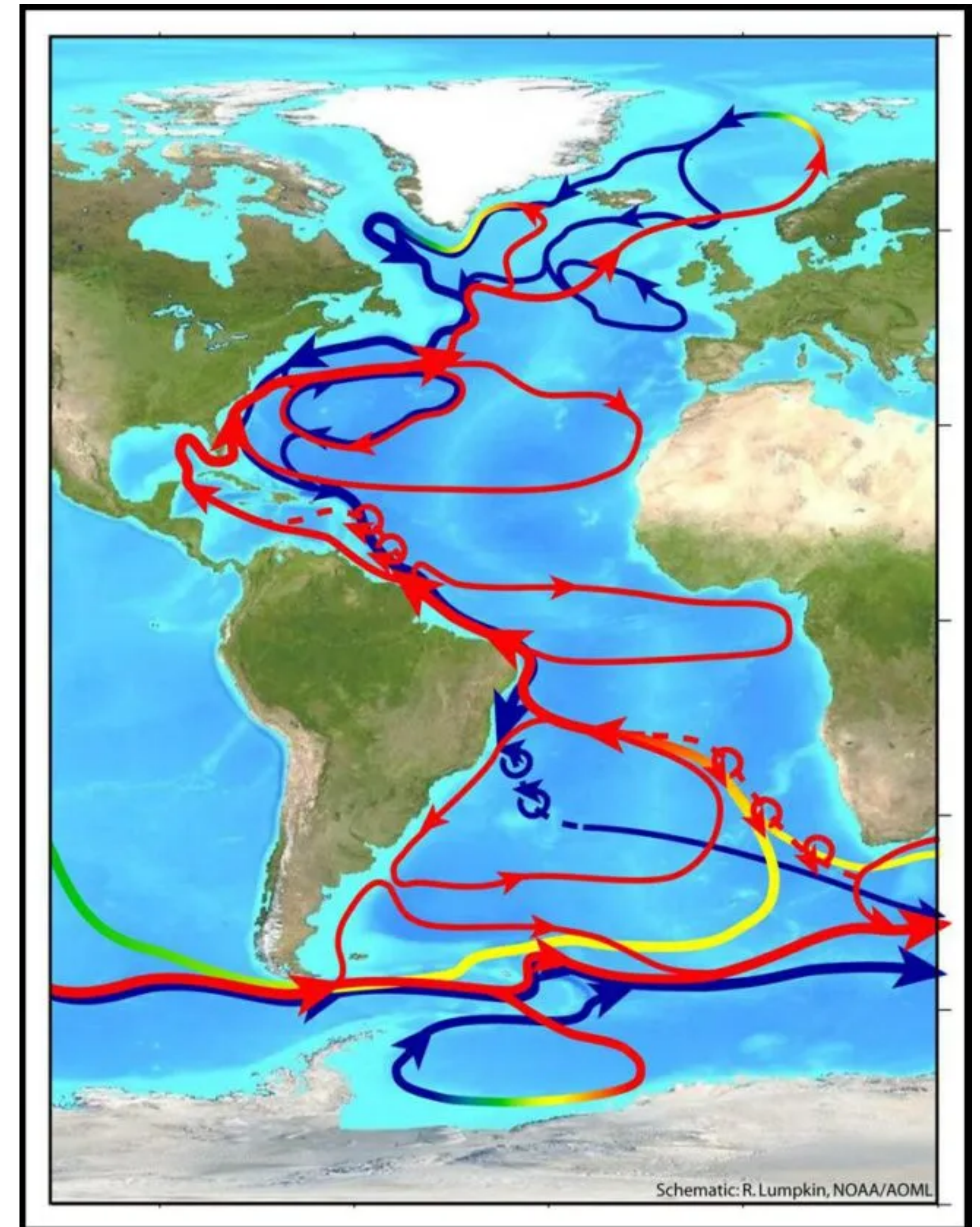
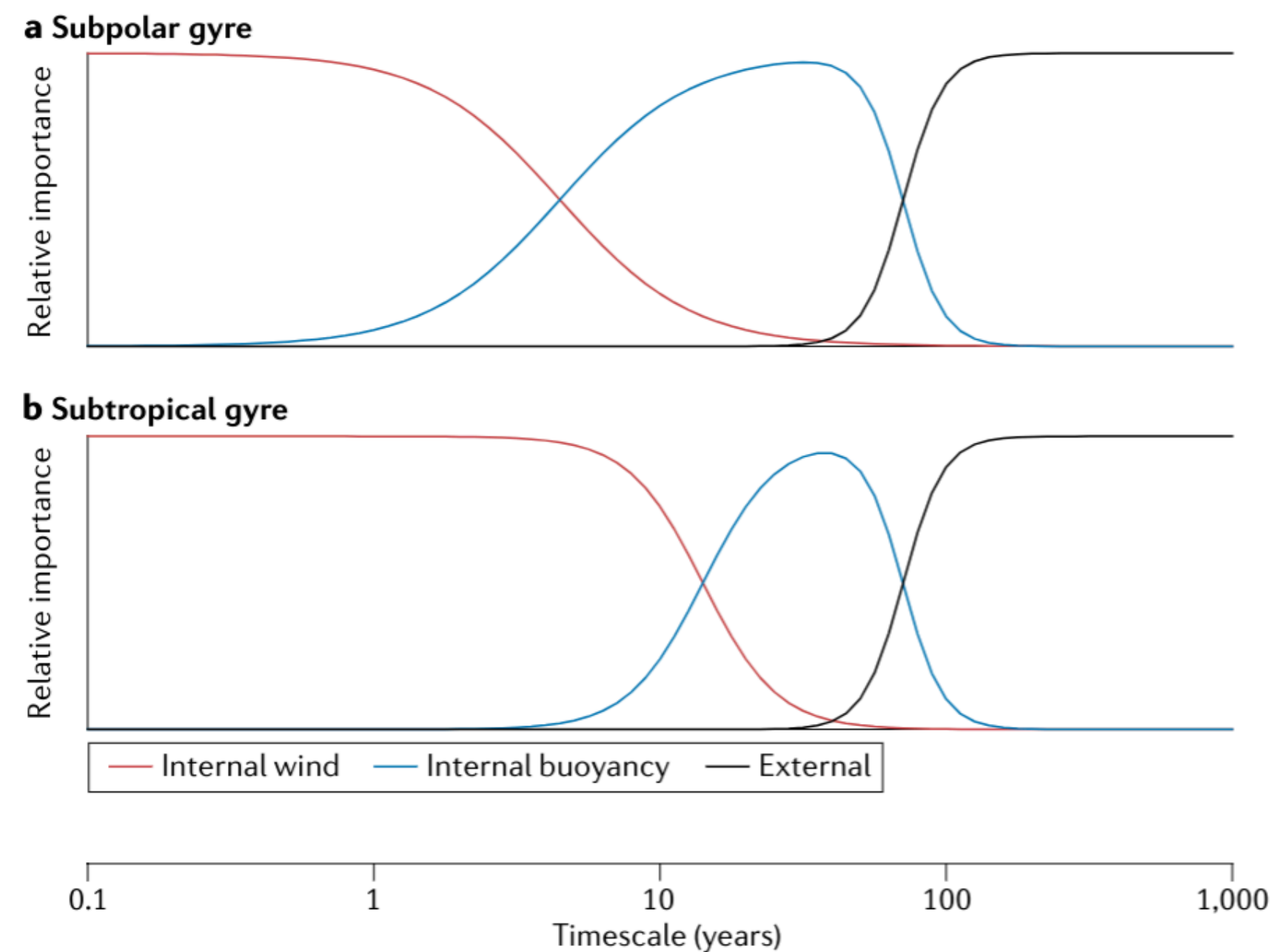
Tracing pathways of decadal AMOC change from atmosphere to ocean

Case study: Variability in the Atlantic Meridional Overturning Circulation

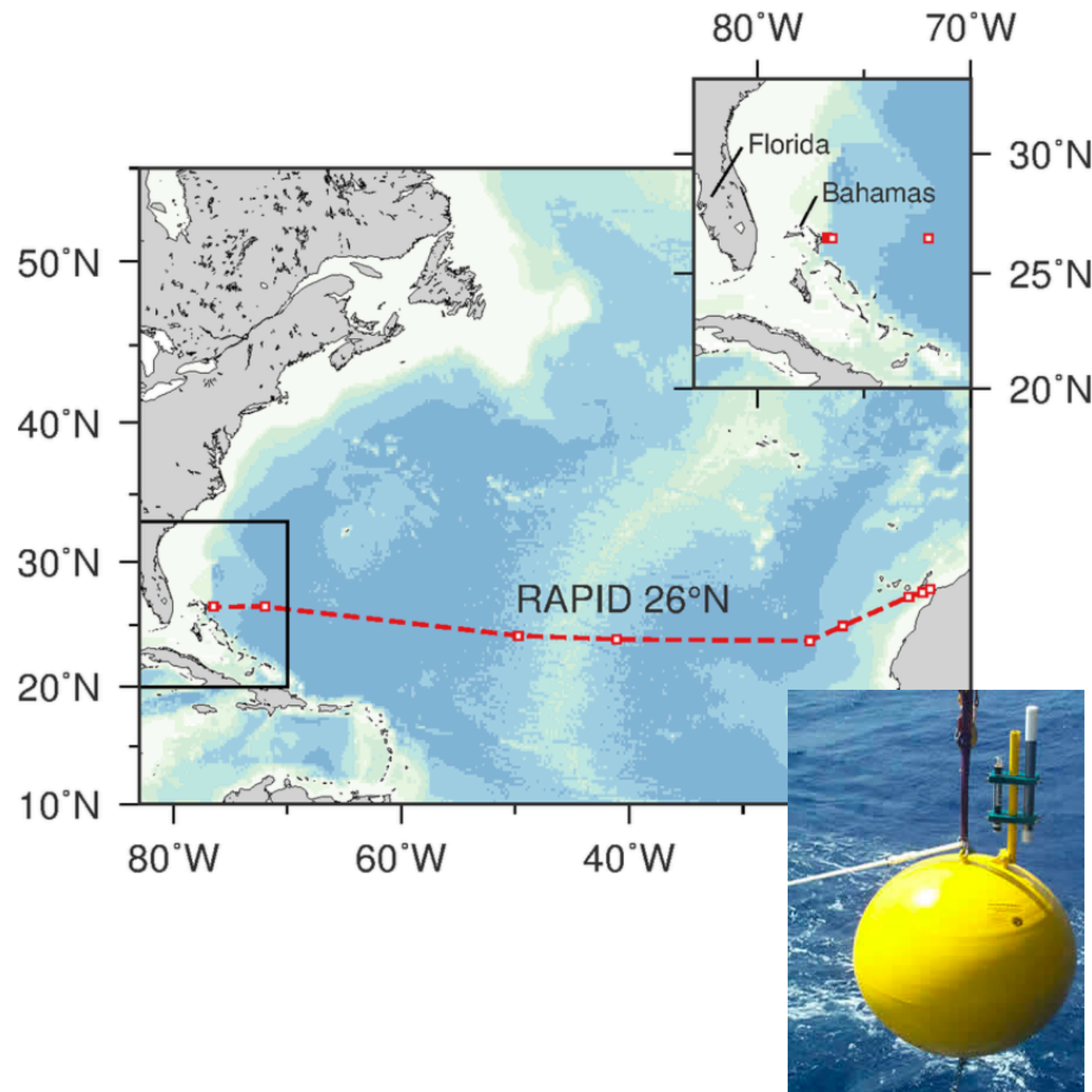
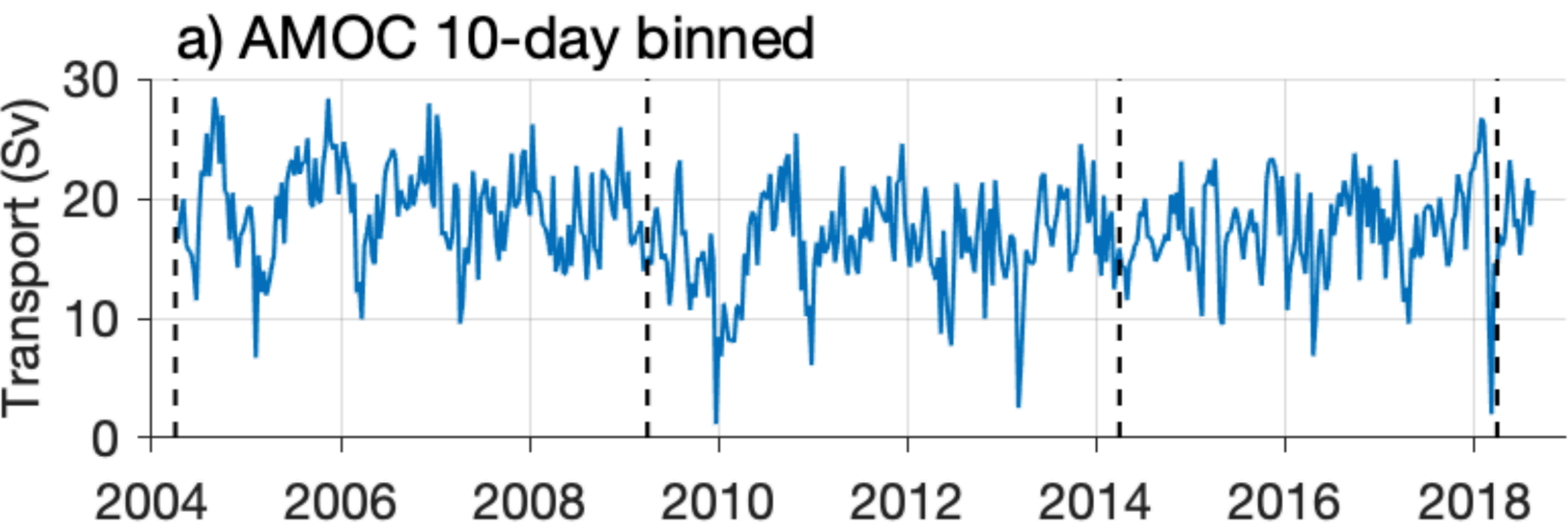
Decadal variability in AMOC influences climate variability

Can mask anthropogenic warming signal

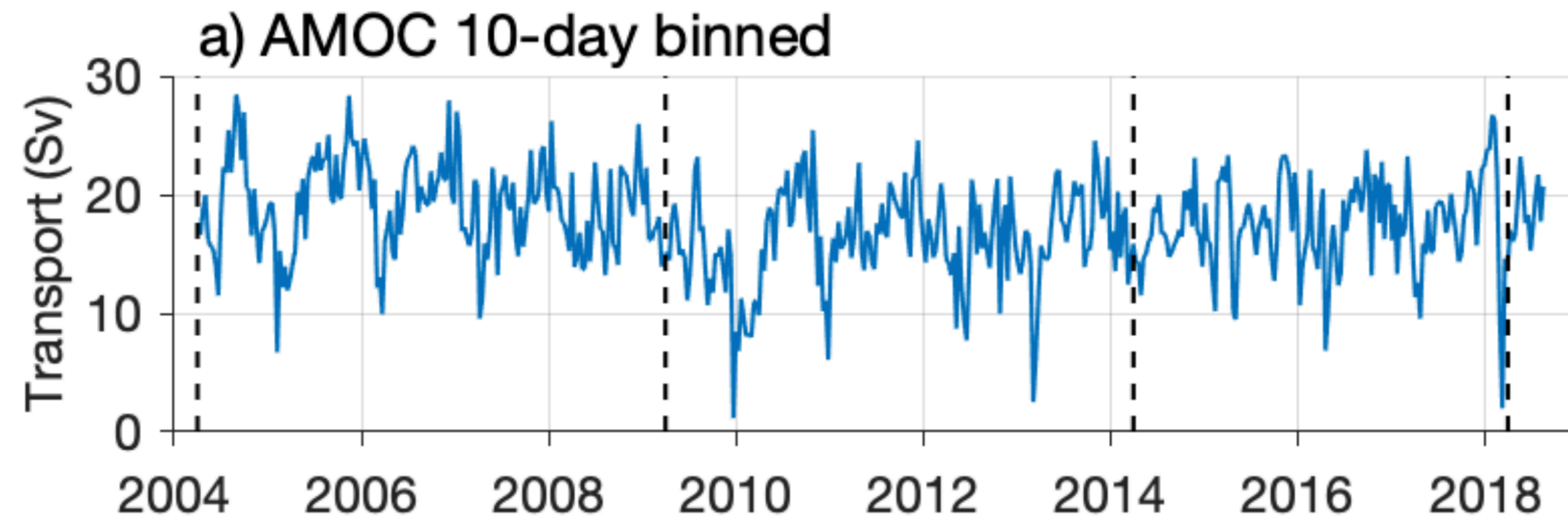
Junction between high-frequency (e.g. wind) and low-frequency (e.g. buoyancy) influences



Observing AMOC variability



Representing AMOC variability as a random process



$$X_{\text{amoc}} = X_{\tau} + X_b + \dots$$

$$\text{var}(X_{\text{amoc}}) = \text{var}(X_{\tau}) + \text{var}(X_b) + 2\text{cov}(X_b, X_{\tau}) + \dots$$

Representing AMOC variability as a random process

The zero-order result here is that a modern ocean–ice GCM, when least squares fit to the 2-decade-long global datasets available since 1992, produces a dynamically consistent estimate of the Atlantic MOC, one which is indistinguishable from a stationary Gaussian red-noise process. With the benefit of hindsight, the result is unsurprising: a system with long memory is subject to continuous small stochastic disturbances by external processes (winds, precipitation, etc.), which themselves

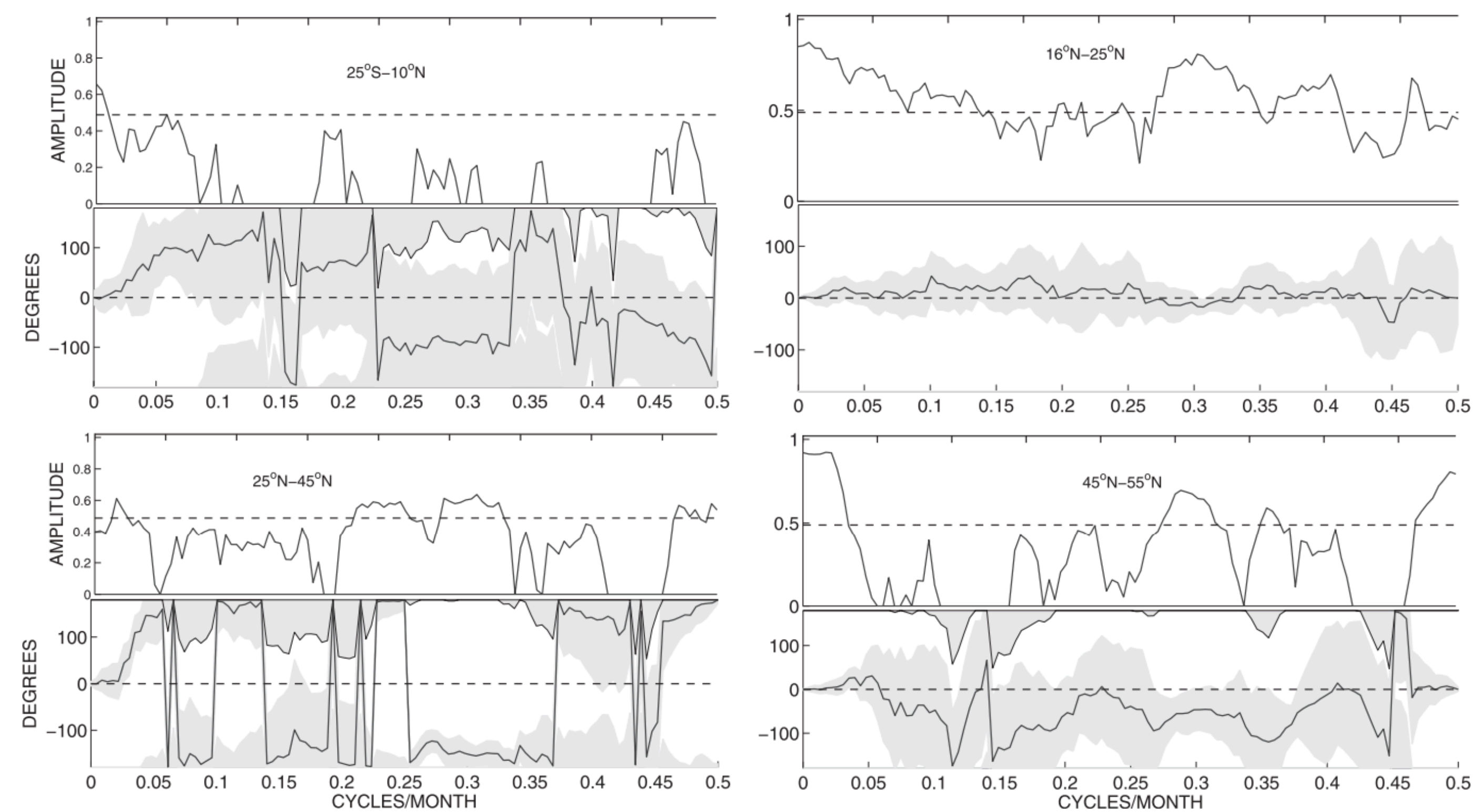


FIG. 9. Coherence magnitude and phase of the AMOC between four pairs of nearest-neighbor latitudes. The

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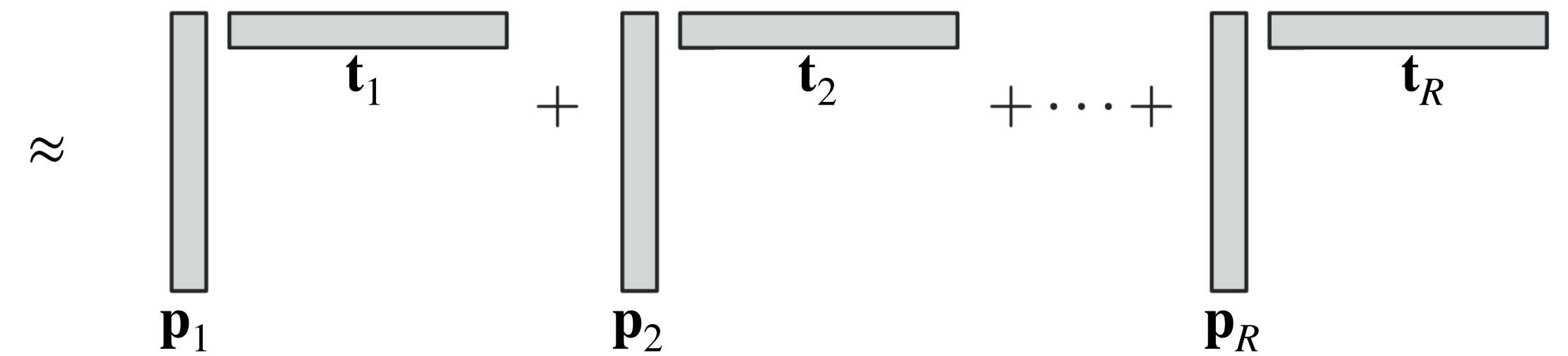
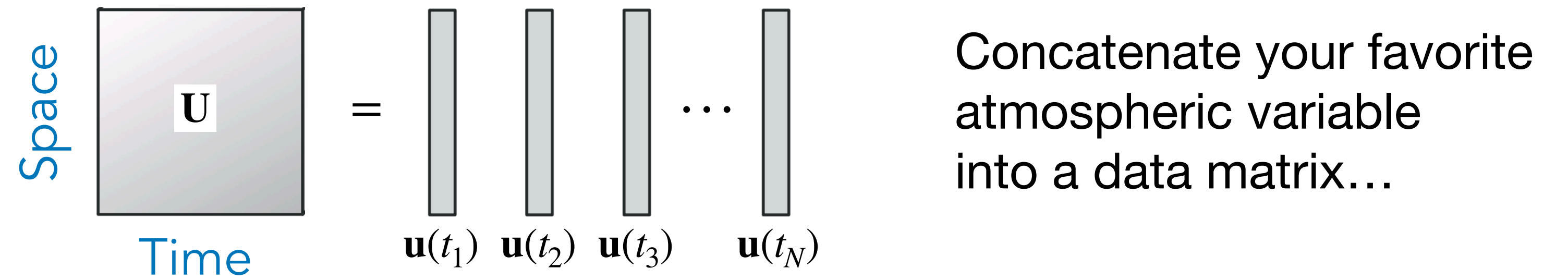
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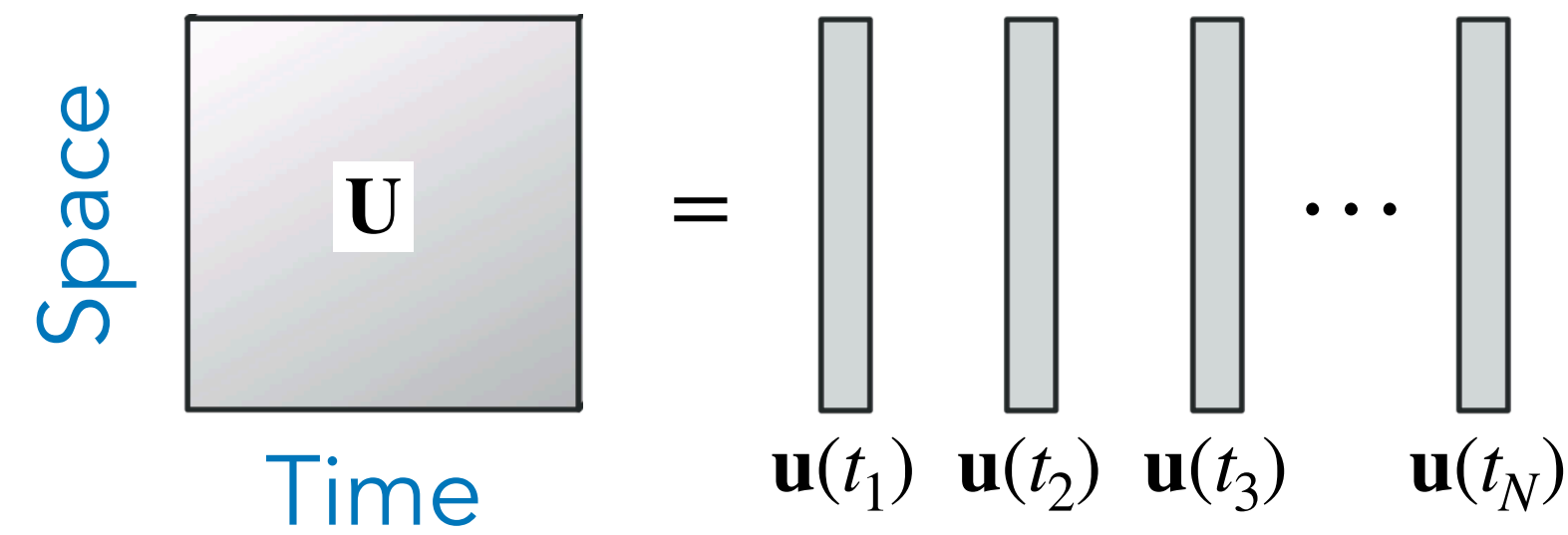
The leading **EOF**
answers the question:

**What atmospheric
pattern accounts for
the greatest fraction
of total atmospheric
variance?**

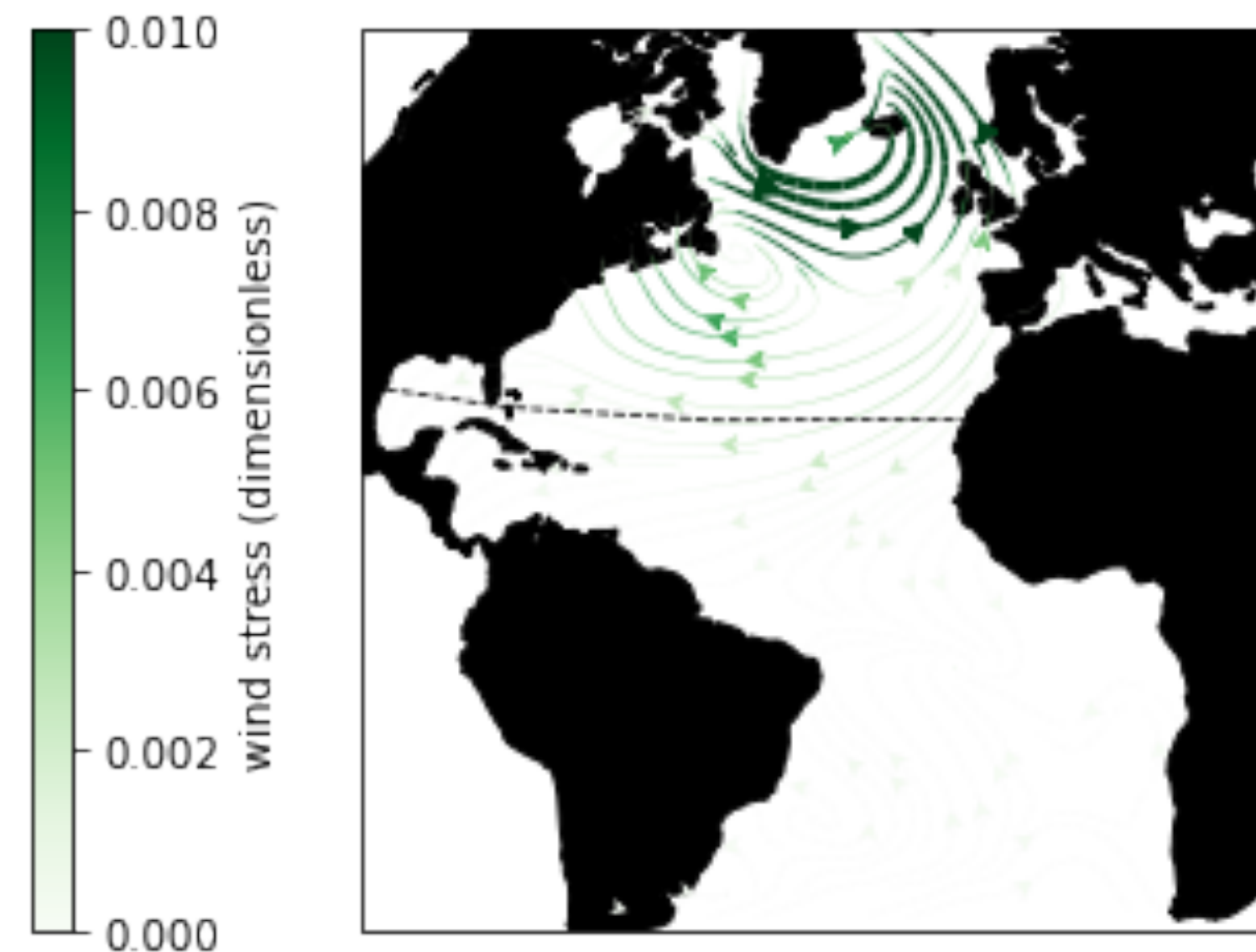
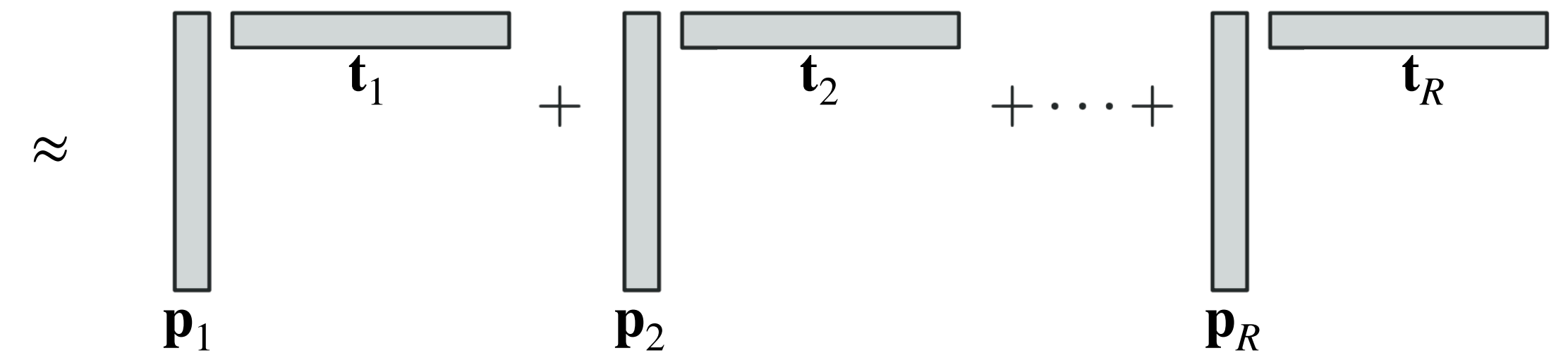


The leading **EOF**
answers the question:

**What atmospheric
pattern accounts for
the greatest fraction
of total atmospheric
variance?**



Concatenate your favorite
atmospheric variable
into a data matrix...



The leading EOF of
wind stress in the ECCO
v4r4 state estimate.

An ocean perspective:

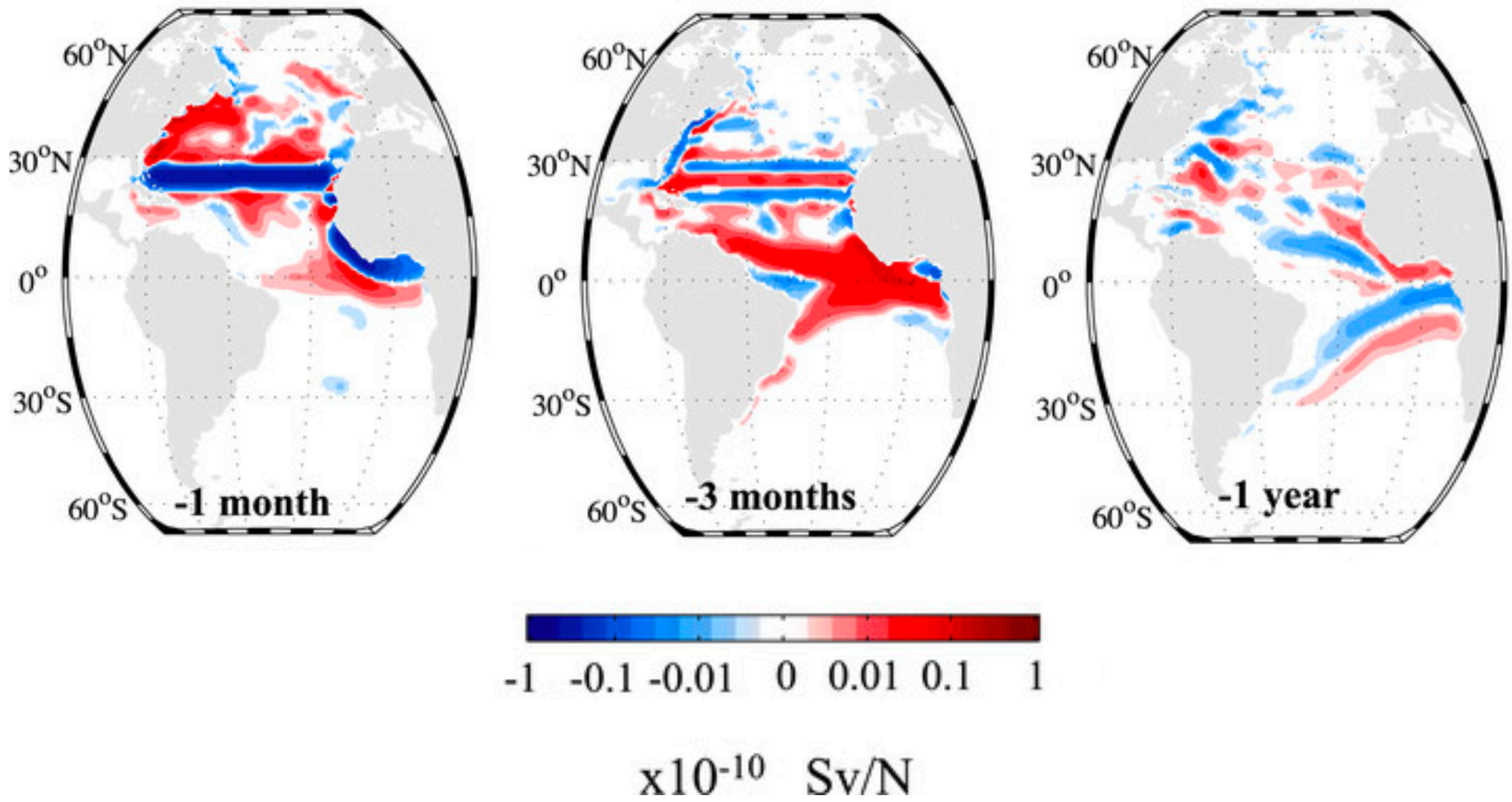
What atmospheric
pattern would **most
efficiently excite
ocean variability?**

Ocean model adjoint sensitivities diagnose dominant drivers

$$\mathbf{s} = \frac{\partial x}{\partial \mathbf{u}}$$

AMOC strength
@ 26N in January

Zonal wind stress



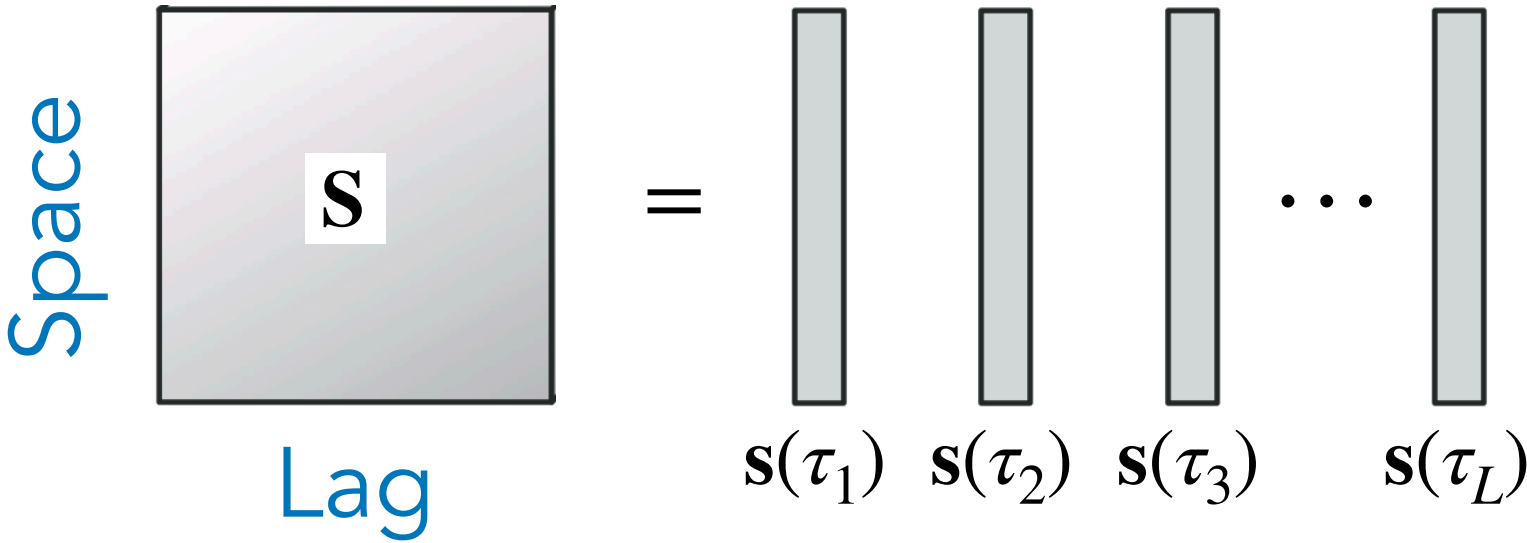
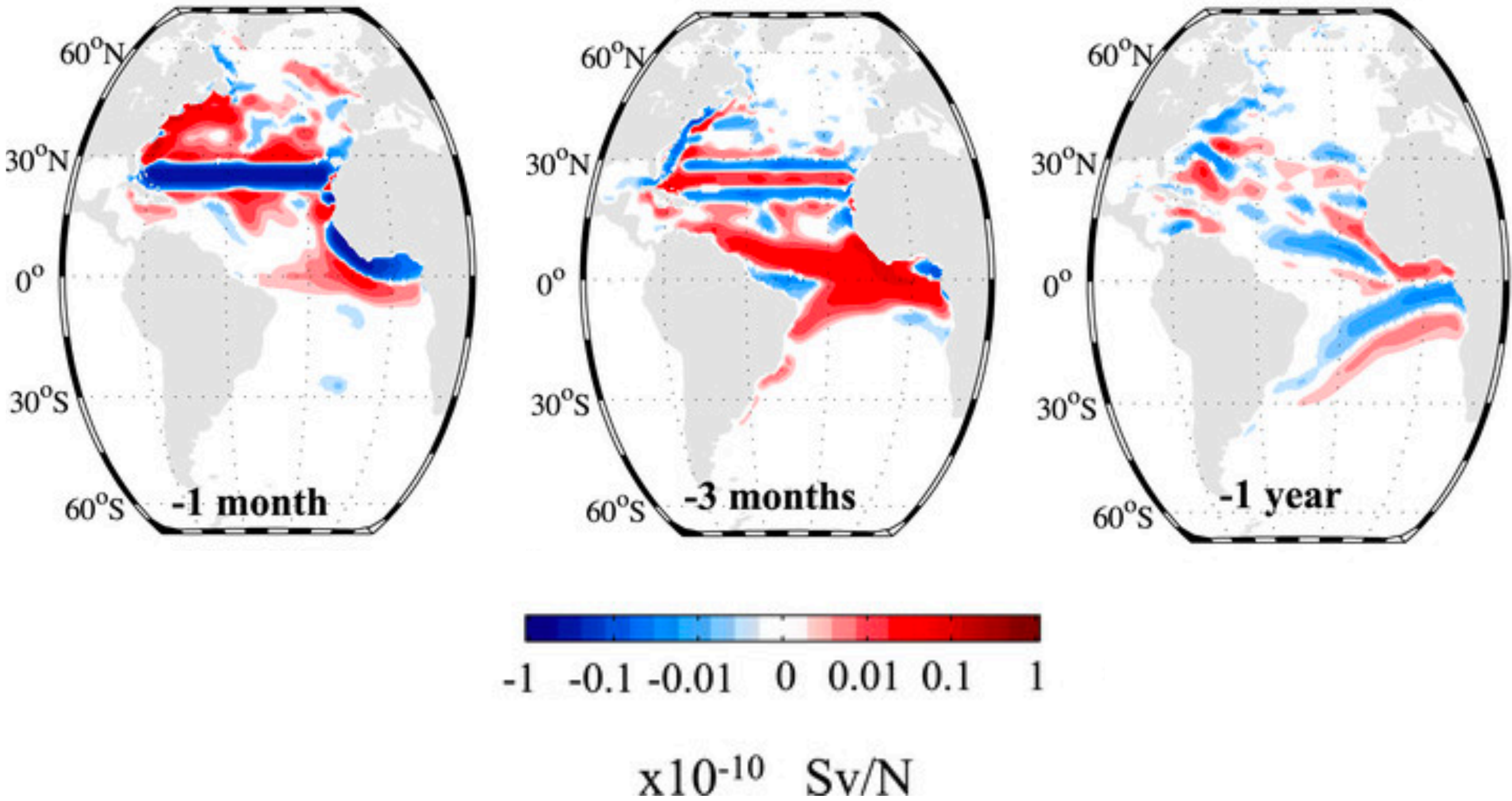
Pillar et al. 2016
Also Heimbach and Wunsch 2011; Jones et al. 2018;
Kostov et al. 2019, 2021; Fukumori et al. 2021; Stephenson
and Sevellec 2020, 2021

Ocean model adjoint sensitivities diagnose dominant drivers

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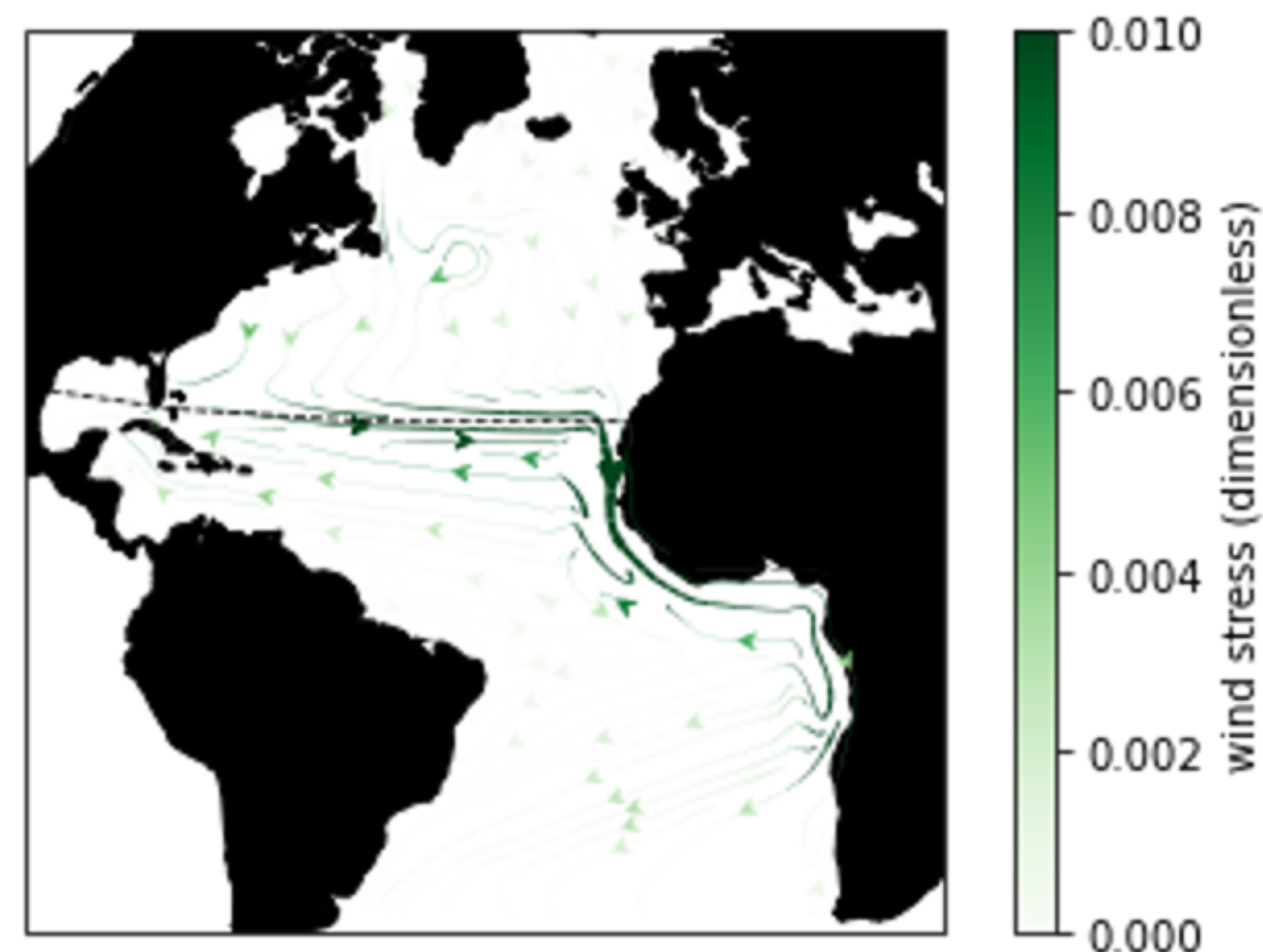
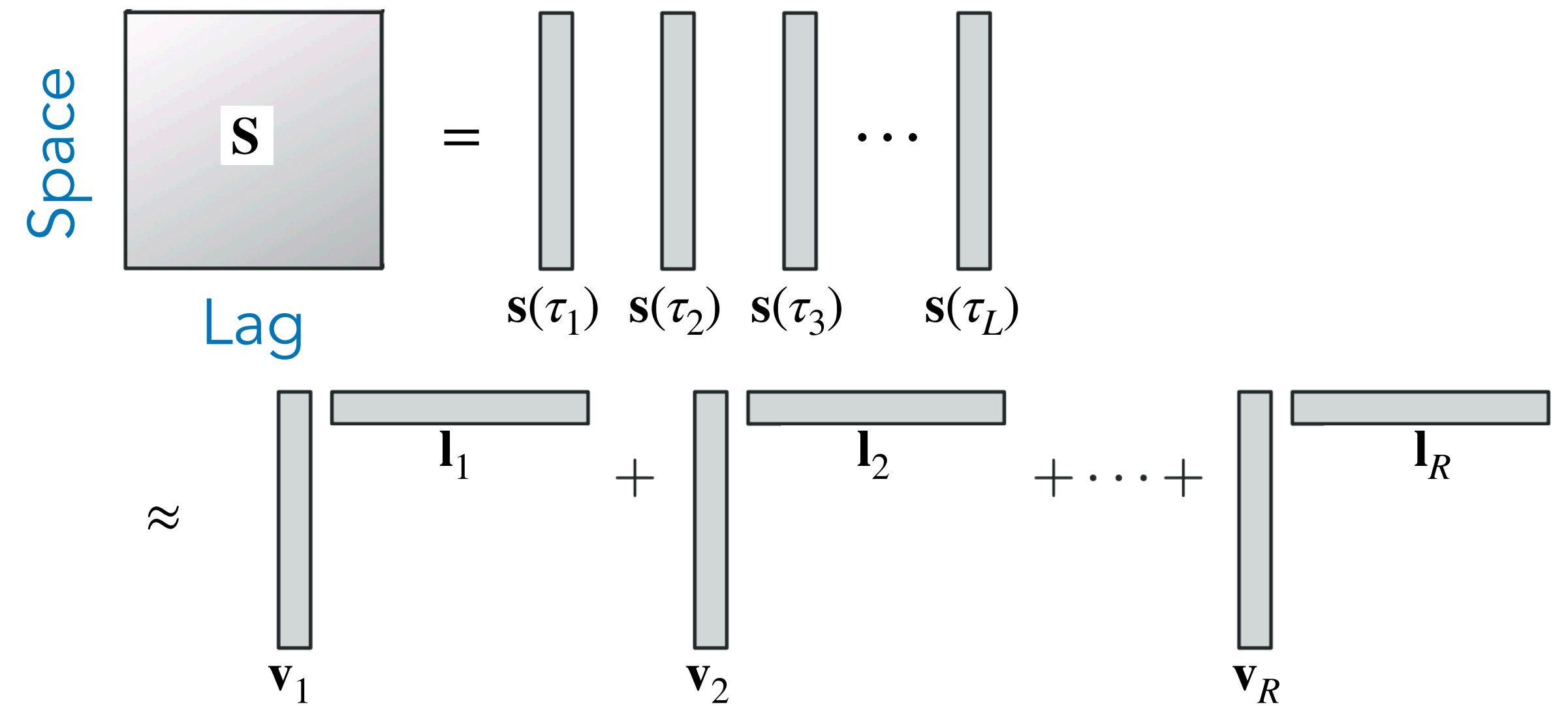


Pillar et al. 2016
Also Heimbach and Wunsch 2011; Jones et al. 2018;
Kostov et al. 2019, 2021; Fukumori et al. 2021; Stephenson
and Sevellec 2020, 2021

What atmospheric pattern would **most efficiently excite ocean variability**?

The leading **stochastic optimal**.

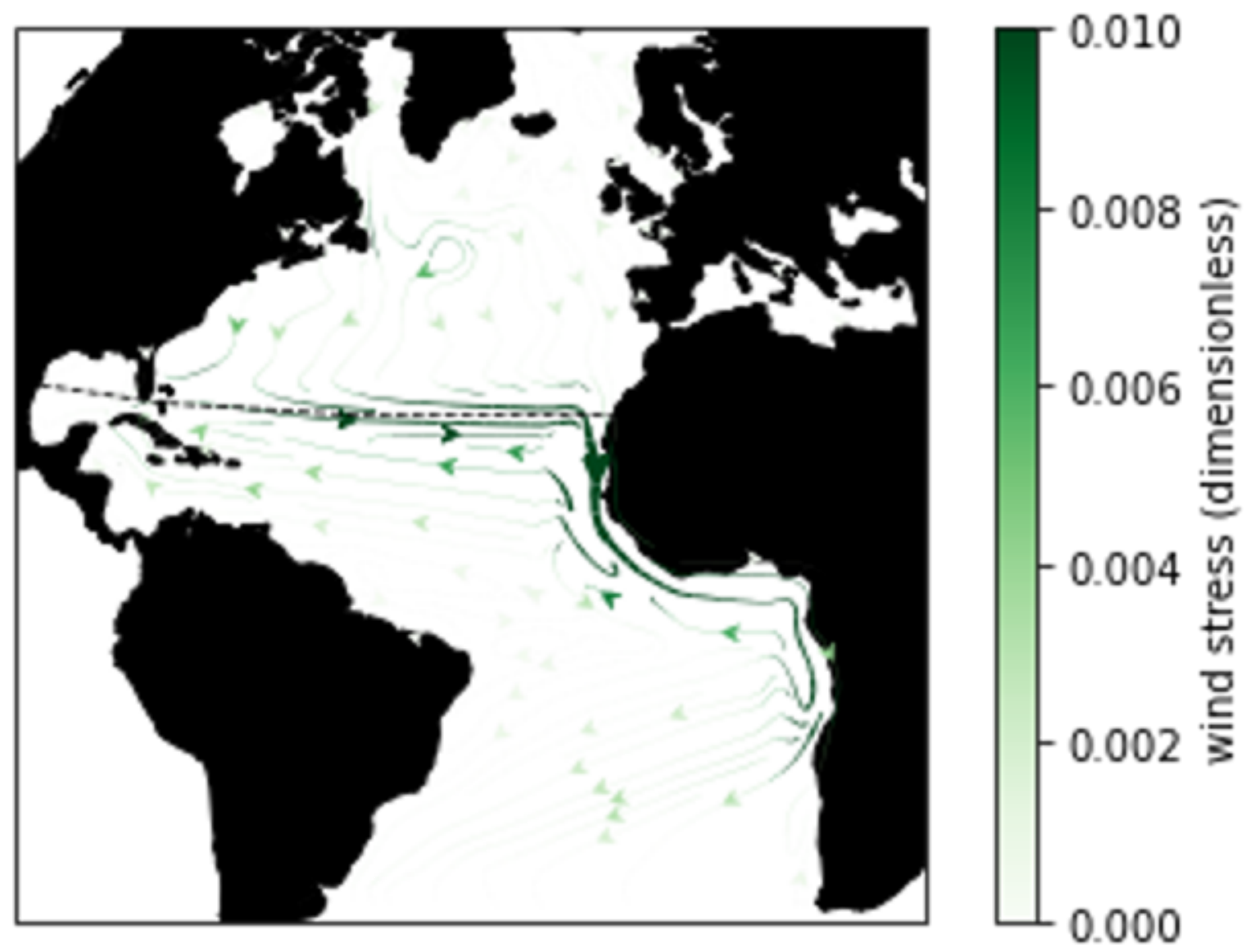
*Farrell and Ioannou
1993, 1996; Kleeman
and Moore, 1997*



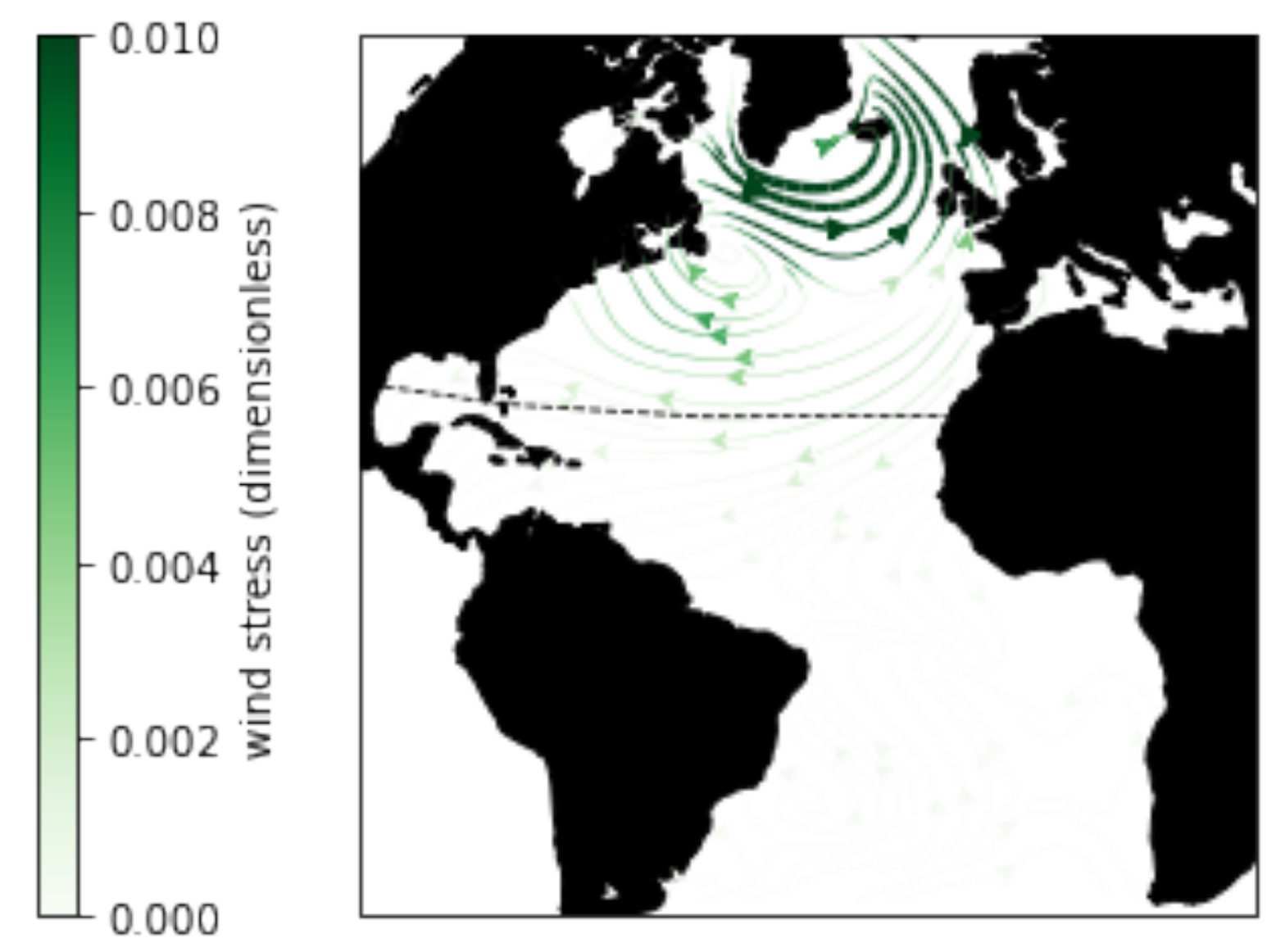
\mathbf{v}_1 for wind stress in ECCO v4r4

An interpretive quandary

What the ocean “wants”

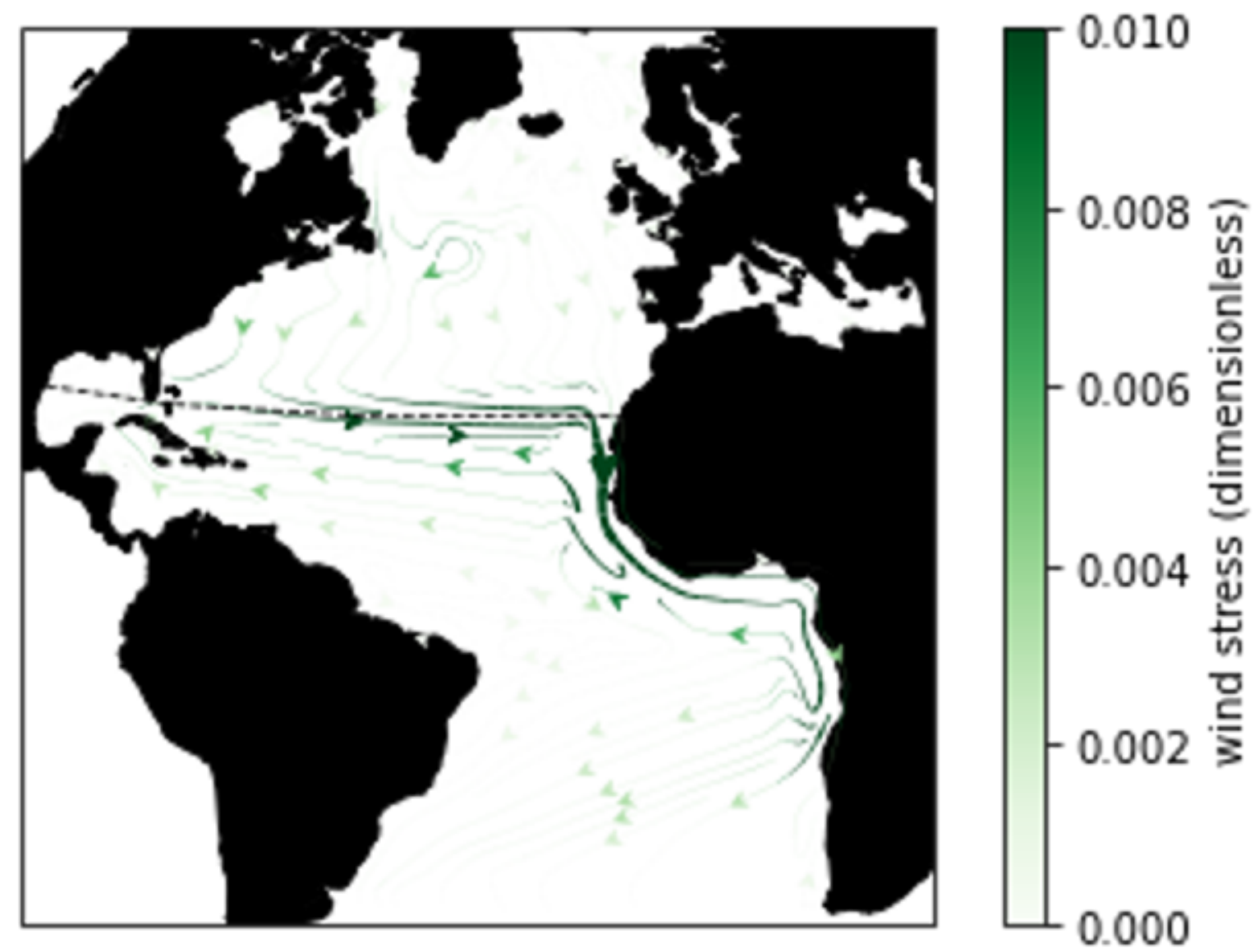


What the ocean “gets”

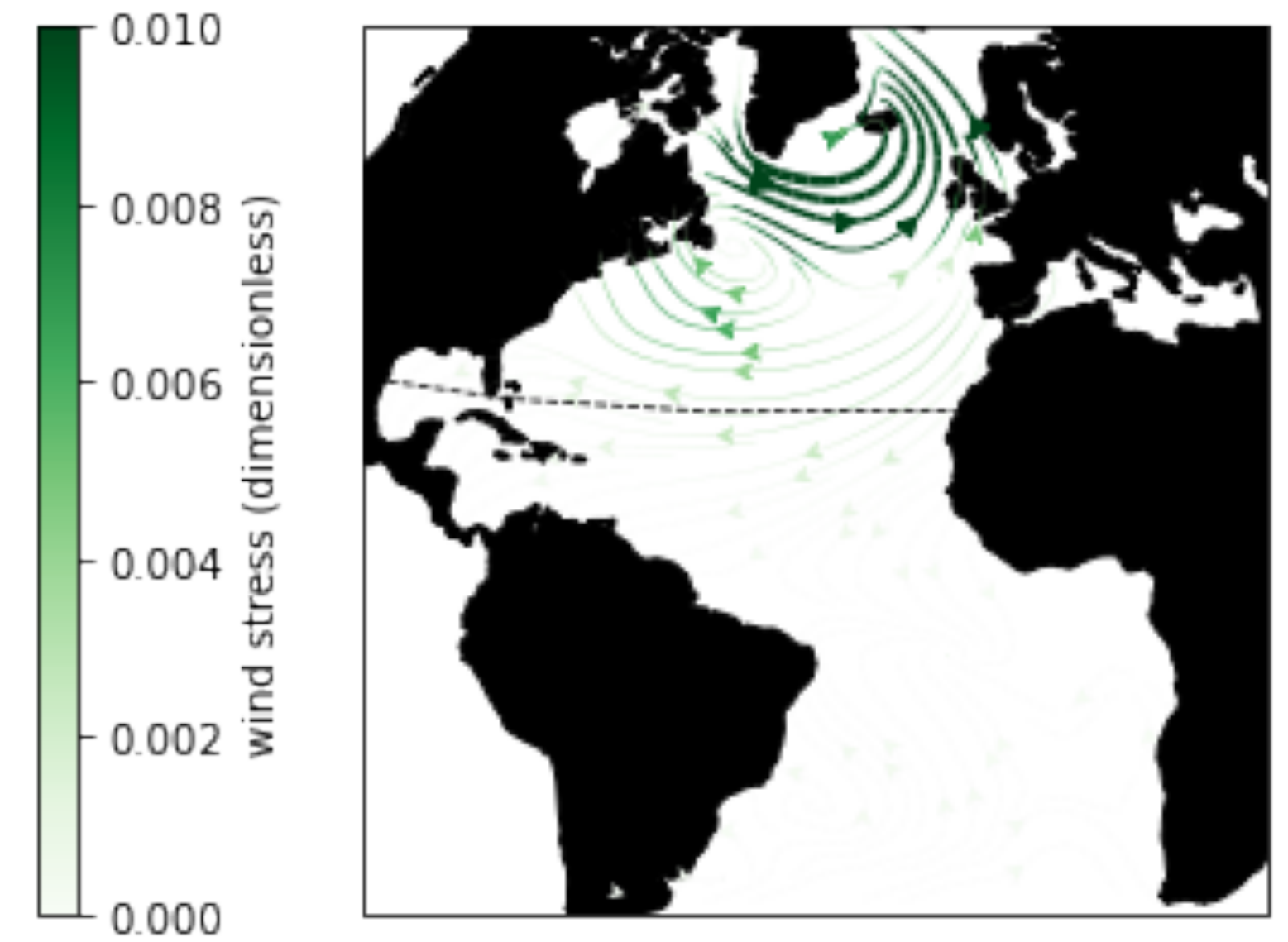


An interpretive quandary

What the ocean “wants”

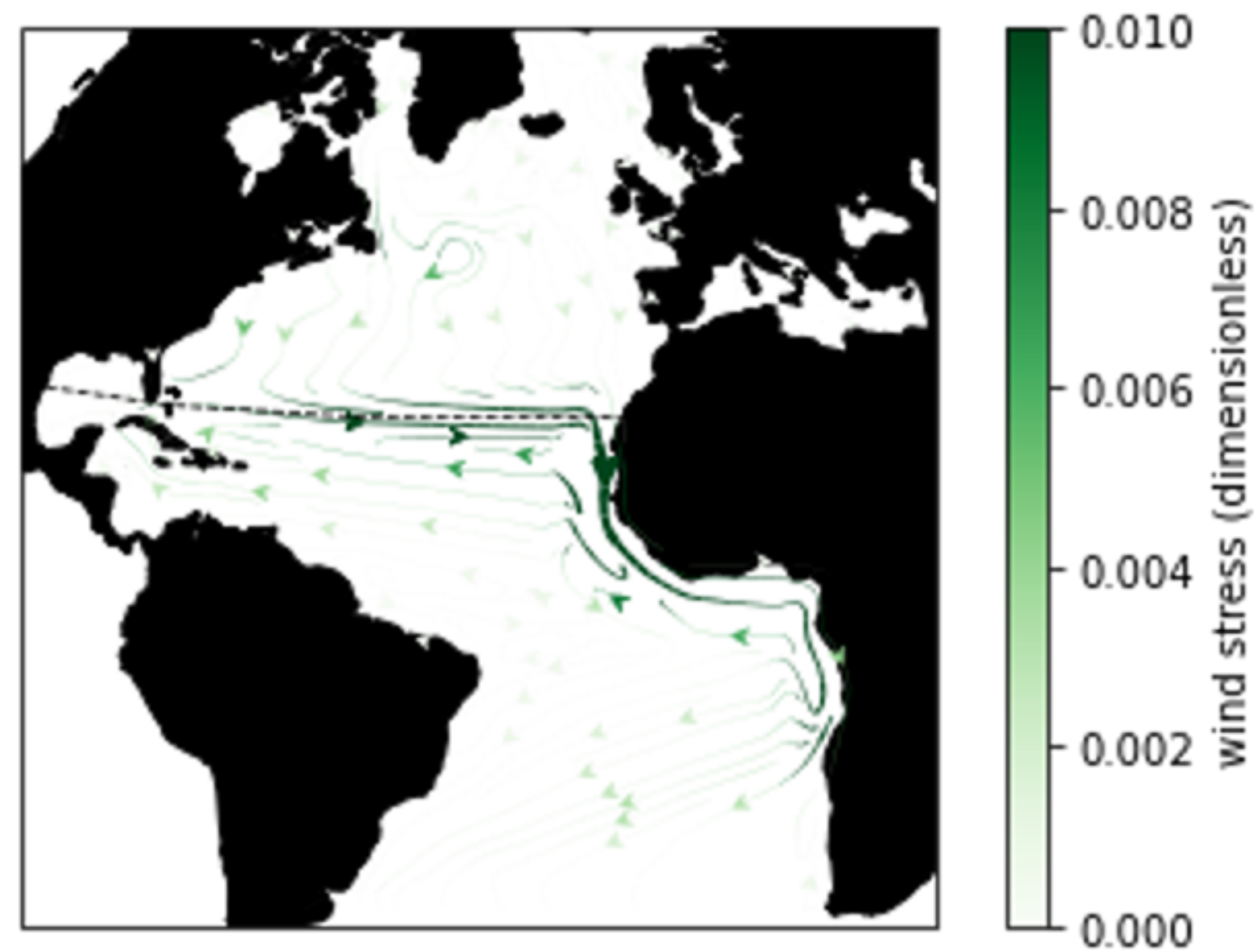


What the ocean “gets”

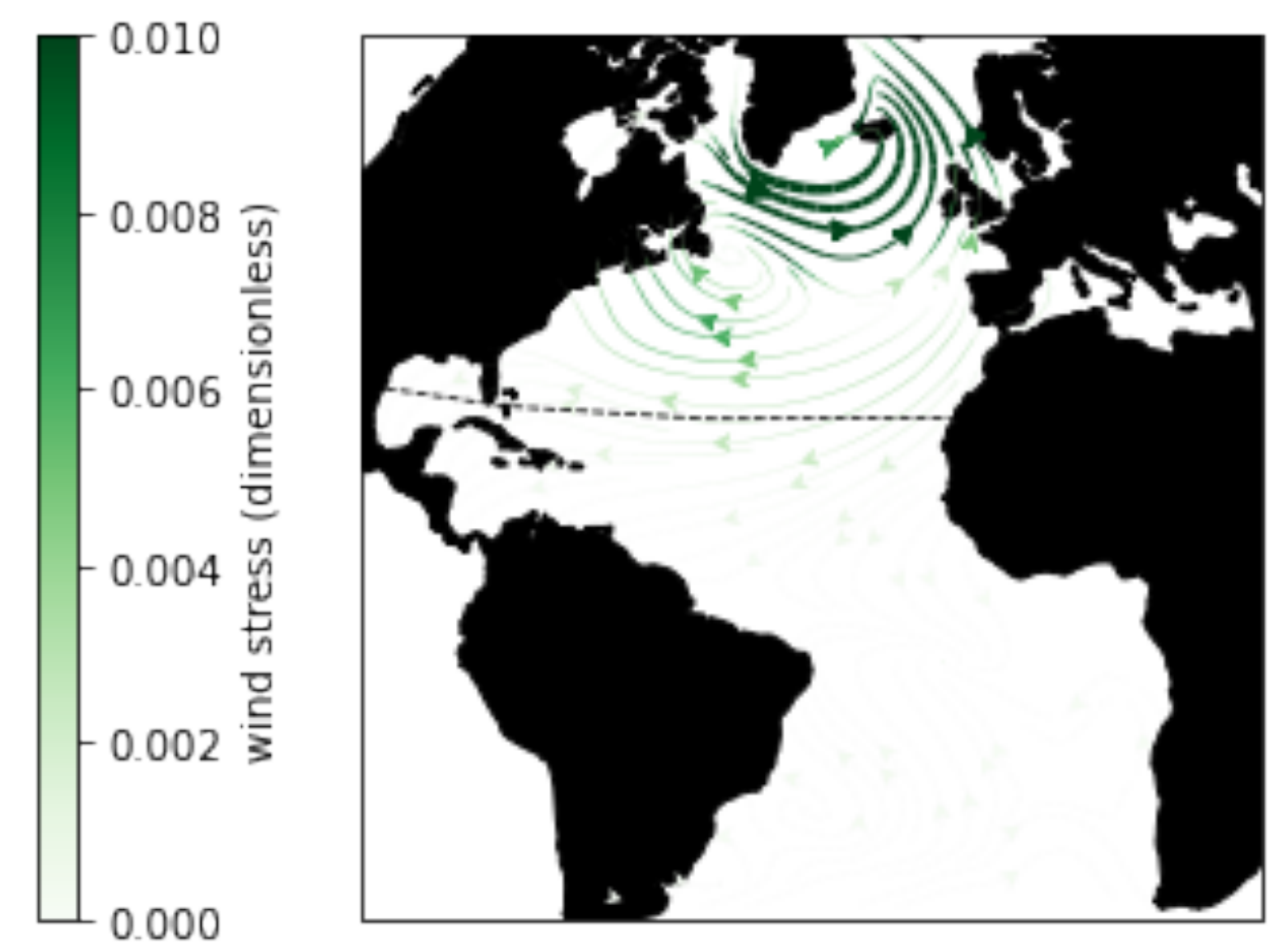


An interpretive quandary

What the ocean “wants”



What the ocean “gets”



Is the leading EOF the **leading driver** of variability in this ocean quantity?

Is the leading stochastic optimal really the most important **mechanism** for changing the ocean?

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The math slide! Deriving dynamics-weighted principal components

$$\mathbf{s} = \frac{\partial x}{\partial \mathbf{u}}$$

Definition of adjoint sensitivity

The math slide! Deriving dynamics-weighted principal components

$$\mathbf{s} = \frac{\partial x}{\partial \mathbf{u}}$$

Definition of adjoint sensitivity

$$\delta x(t) \approx \sum_{i=1}^{N_\tau} \mathbf{s}(\tau_i)^\top \delta \mathbf{u}(t - \tau_i)$$

To change AMOC strength, we
can make a control change $\delta \mathbf{u}$

The math slide! Deriving dynamics-weighted principal components

$$\mathbf{s} = \frac{\partial x}{\partial \mathbf{u}}$$

Definition of adjoint sensitivity

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To change AMOC strength, we can make a control change $\delta \mathbf{u}$

$$\sigma_\Sigma^2 = \left\langle \left(\delta x(t) \right)^2 \right\rangle$$

The variance of the quantity of interest

The math slide! Deriving dynamics-weighted principal components

$$\mathbf{s} = \frac{\partial x}{\partial \mathbf{u}}$$

Definition of adjoint sensitivity

$$\delta x(t) \approx \sum_{i=1}^{N_\tau} \mathbf{s}(\tau_i)^\top \delta \mathbf{u}(t - \tau_i)$$

To change AMOC strength, we can make a control change $\delta \mathbf{u}$

$$\sigma_\Sigma^2 = \left\langle \left(\delta x(t) \right)^2 \right\rangle$$

The variance of the quantity of interest

$$= \sum_{i=1}^{N_\tau} \sum_{j=1}^{N_\tau} \mathbf{s}(\tau_i)^\top \left\langle \delta \mathbf{u}(t - \tau_i) \delta \mathbf{u}^\top(t - \tau_j) \right\rangle \mathbf{s}(\tau_j)$$

Substitution gets a bit sticky...

$$= \mathbf{tr}(\mathbf{C}\mathbf{Z})$$

$$\mathbf{Z} = \mathbf{S}\mathbf{S}^\top$$

Atmospheric
spatial covariance

...but is simplified by two assumptions (see also Kleeman and Moore, 1997):

1. Flux covariances are separable in space and time
2. Sensitivities are stationary

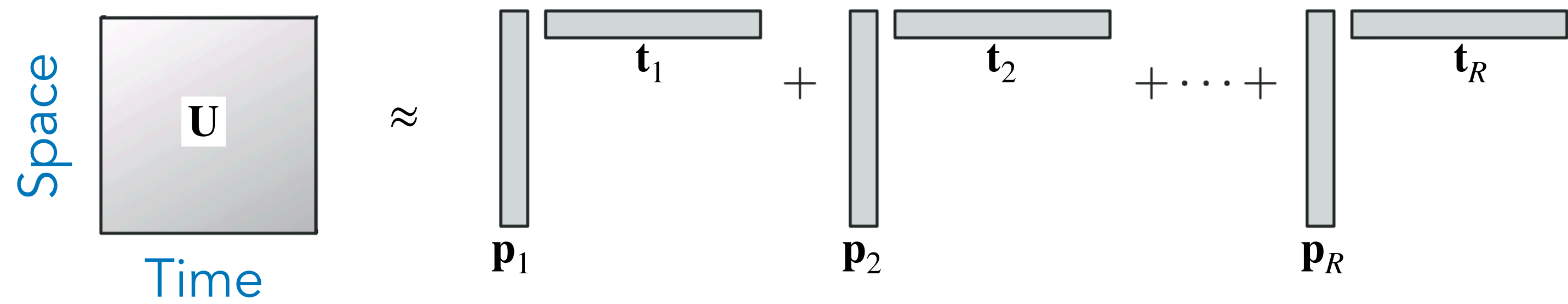
The math slide! Deriving dynamics-weighted principal components

$$\sigma_{\Sigma}^2 = \text{tr}(\mathbf{C}\mathbf{Z})$$

$$\mathbf{U} = \sum \mathbf{p}_k \mathbf{t}_k^{\top}$$

$$\sigma_{\Sigma}^2 = \sum \sigma_k^2$$

- Our requirements:
- 1. An EOF-like decomposition
 - 2. Contributions to ocean variance that add (no cross terms)



The math slide! Deriving dynamics-weighted principal components

$$\sigma_{\Sigma}^2 = \text{tr}(\mathbf{C}\mathbf{Z})$$

$$\mathbf{U} = \sum \mathbf{p}_k \mathbf{t}_k^{\top}$$

$$\sigma_{\Sigma}^2 = \sum \sigma_k^2$$

$$\mathbf{S}^{\top} \mathbf{U} = \mathbf{L} \mathbf{\Gamma} \mathbf{T}^{\top}$$

Contributions to Qol variance

$$\mathbf{P} = \mathbf{U} \mathbf{T}$$

↑
Spatial patterns ranked
by their contribution to
ocean Qol variance

Our requirements:

1. An EOF-like decomposition
2. Contributions to ocean variance that add (no cross terms)

...yields an **SVD optimization problem!**

Amounts to computing principal components weighted by adjoint sensitivities.

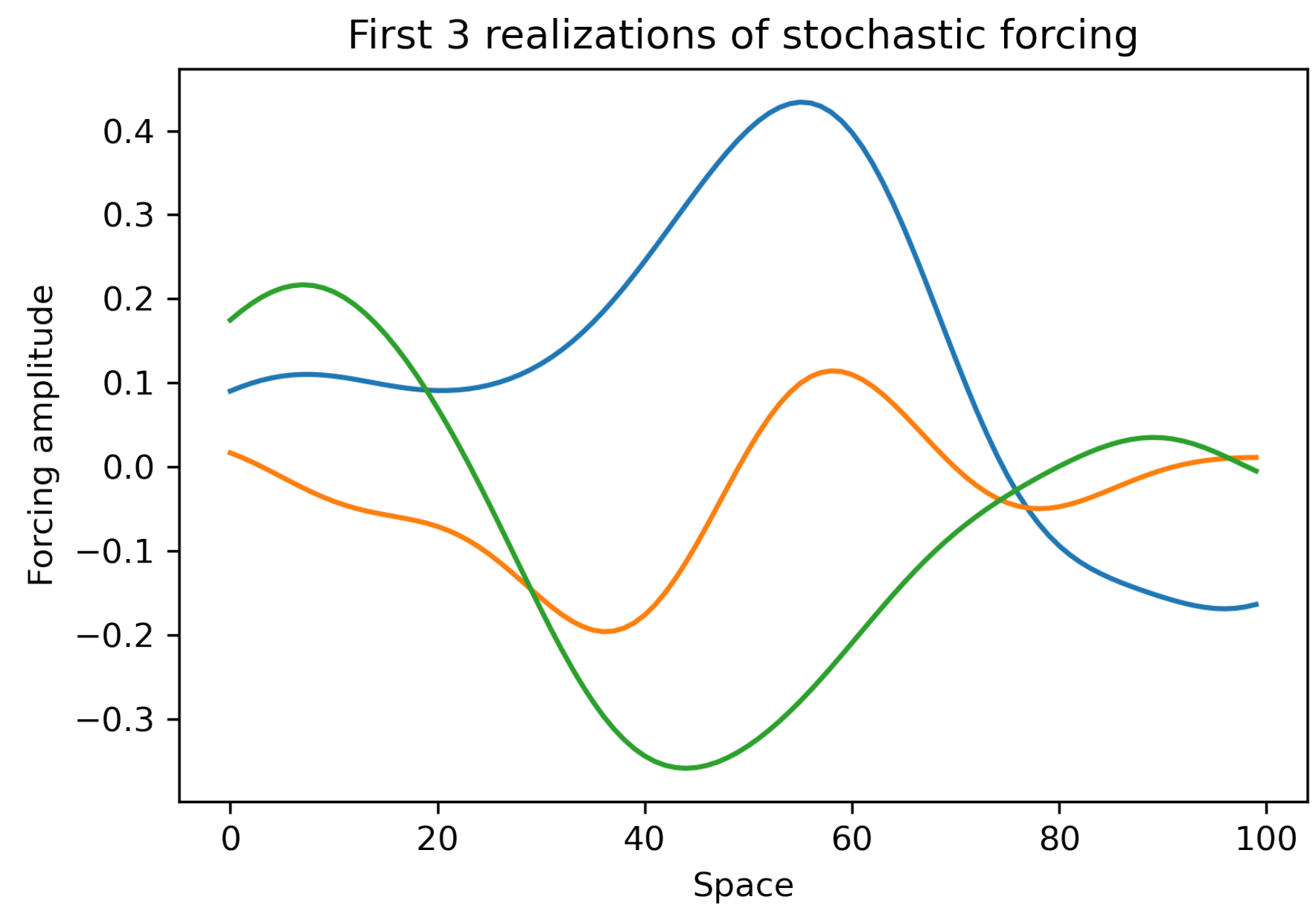
EOF-like, but singular values are ocean Qol variance rather than atmospheric variance.

Patterns are orthogonal in time, but not space.

Recovers EOFs and SOs for limit cases.

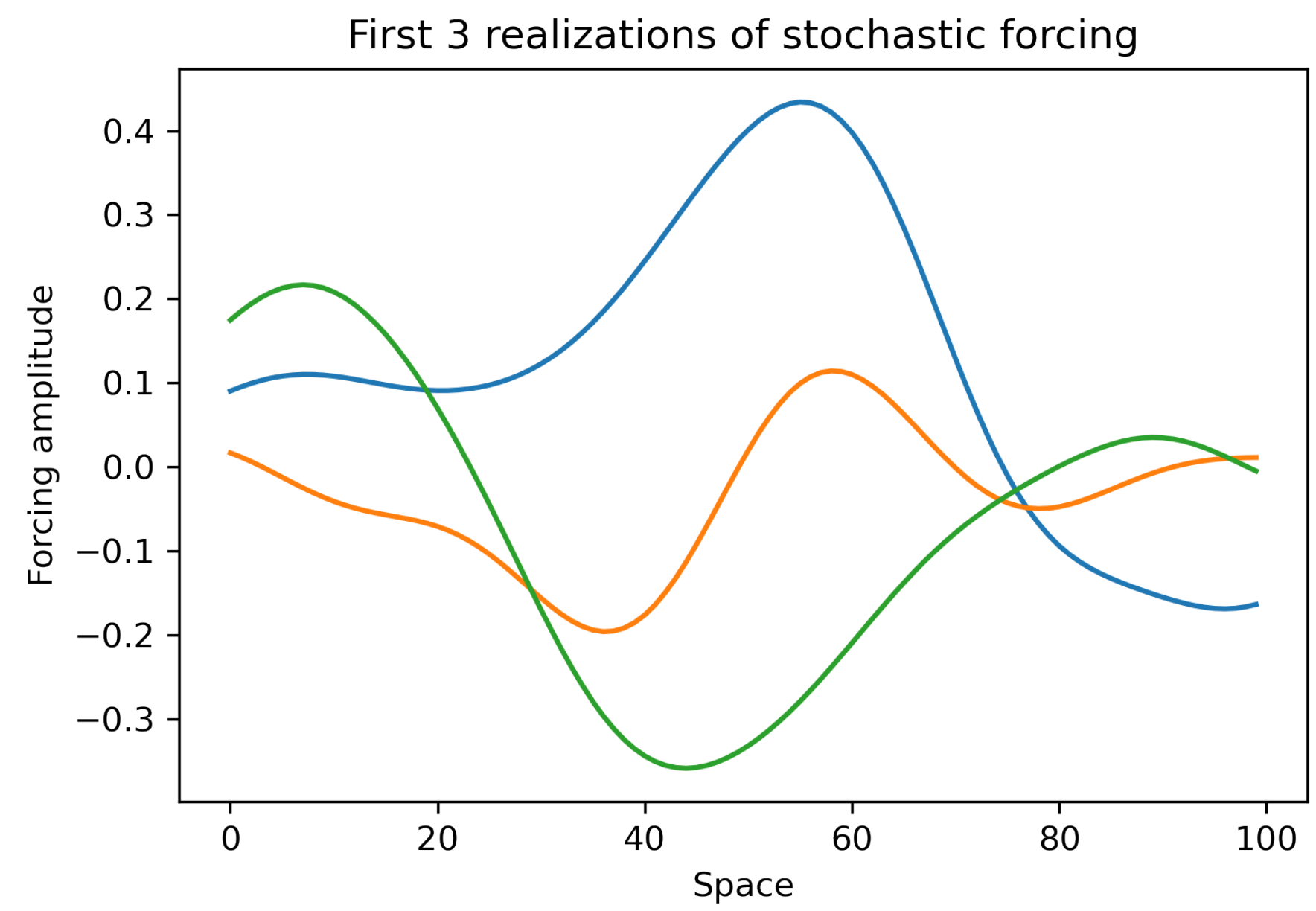
AKA “balanced truncation”: Moore (1981); Farrell and Ioannou (2001); Moore et al. (2022)

Demonstration in a (very) simple system

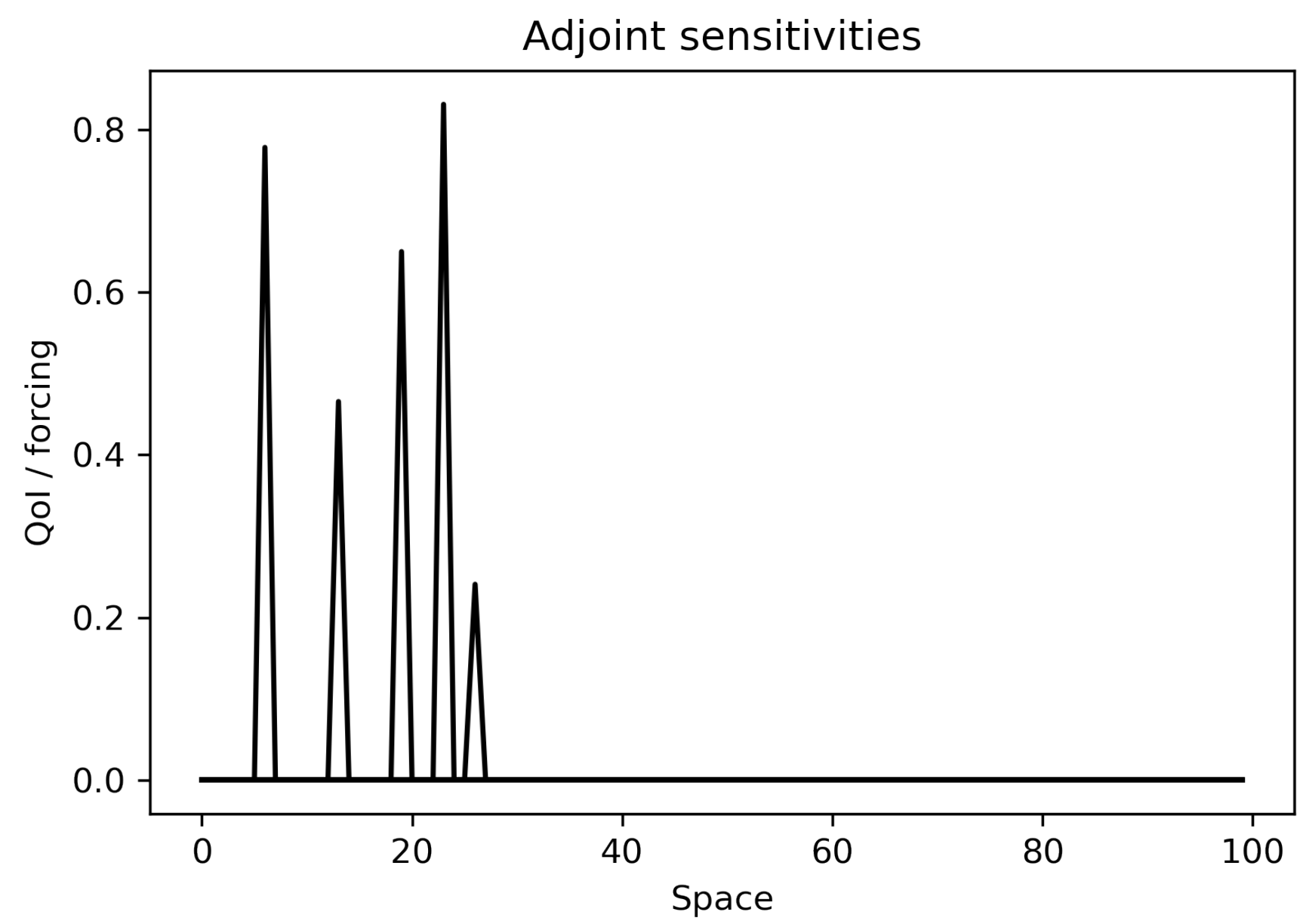


Consider a 1-dimensional system with stochastic forcing that is **smooth in space** and Gaussian **white noise in time**.

Demonstration in a (very) simple system

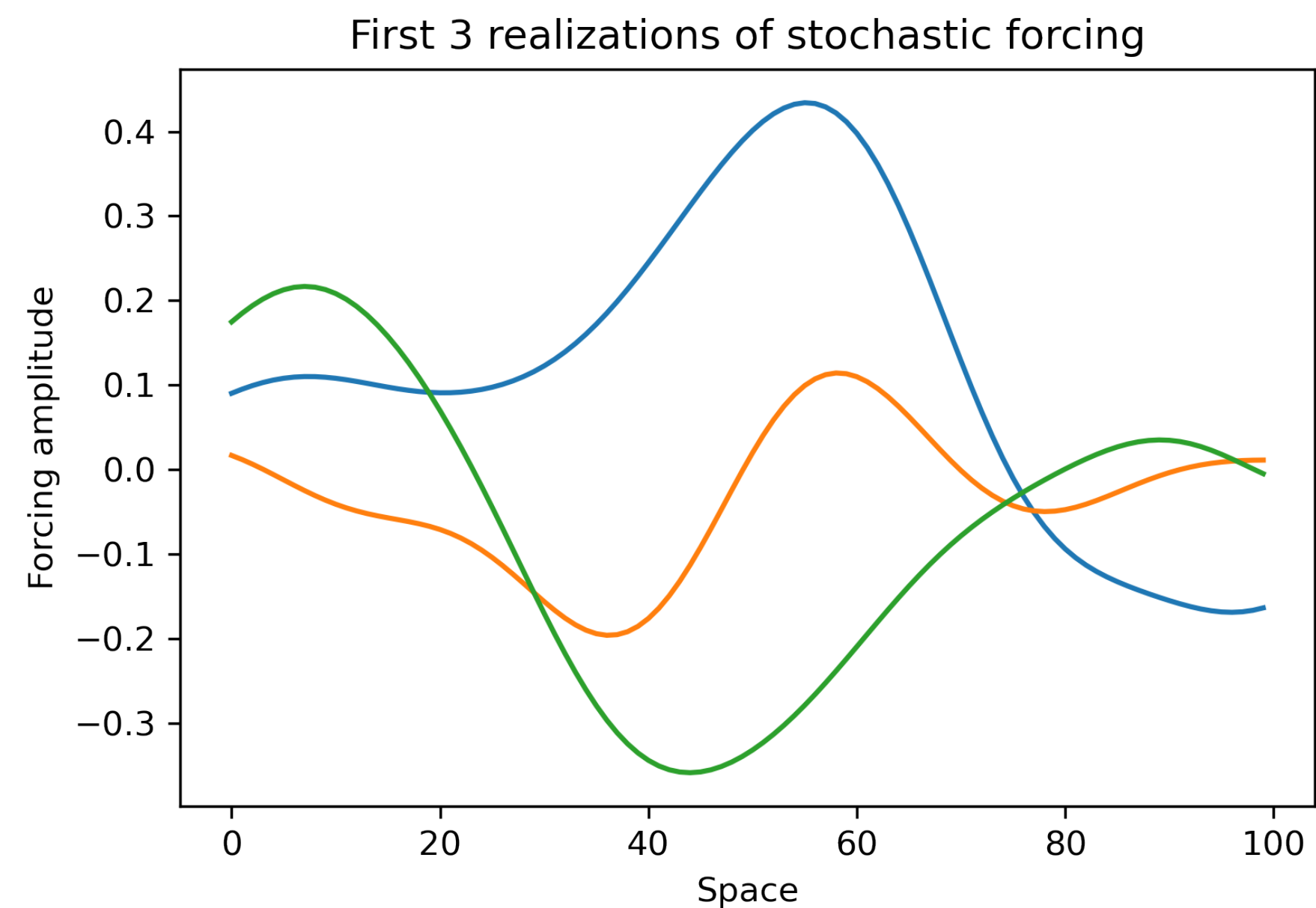


Consider a 1-dimensional system with stochastic forcing that is **smooth in space** and Gaussian **white noise in time**.

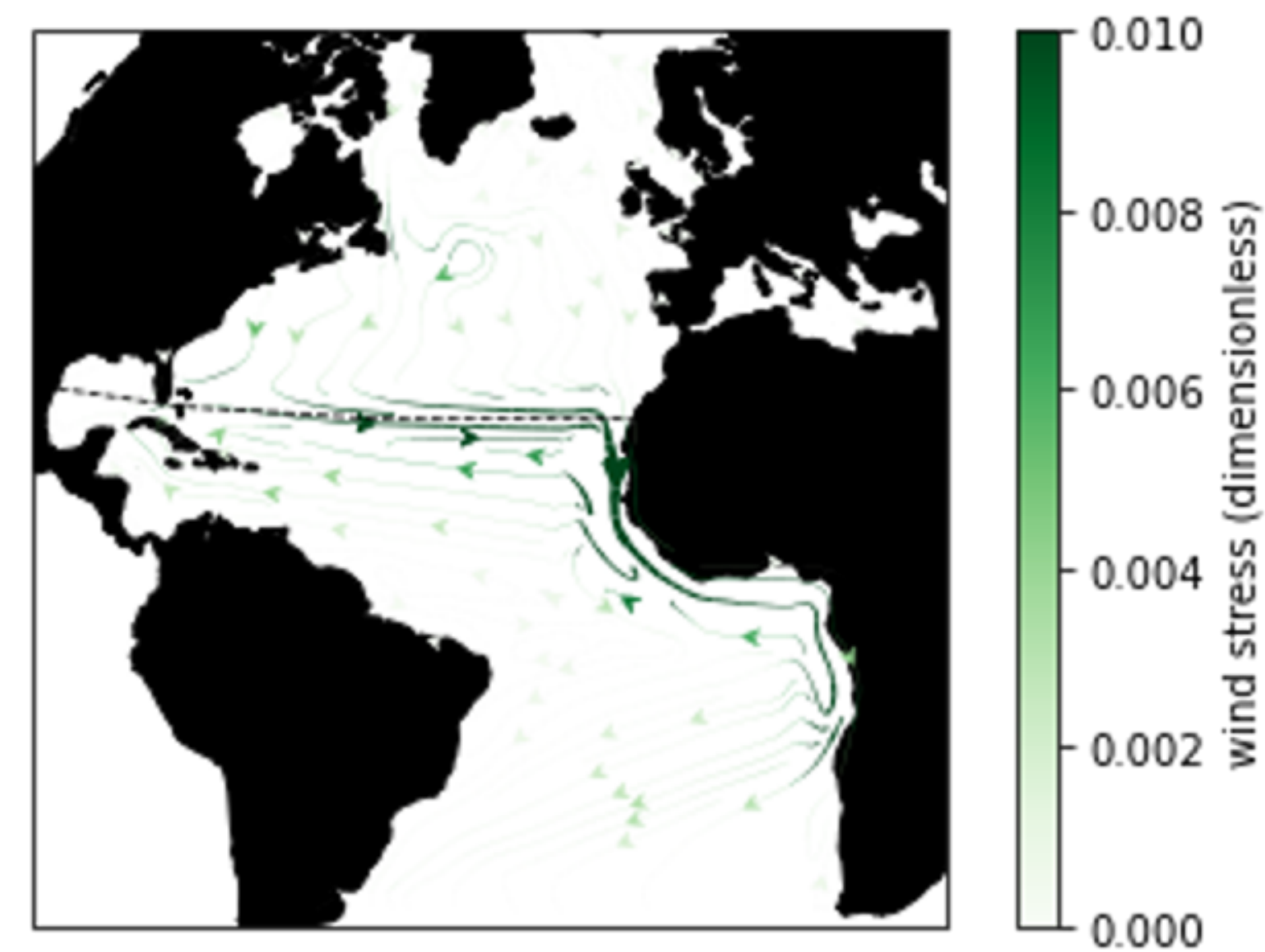
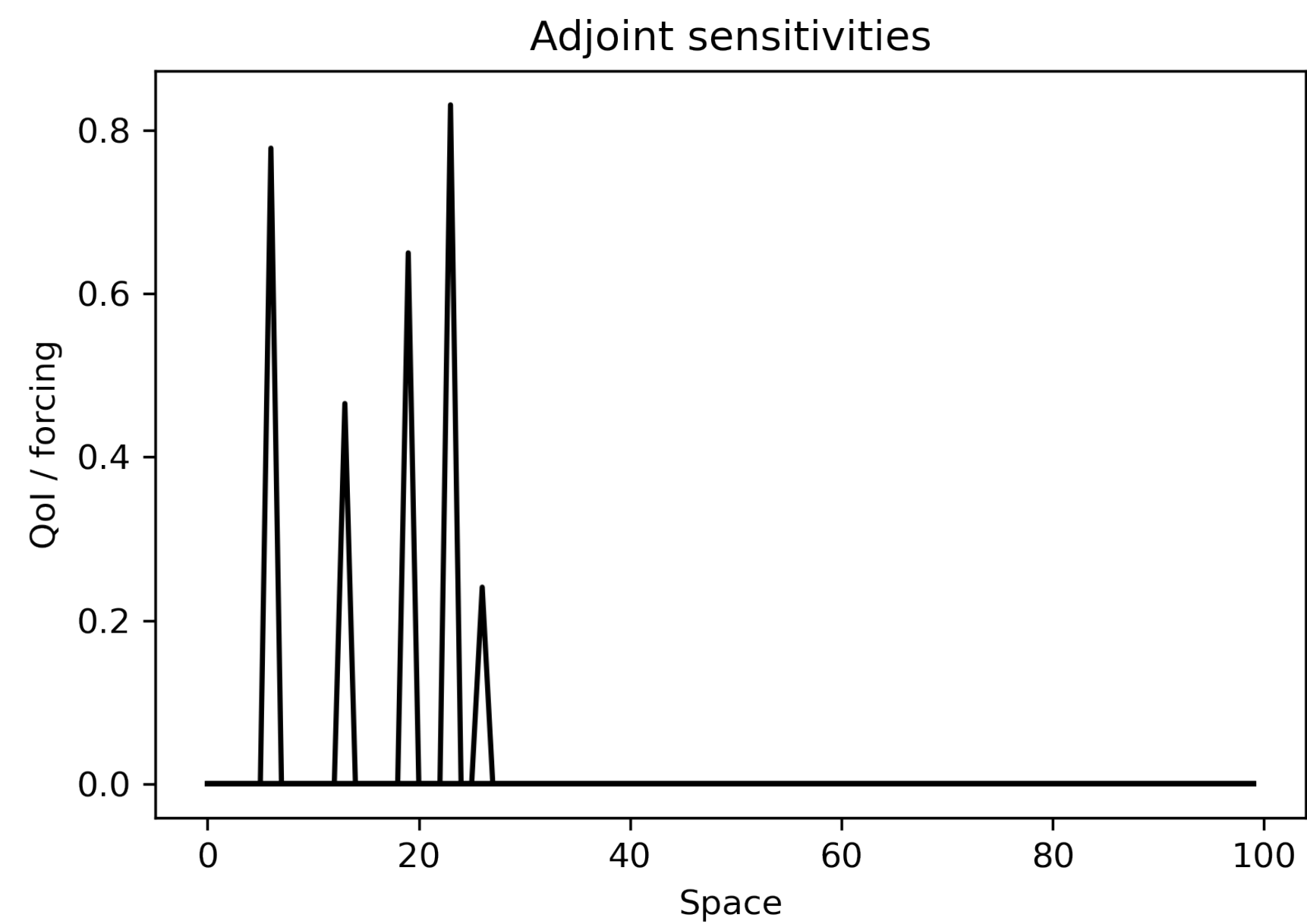
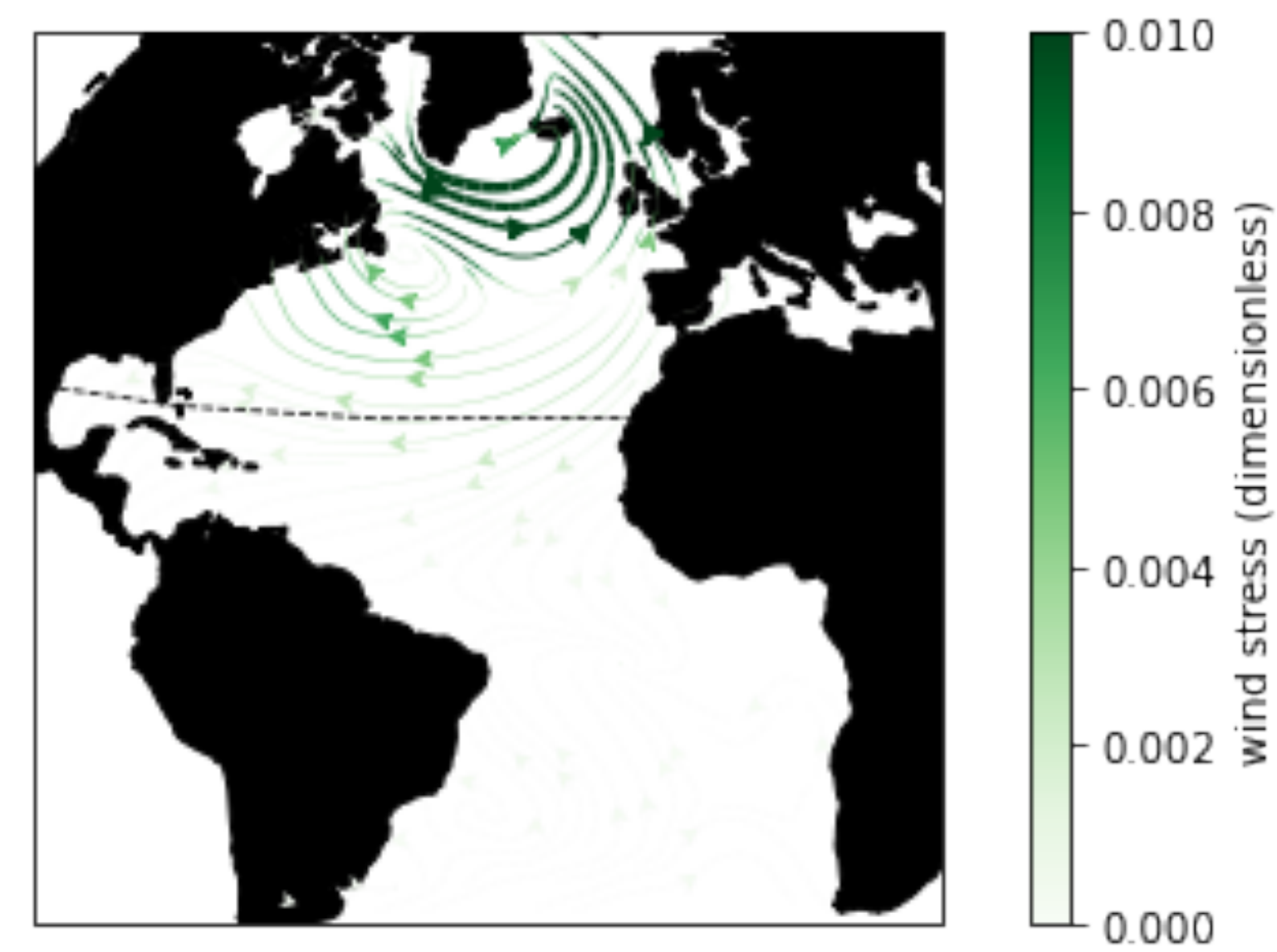


..and adjoint sensitivities of a hypothetical ocean QoI that have **shorter length scales** and are **localized in space**.

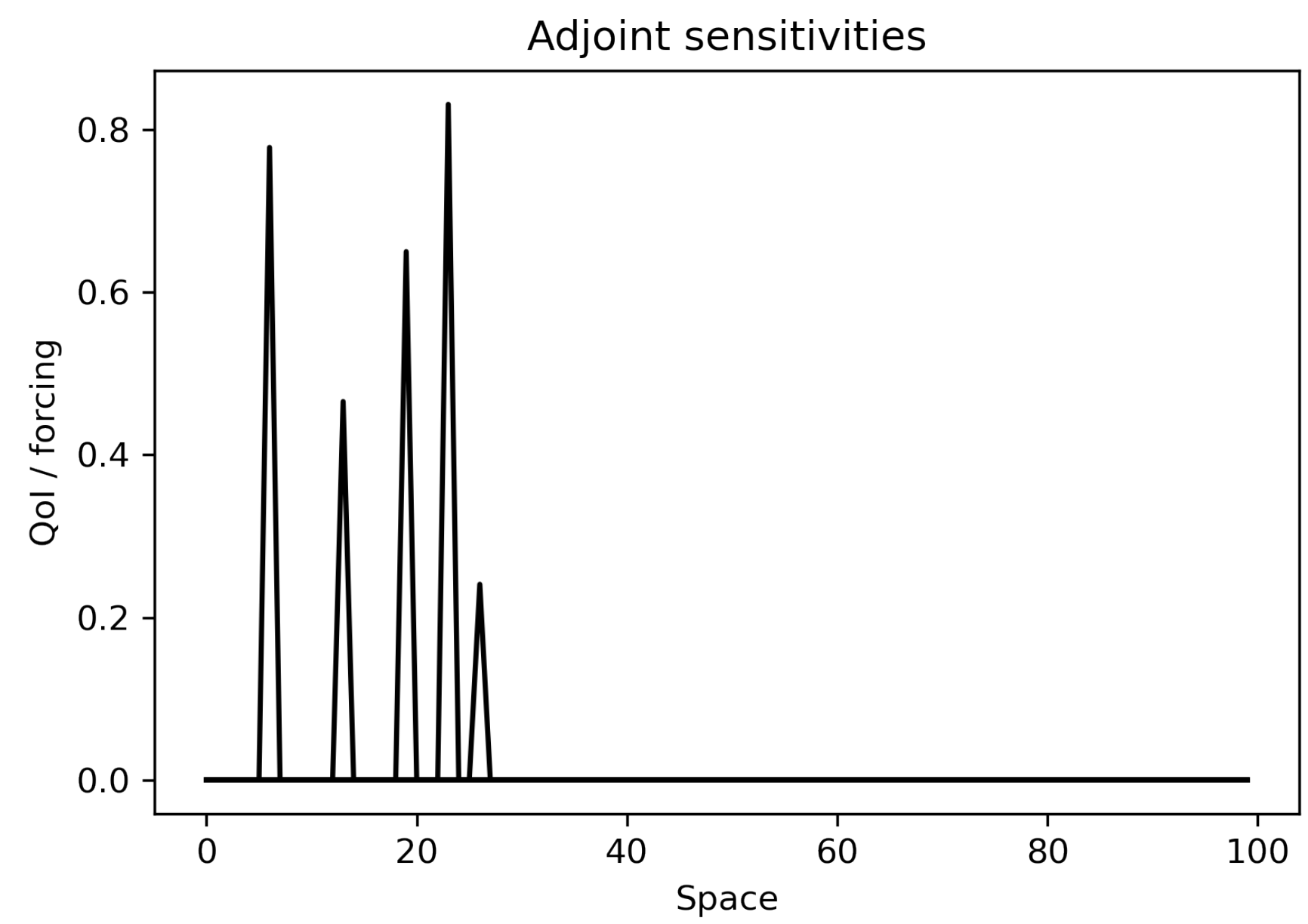
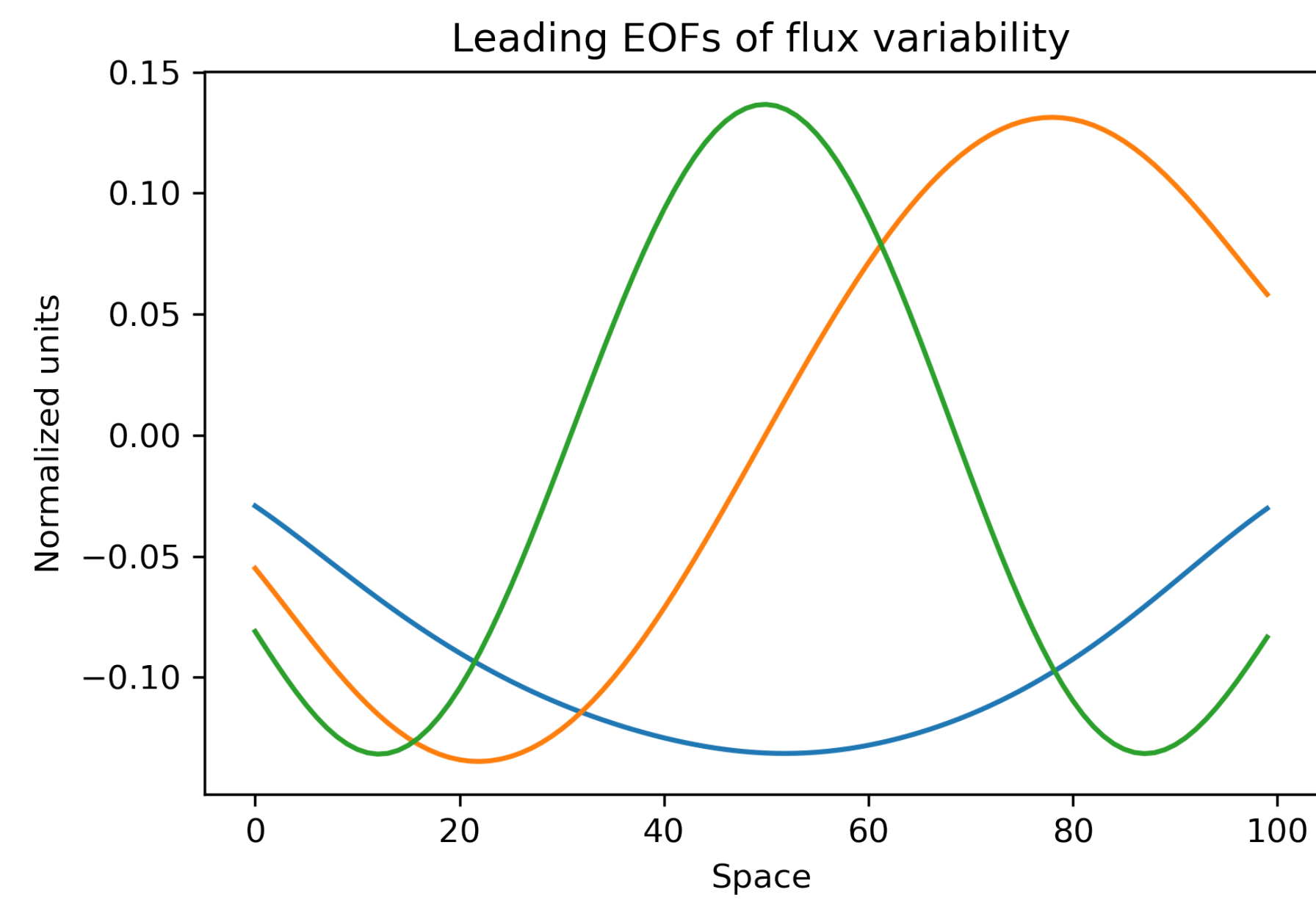
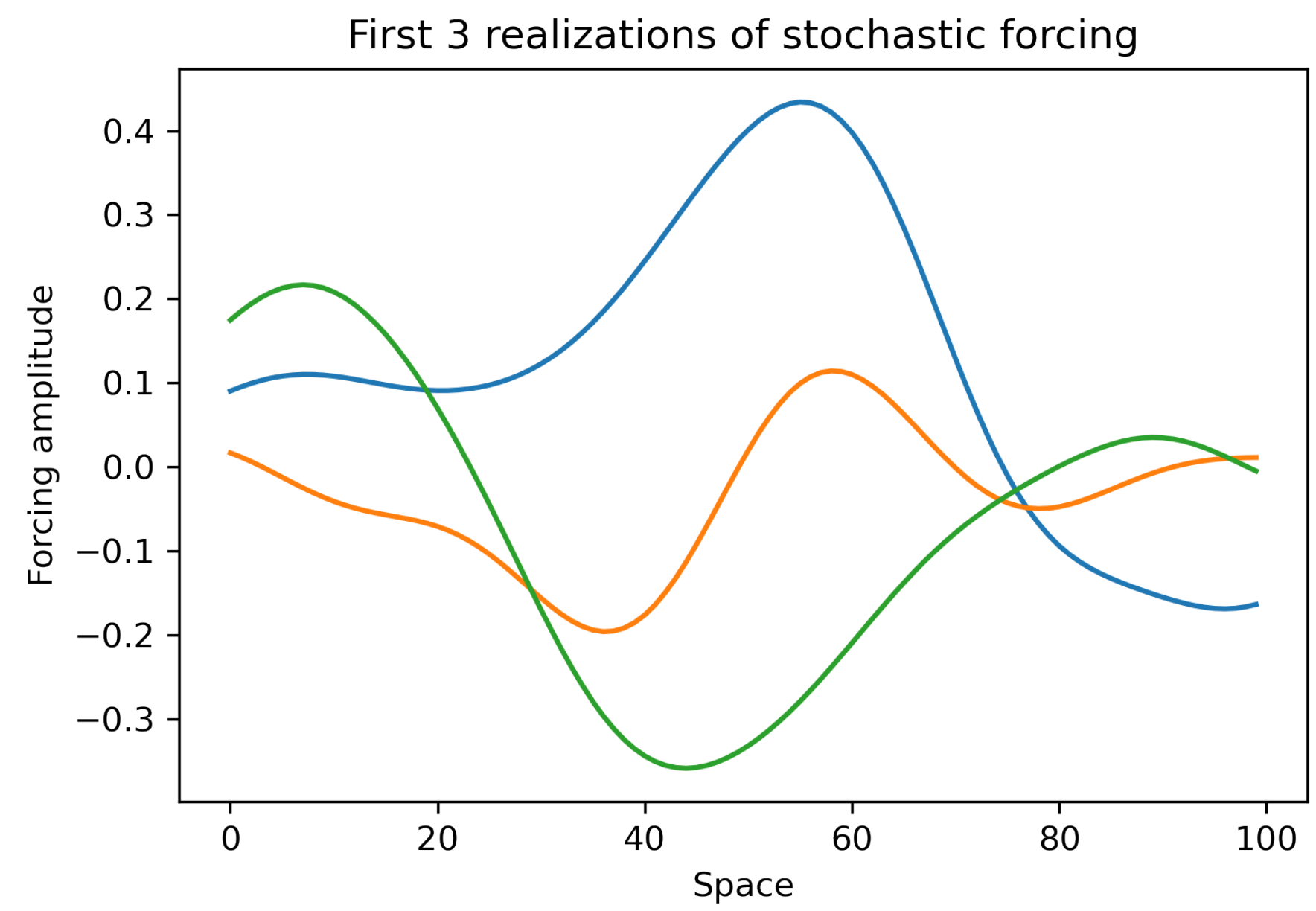
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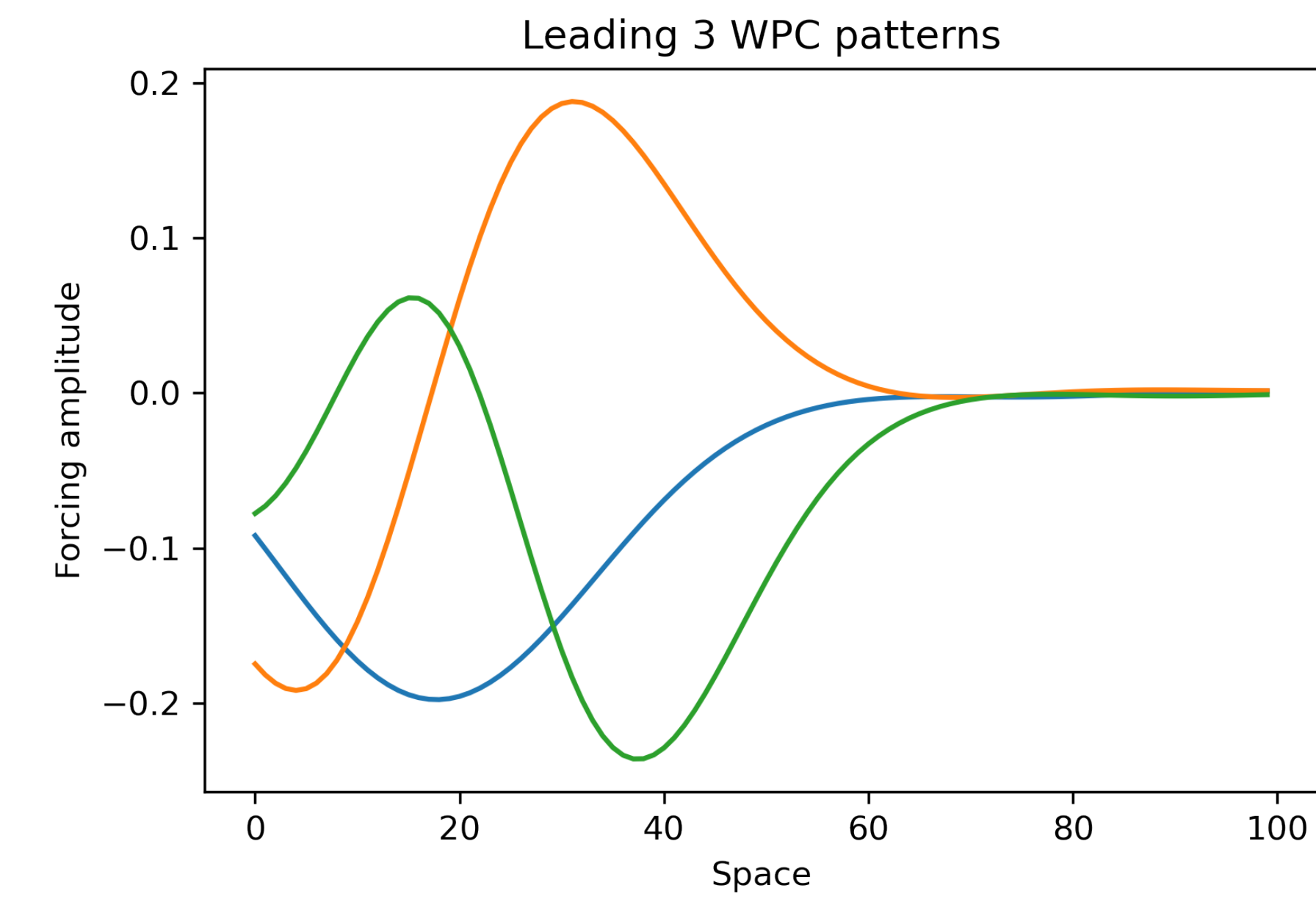
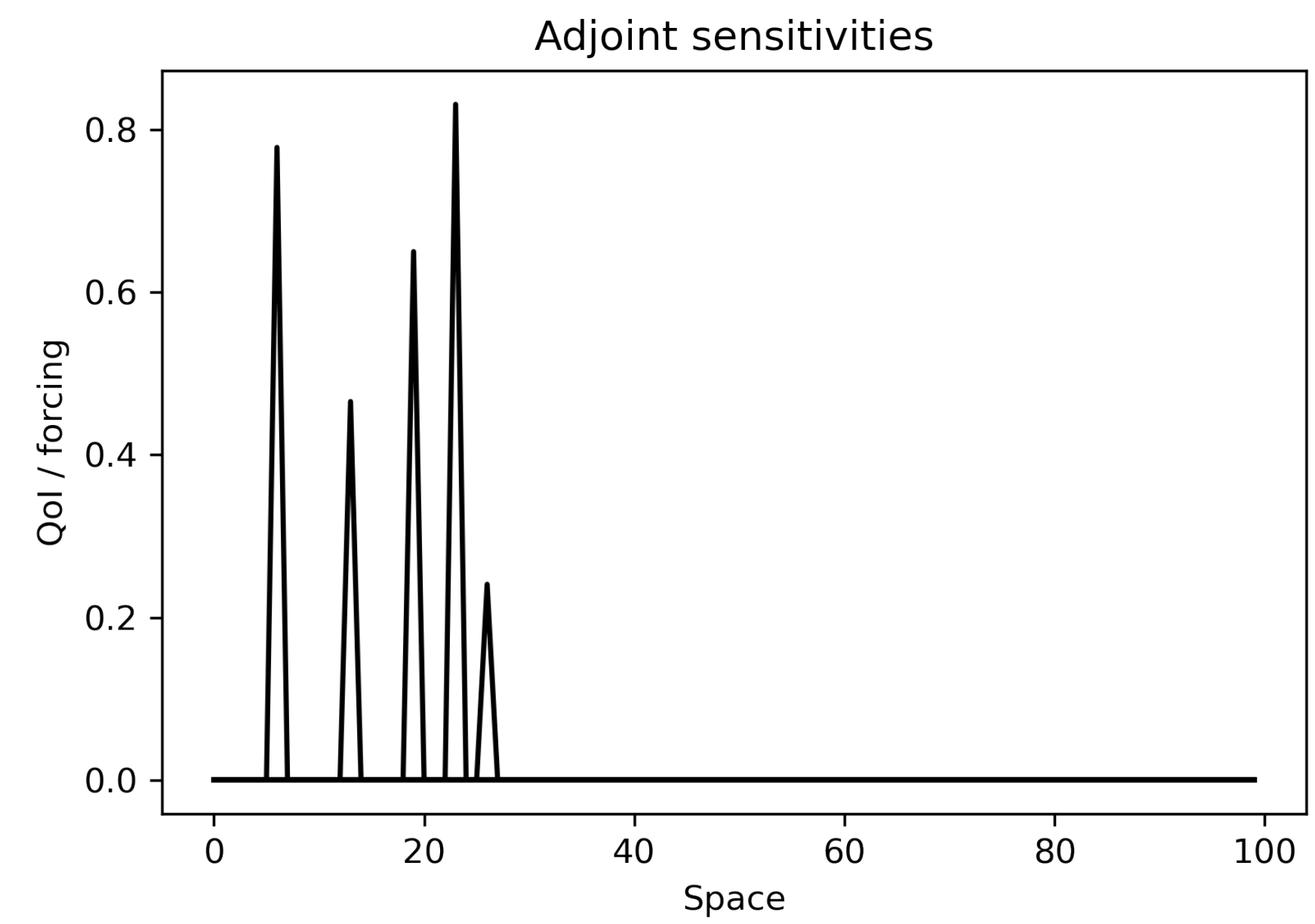
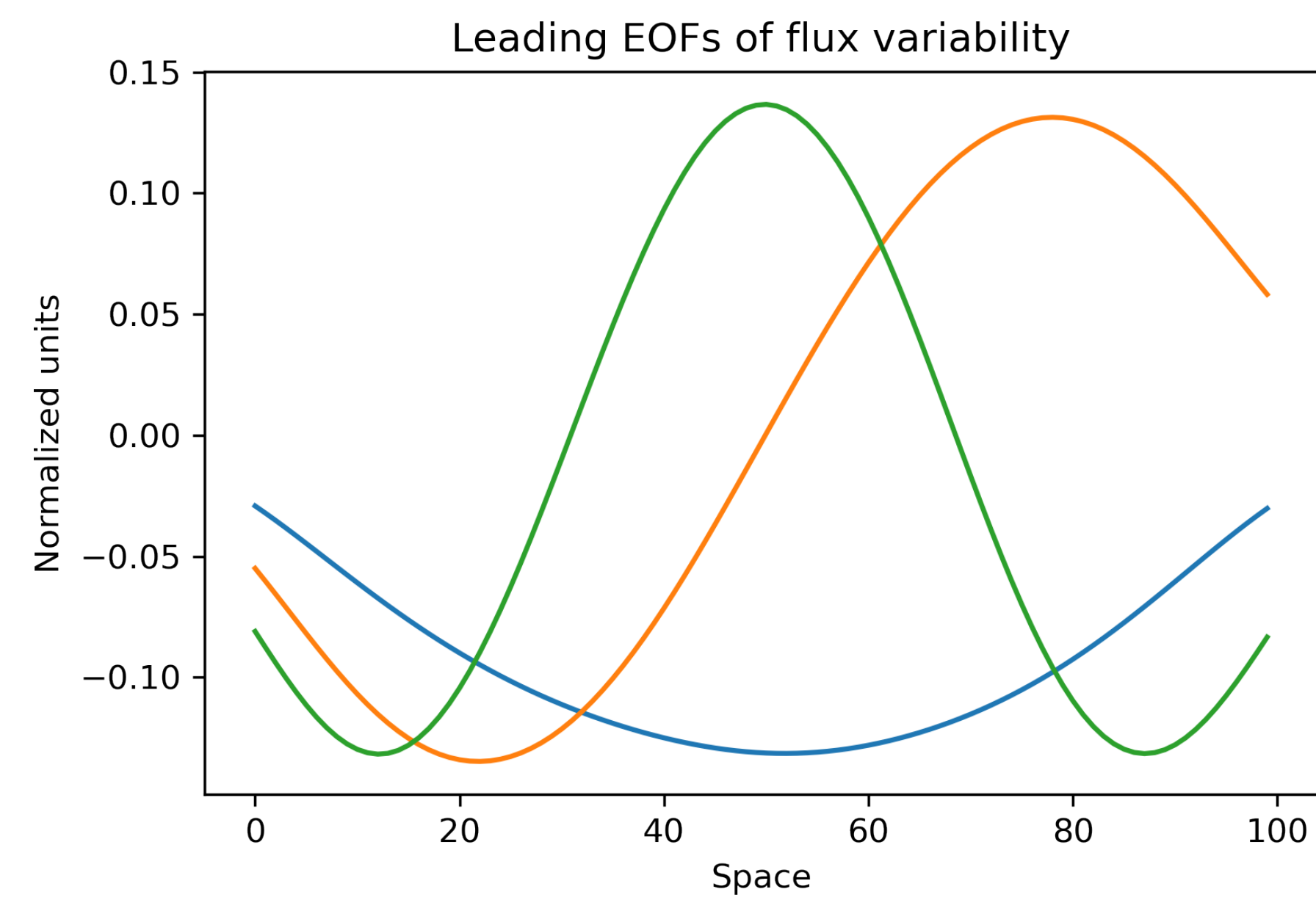
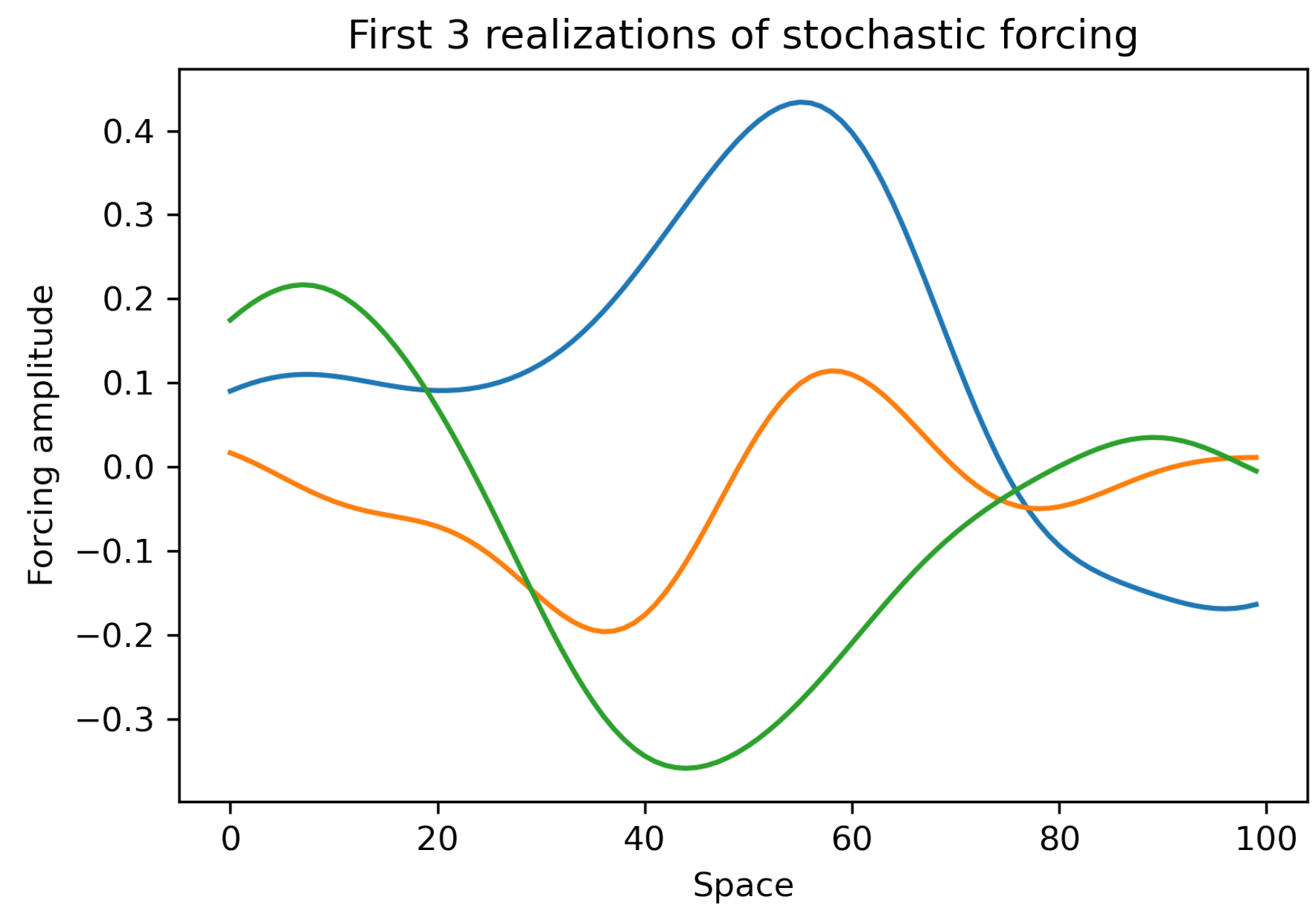
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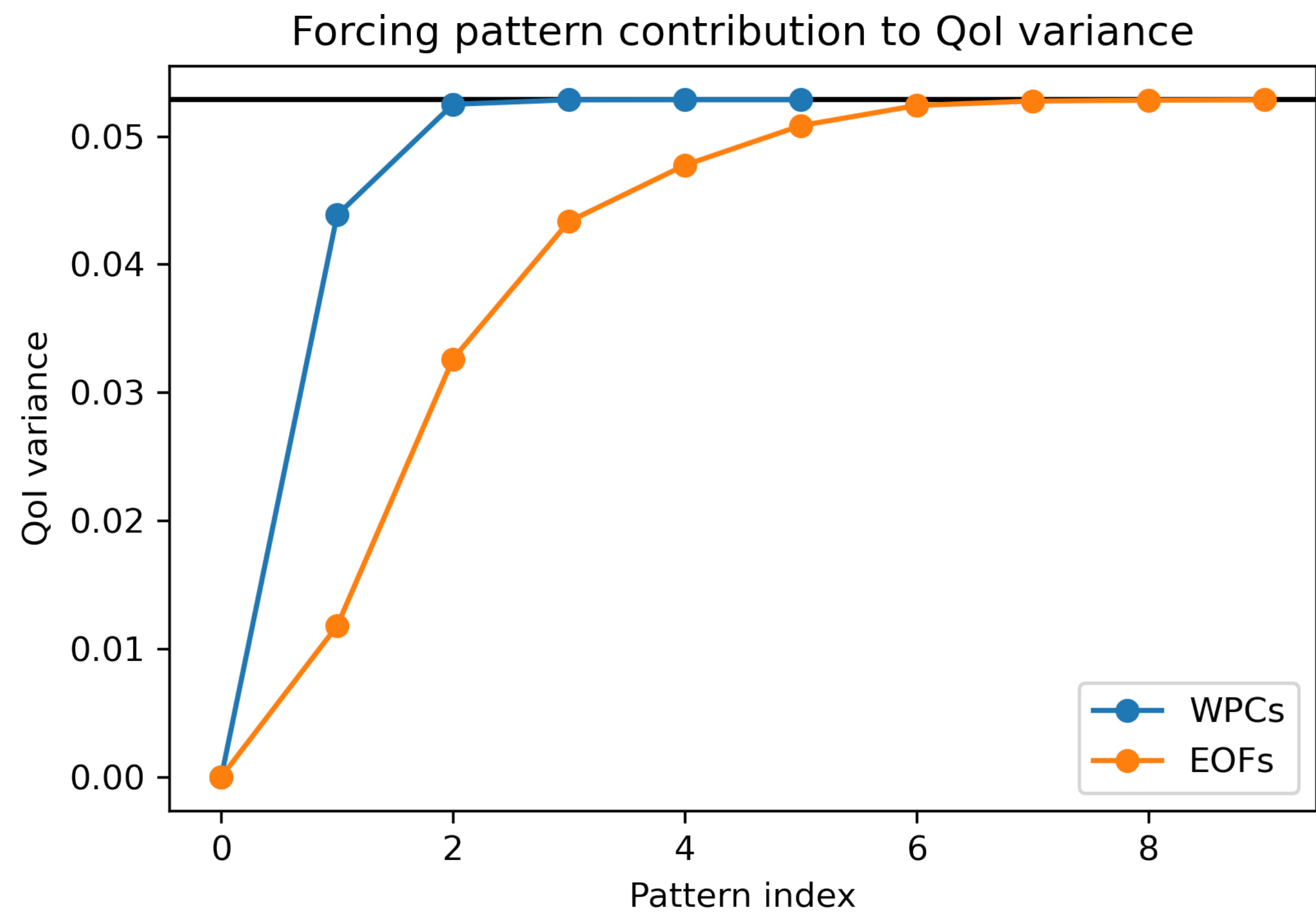
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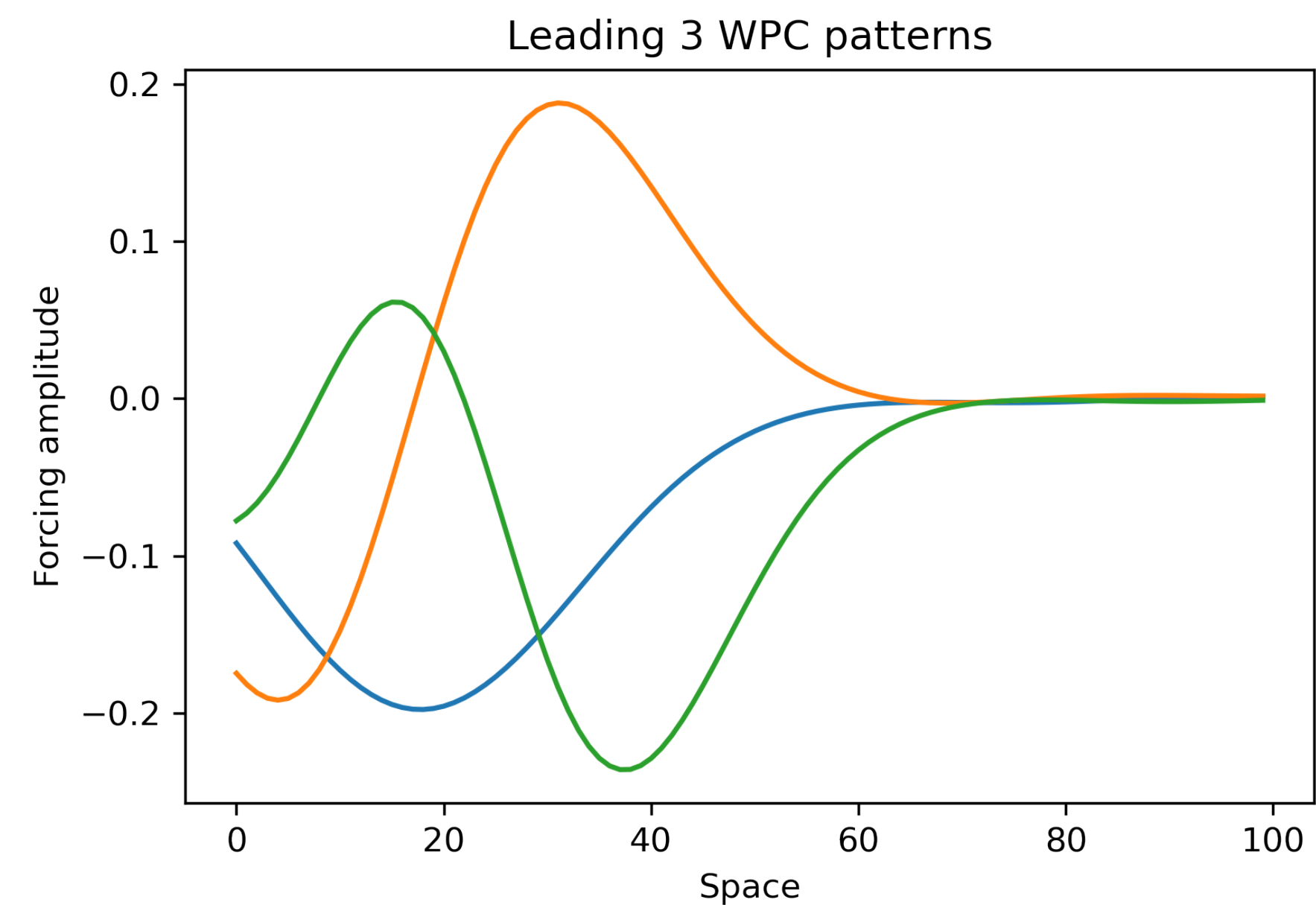
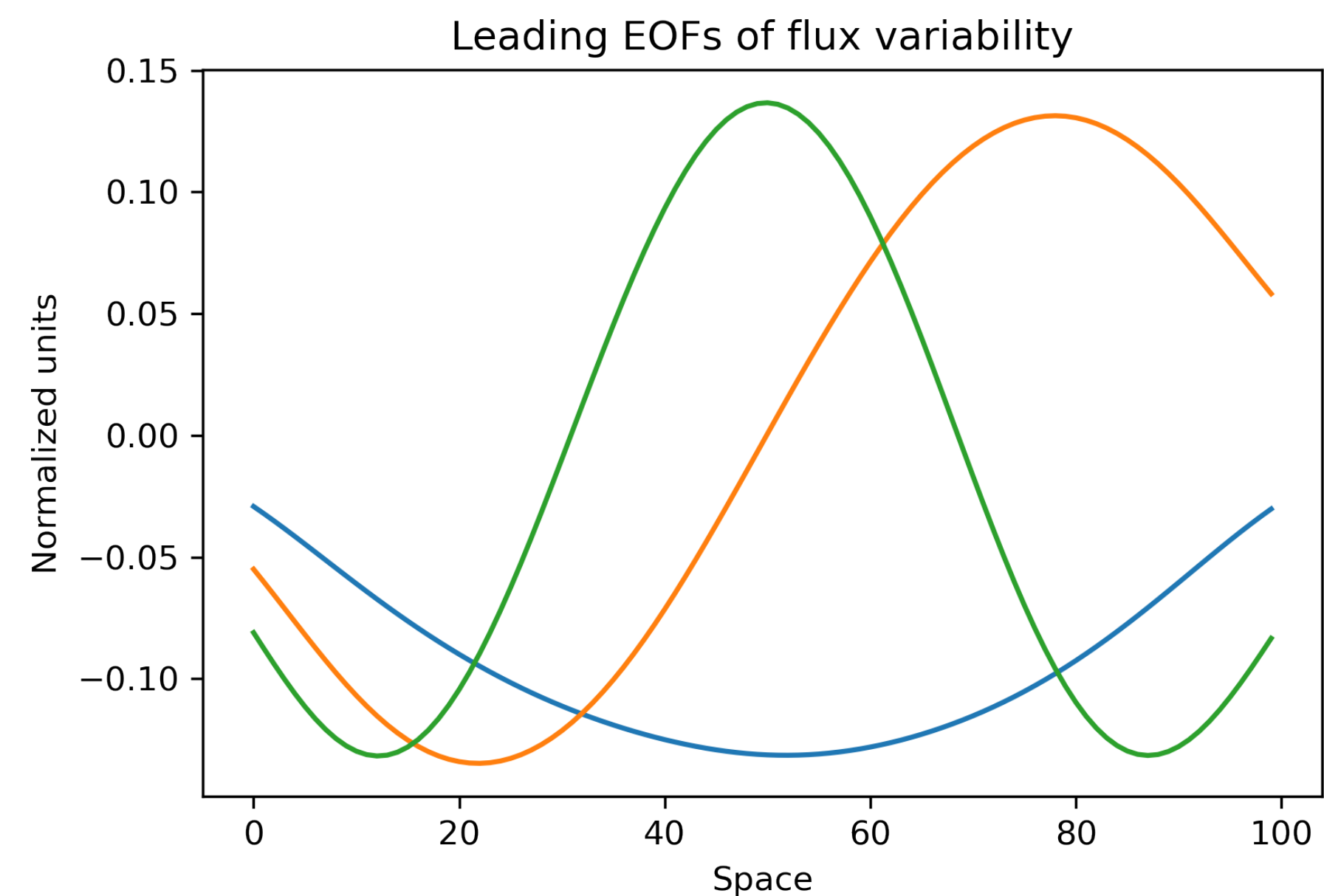
Demonstration in a (very) simple system



Demonstration in a (very) simple system



WPC patterns outperform EOFs at driving QoI variance.



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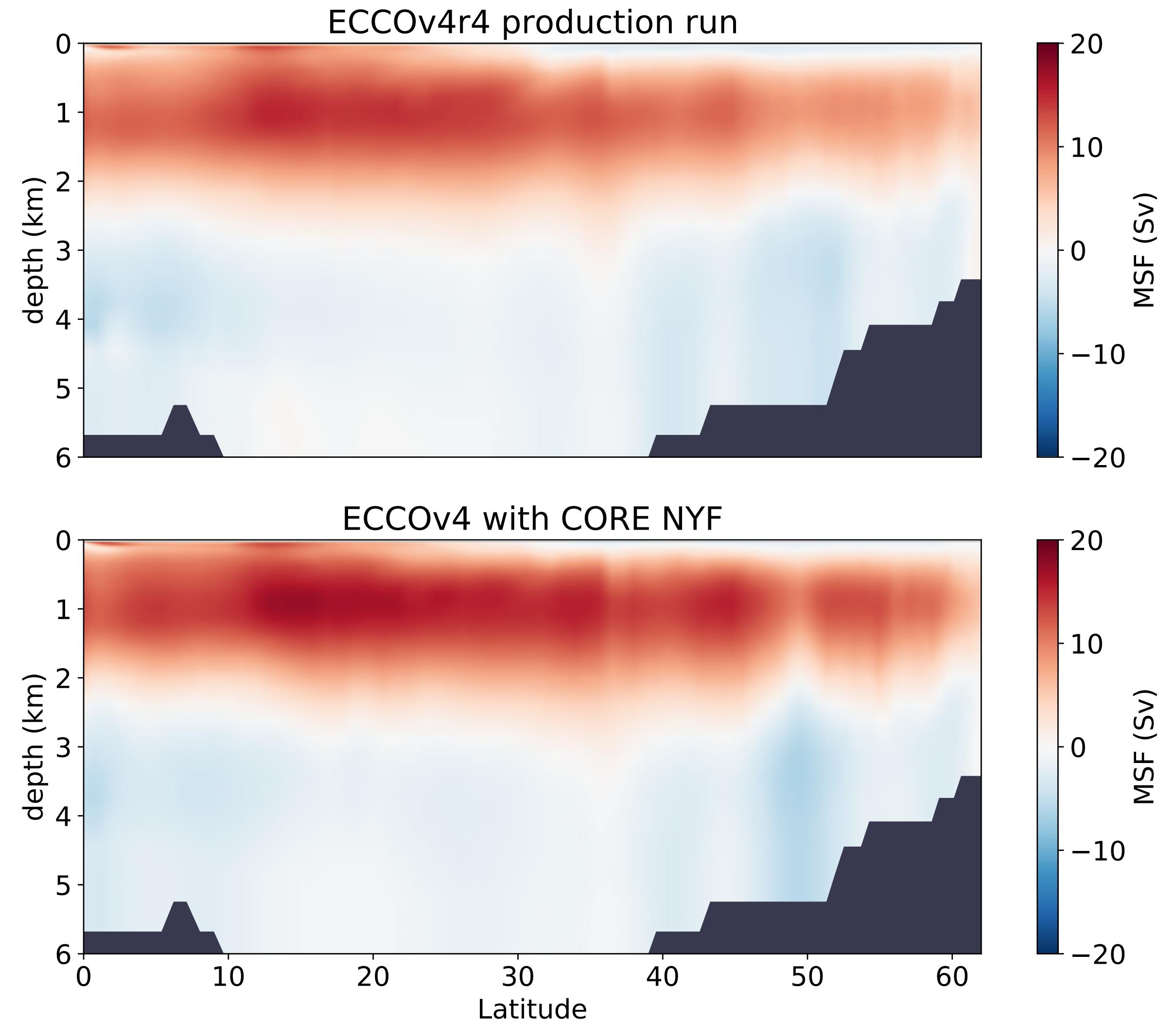
Tracing pathways of decadal AMOC change from atmosphere to ocean

~1° resolution, flux-forced MITgcm ECCO v4 configuration

Ocean and sea ice components spun up under 4800 years following Wolfe et al. (2017).

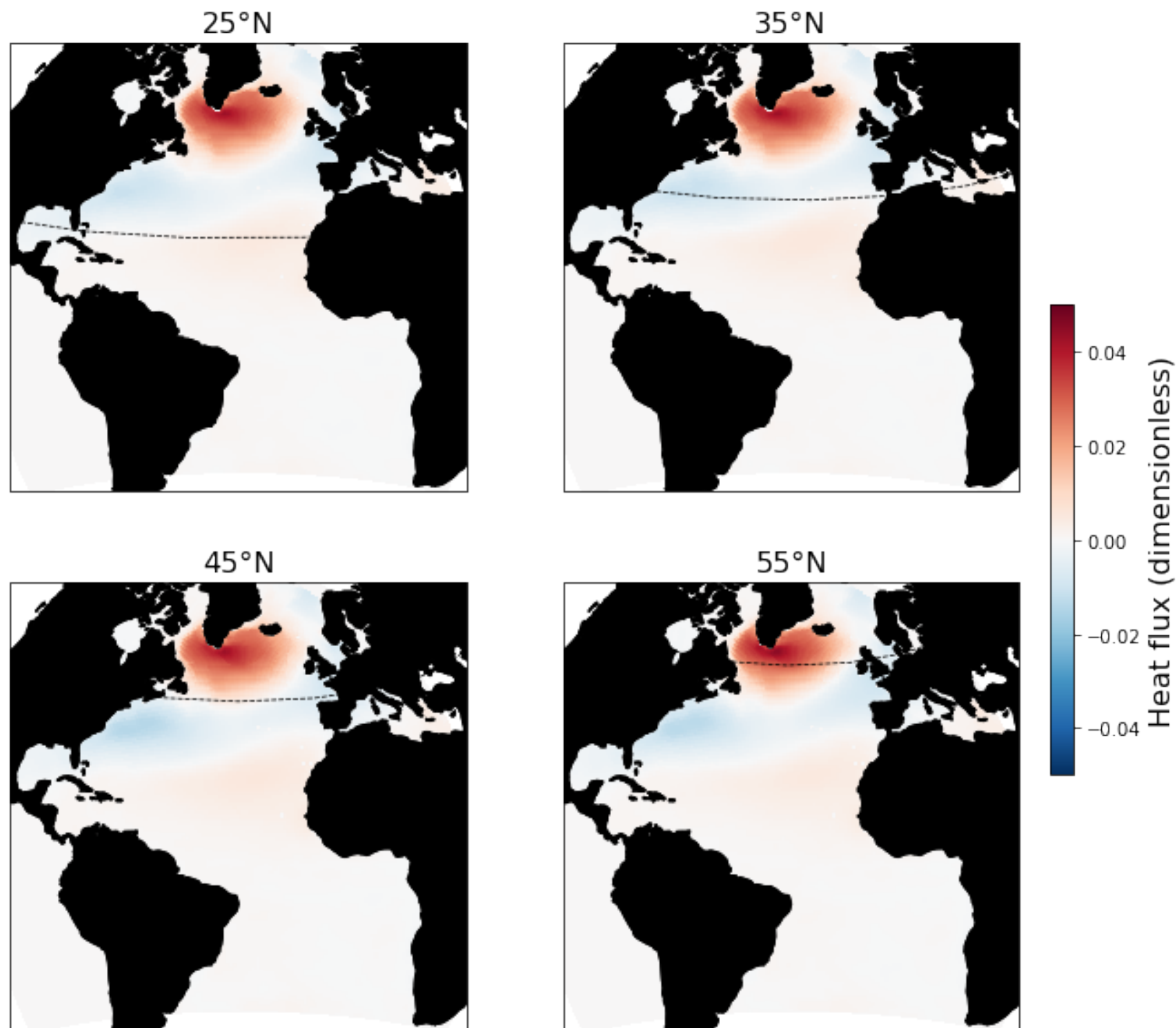
Adjointed and run to compute sensitivities of AMOC transport at climatological maximum depth at **decadal averages** across several latitudes.

Fluxes are 6 hourly from ECCO v4r4.



What are the dominant atmospheric patterns responsible for surface-forced decadal AMOC variability?

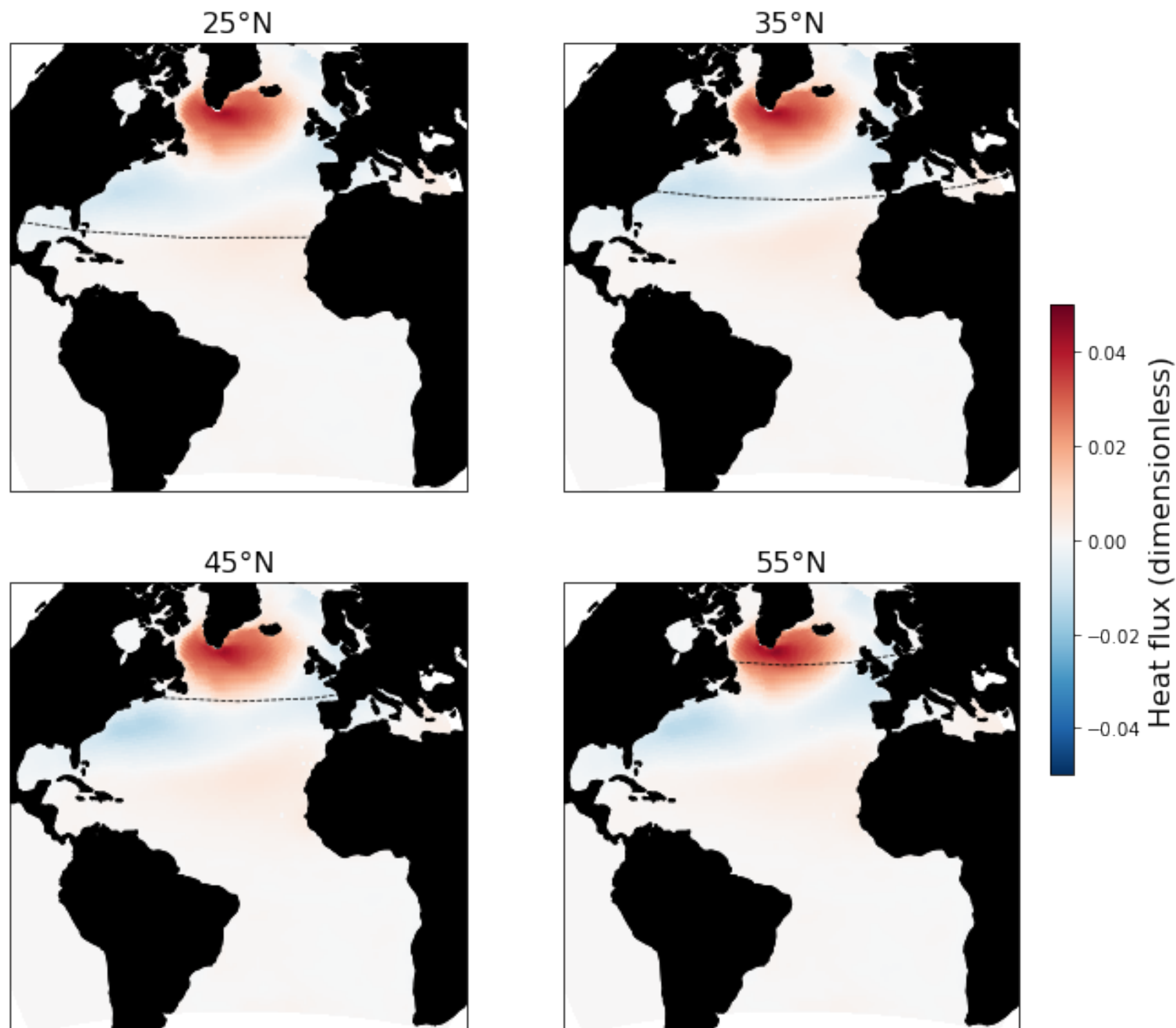
Heat flux patterns (four latitudes):



Leading pattern is almost identical at all four latitudes (>99% agreement)

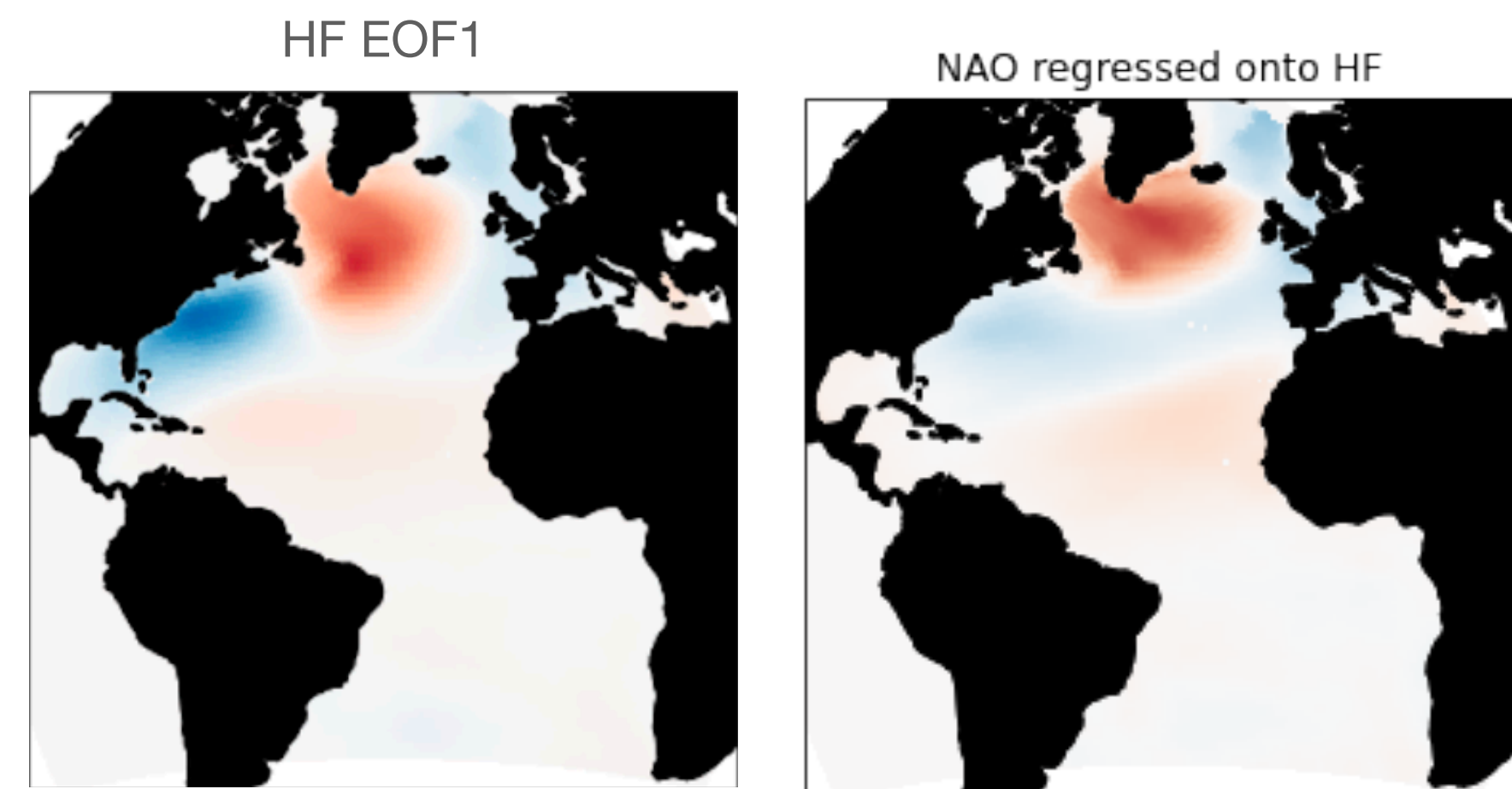
What are the dominant atmospheric patterns responsible for surface-forced decadal AMOC variability?

Heat flux patterns (four latitudes):



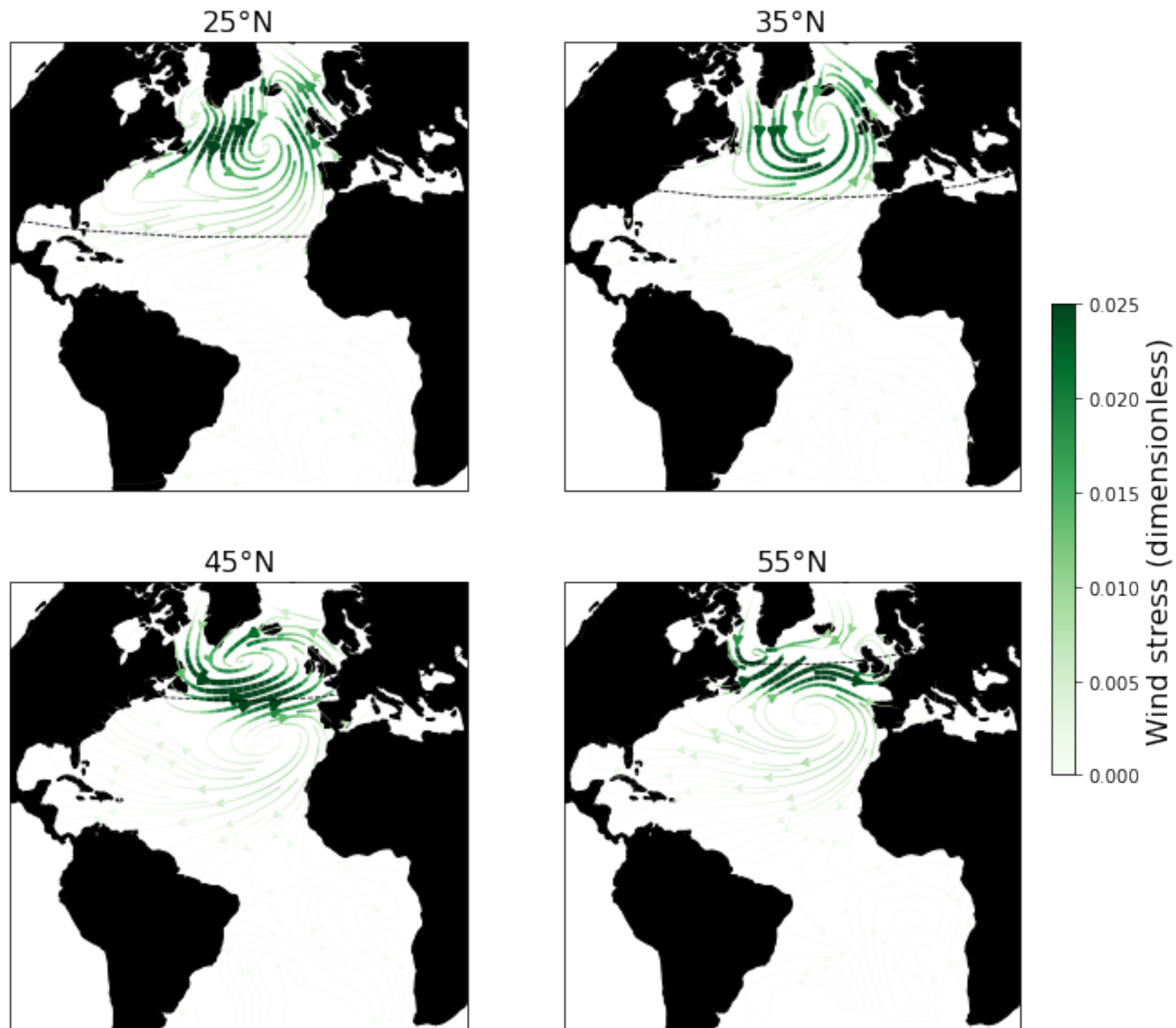
Leading pattern is almost identical at all four latitudes (>99% agreement)
Structurally different from leading EOF pattern

... but highly similar to the heat flux signature of NAO (>90% agreement)



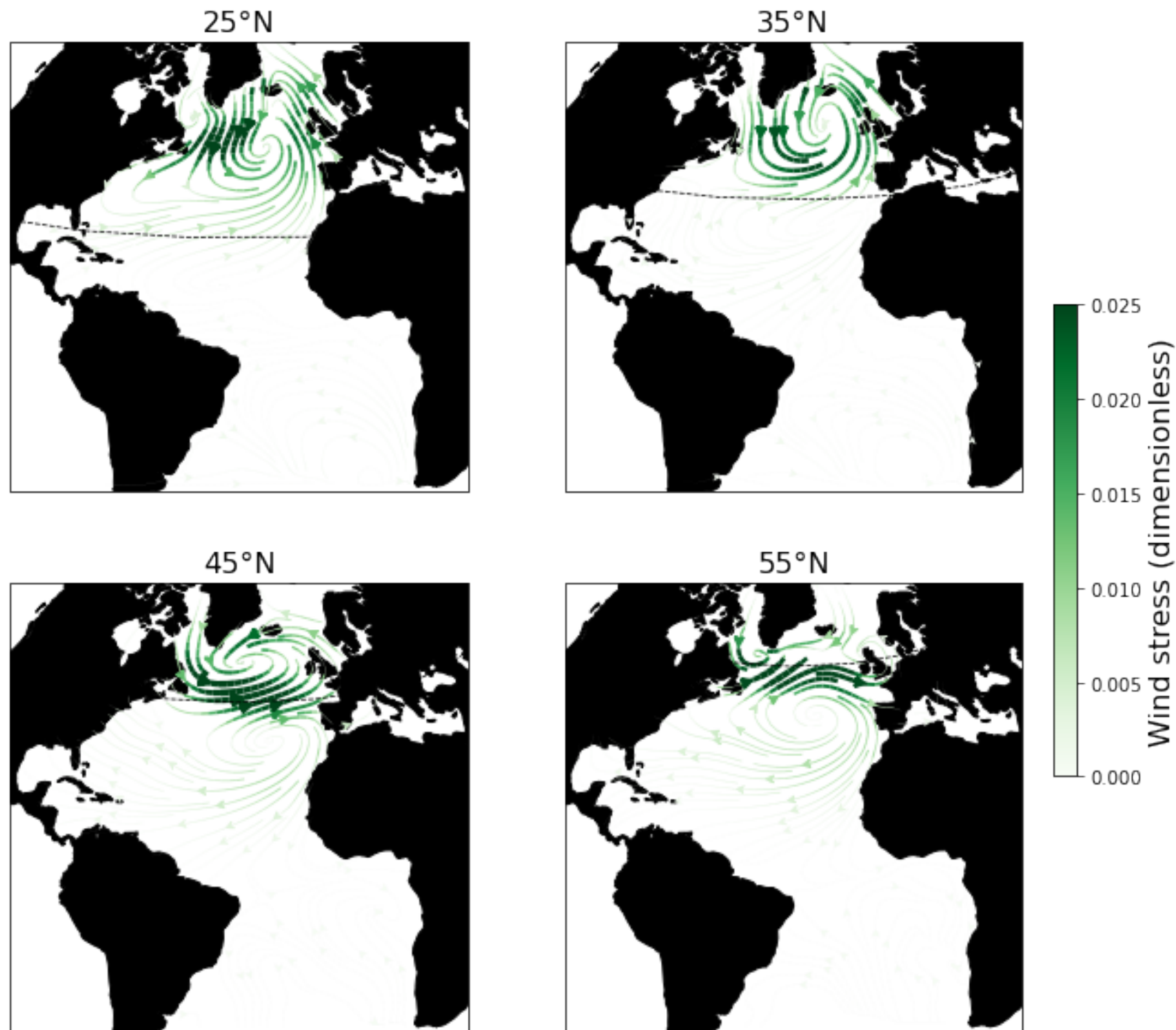
What are the dominant atmospheric patterns responsible for surface-forced decadal AMOC variability?

Leading WPC patterns (four latitudes):

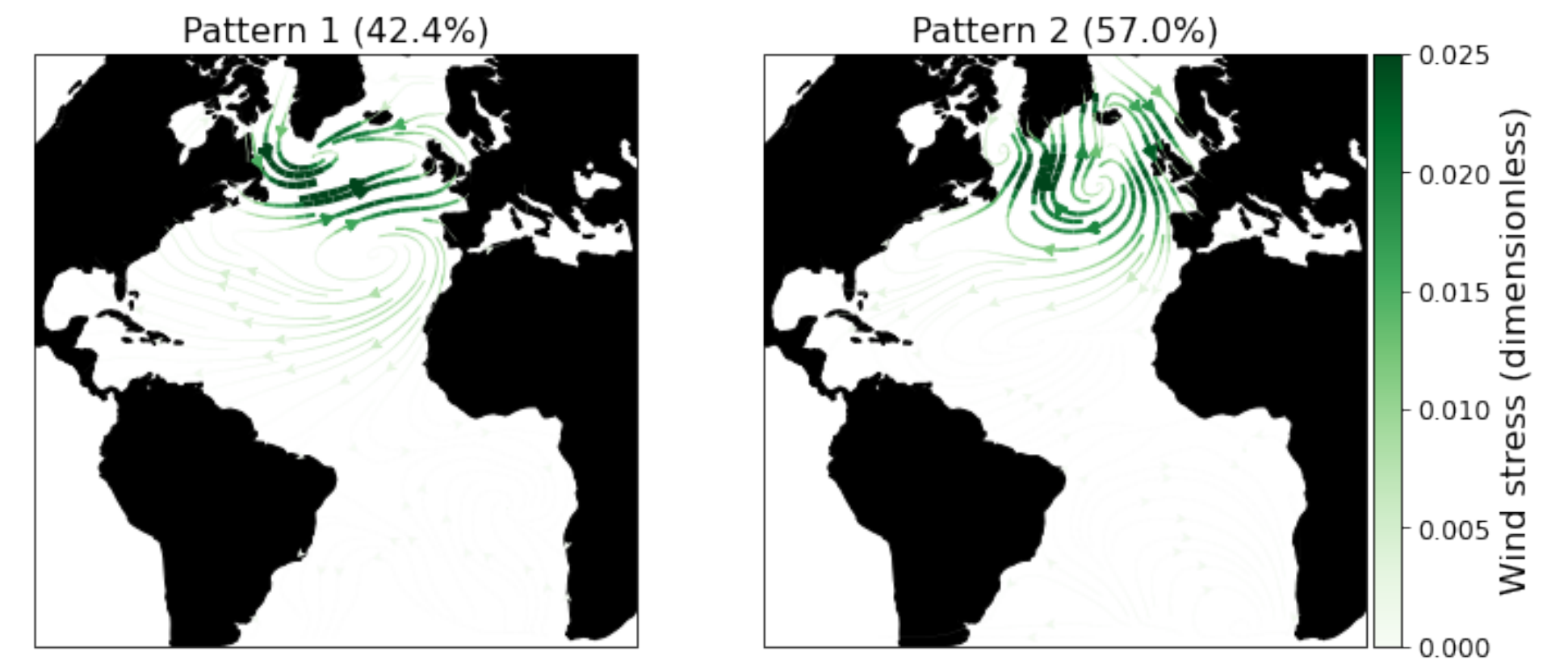


What are the dominant atmospheric patterns responsible for surface-forced decadal AMOC variability?

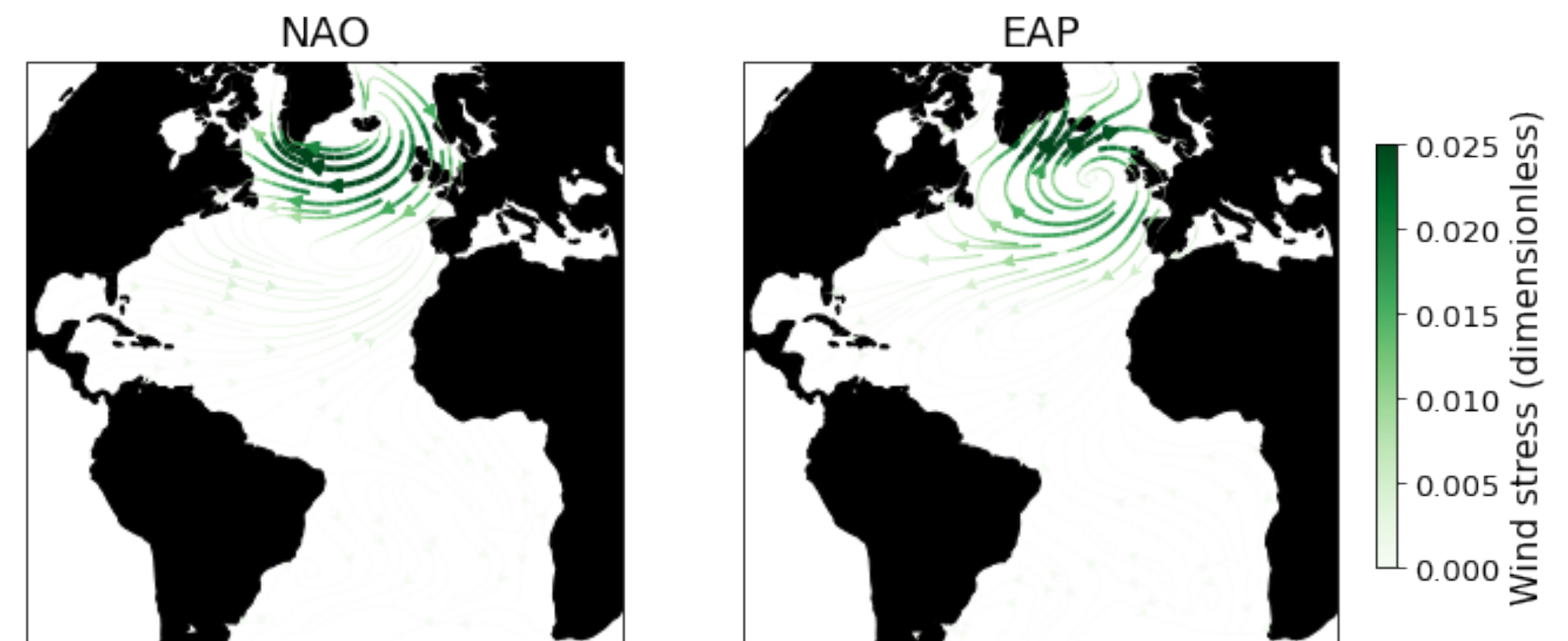
Leading WPC patterns (four latitudes):



Substantially different between AMOC latitudes **but** can be >99% explained by a subpolar pattern (1) and a subtropical pattern (2)



Qualitative similarities to the NAO and EAP:



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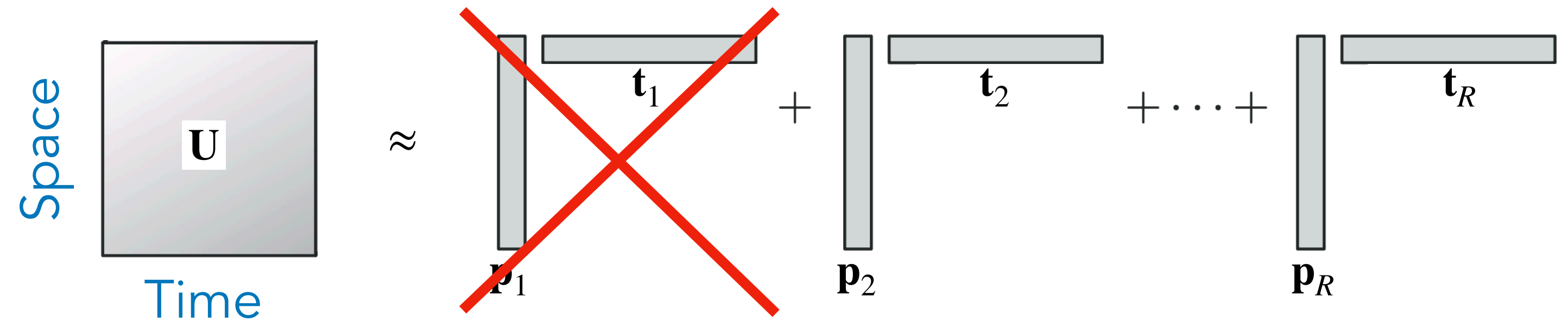
Tracing pathways of decadal AMOC change from atmosphere to ocean

Using ECCO as a climate sandbox

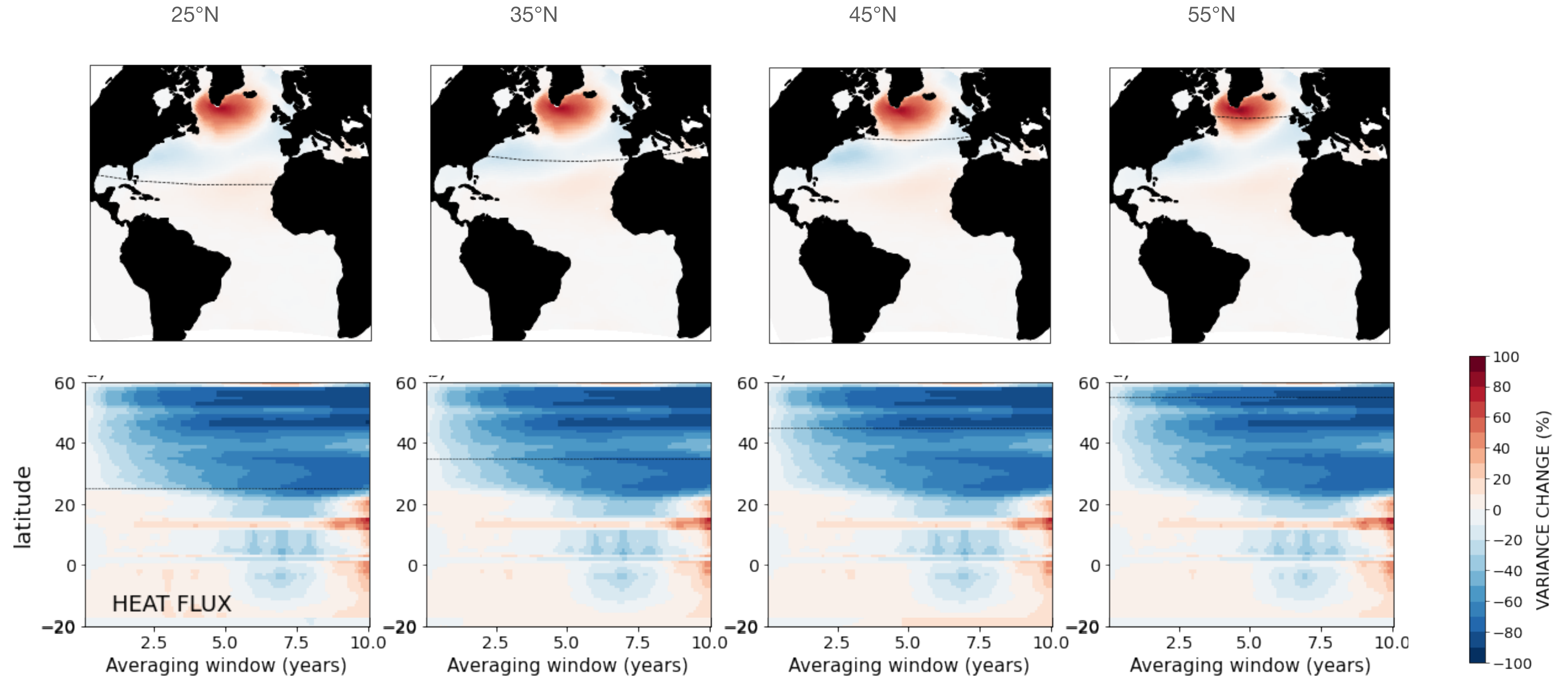
ECCO is a **forward run** of the **MITgcm** that conserves ocean properties.

...so we can make **changes** to the forcing and rerun the MITgcm to evaluate **impacts** and **mechanisms** of atmospheric forcings.

1. Omit a leading WPC pattern

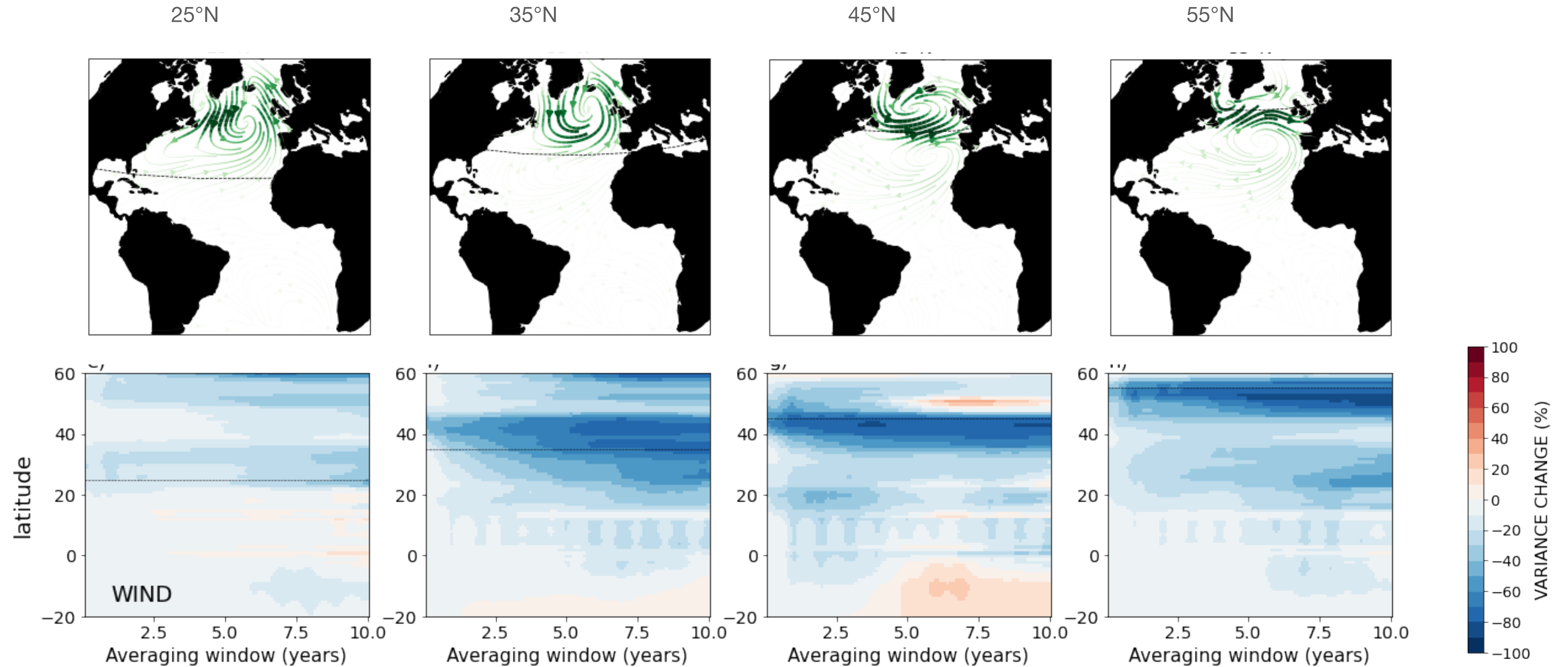


How much AMOC variability do these patterns explain?



Up to **90% change** in variance at the decadal time scale (vs. <30% with the first EOF)
NB: this is in the full nonlinear model (not tangent linear)

How much AMOC variability do these patterns explain?



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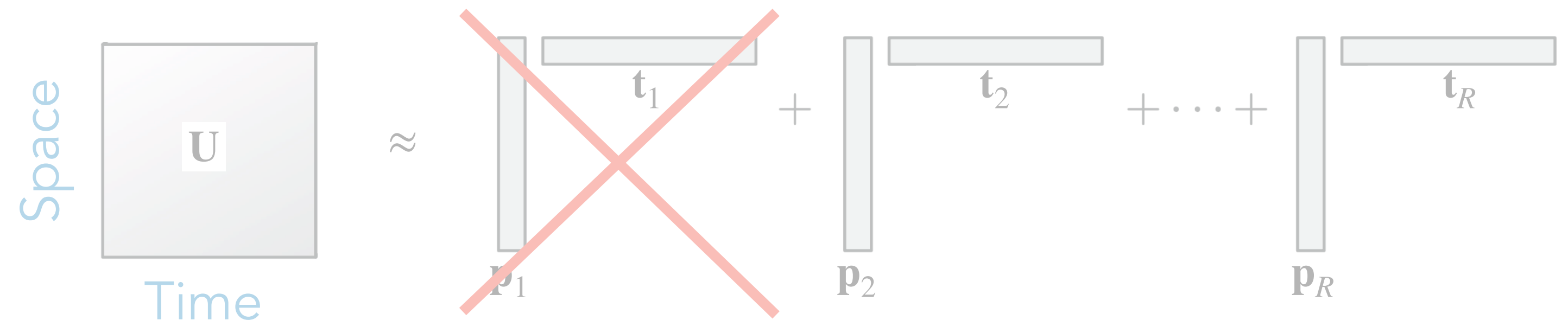
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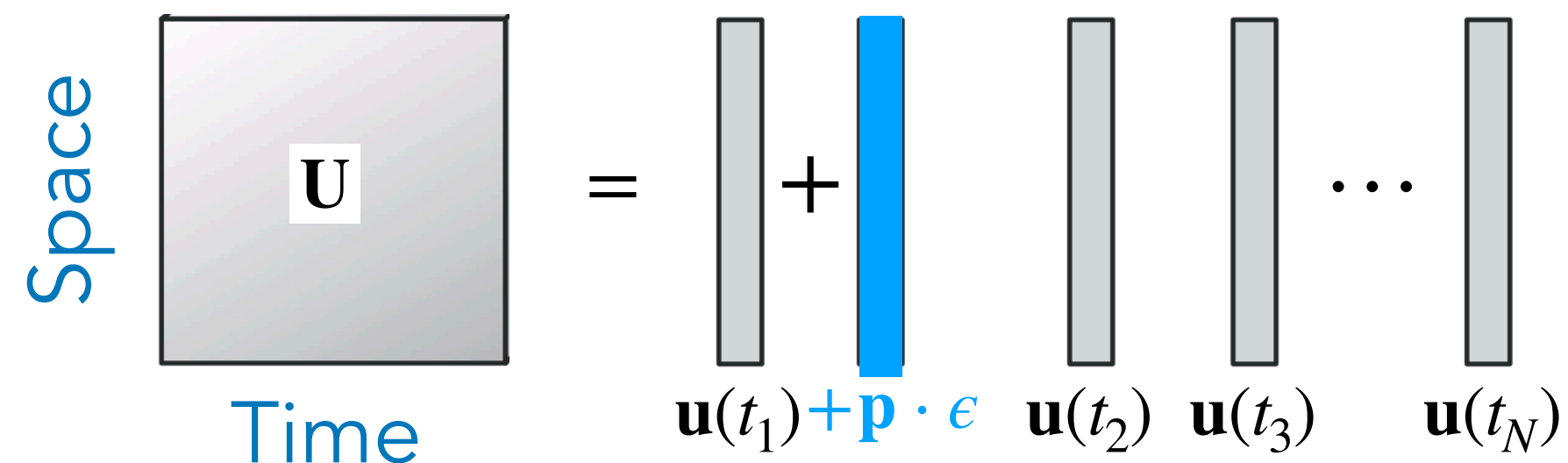
The final product is a **forward run of the MITgcm** that conserves ocean properties.

...so we can make **changes** to the forcing and rerun the MITgcm to evaluate **impacts** and **mechanisms** of atmospheric forcings.

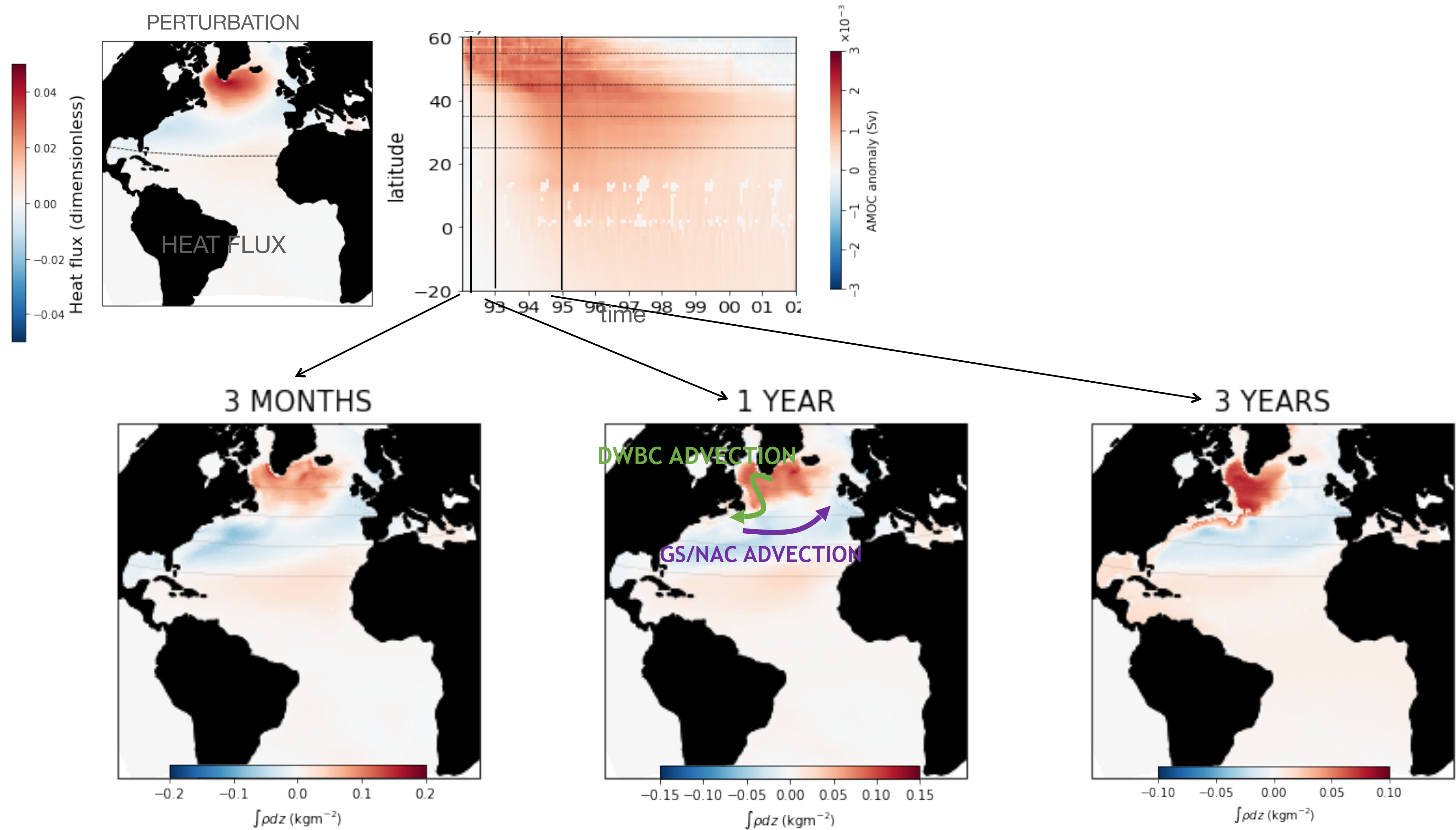
1. Omit a leading WPC pattern



2. Perturb model forcings with the WPC pattern

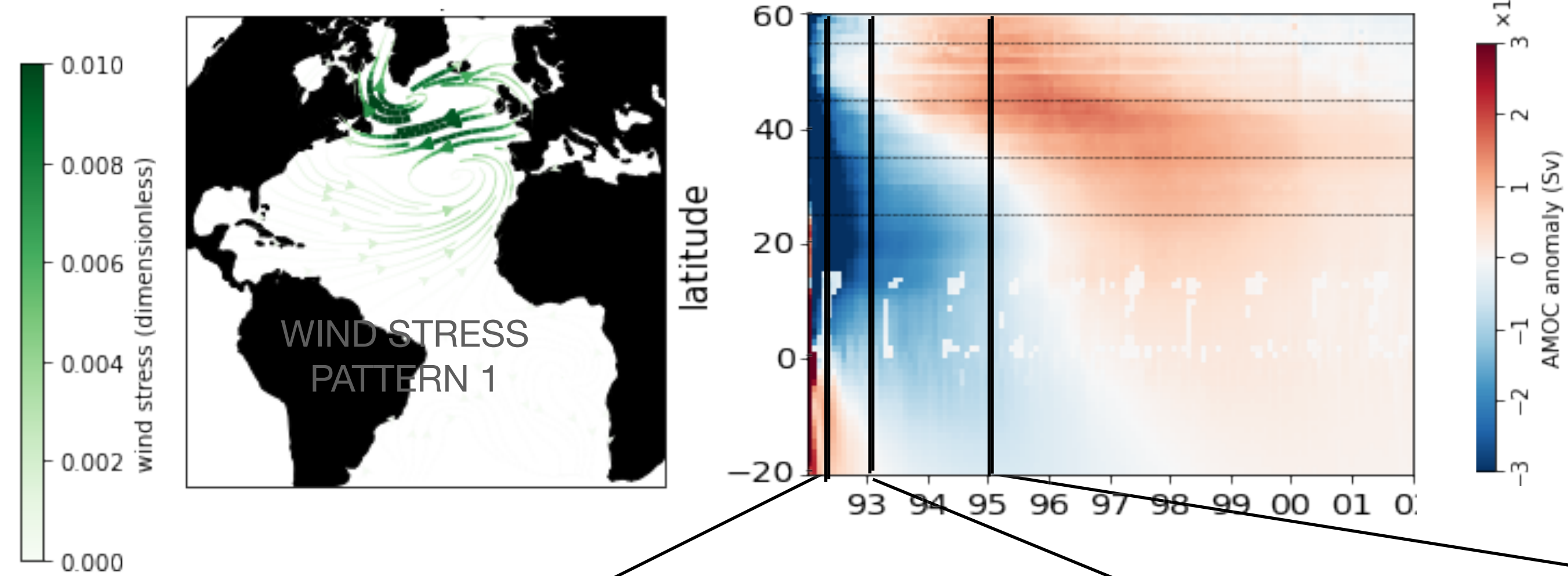


How are low-frequency AMOC anomalies established?

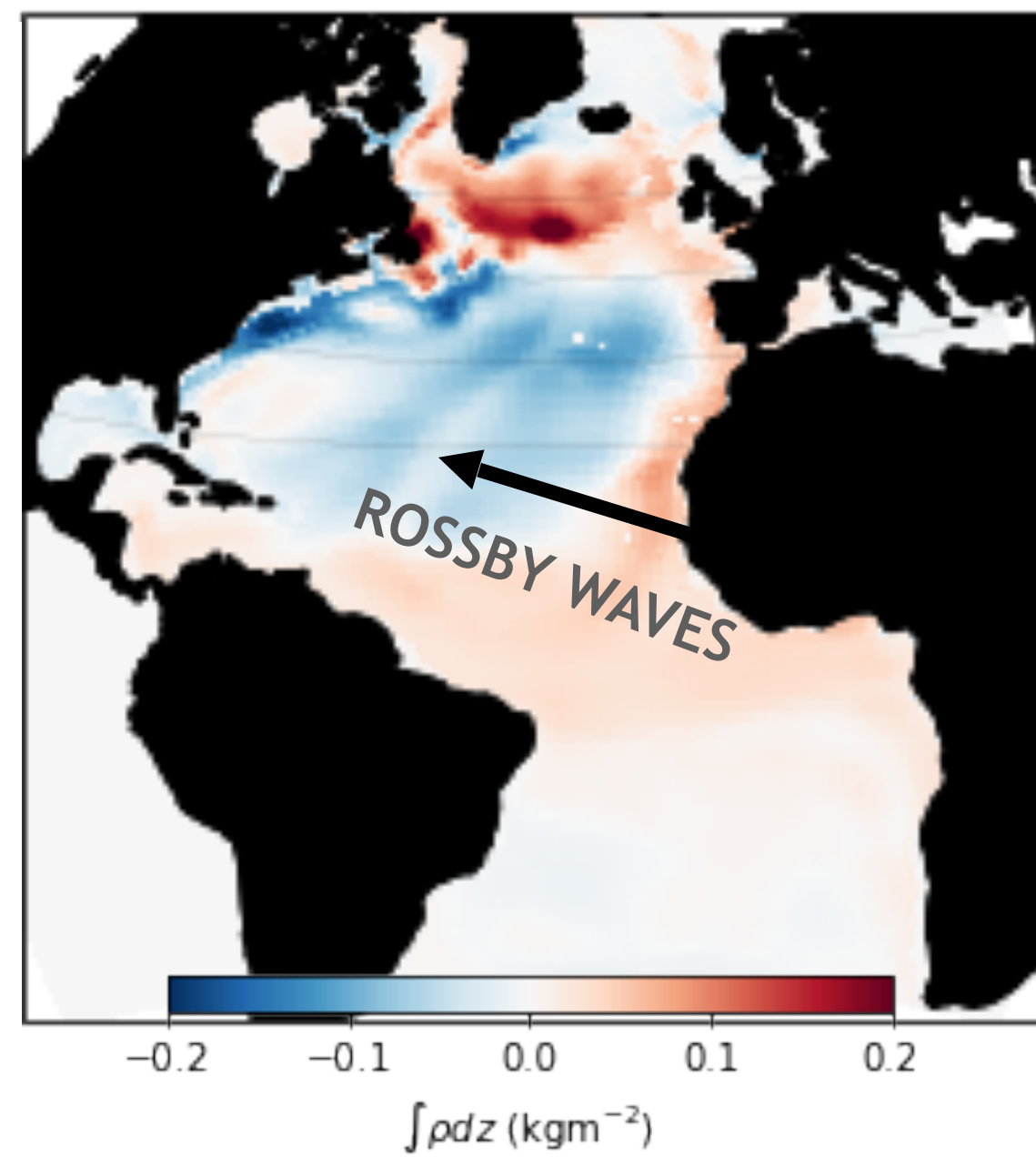


How are low-frequency AMOC anomalies established?

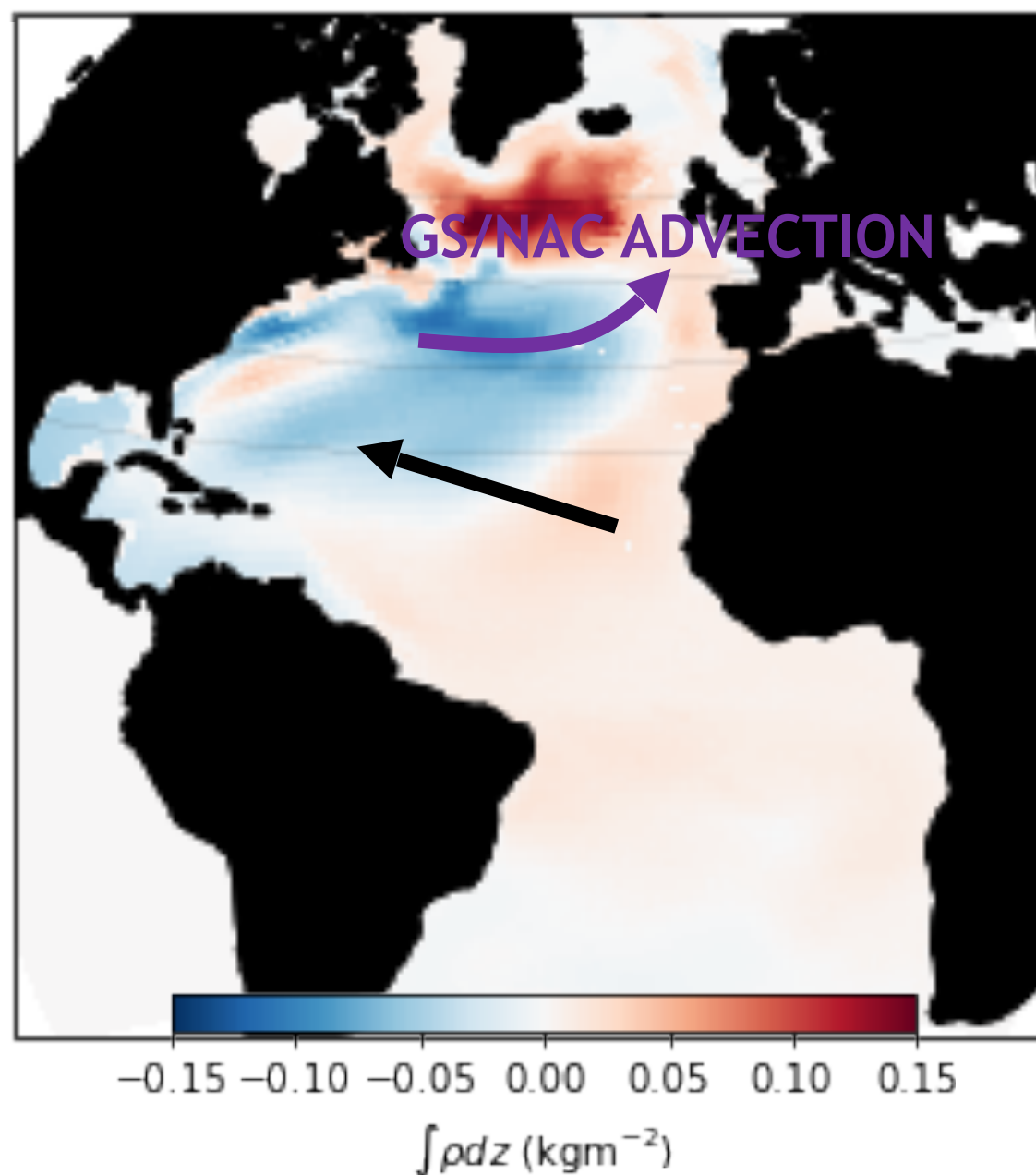
PERTURBATION



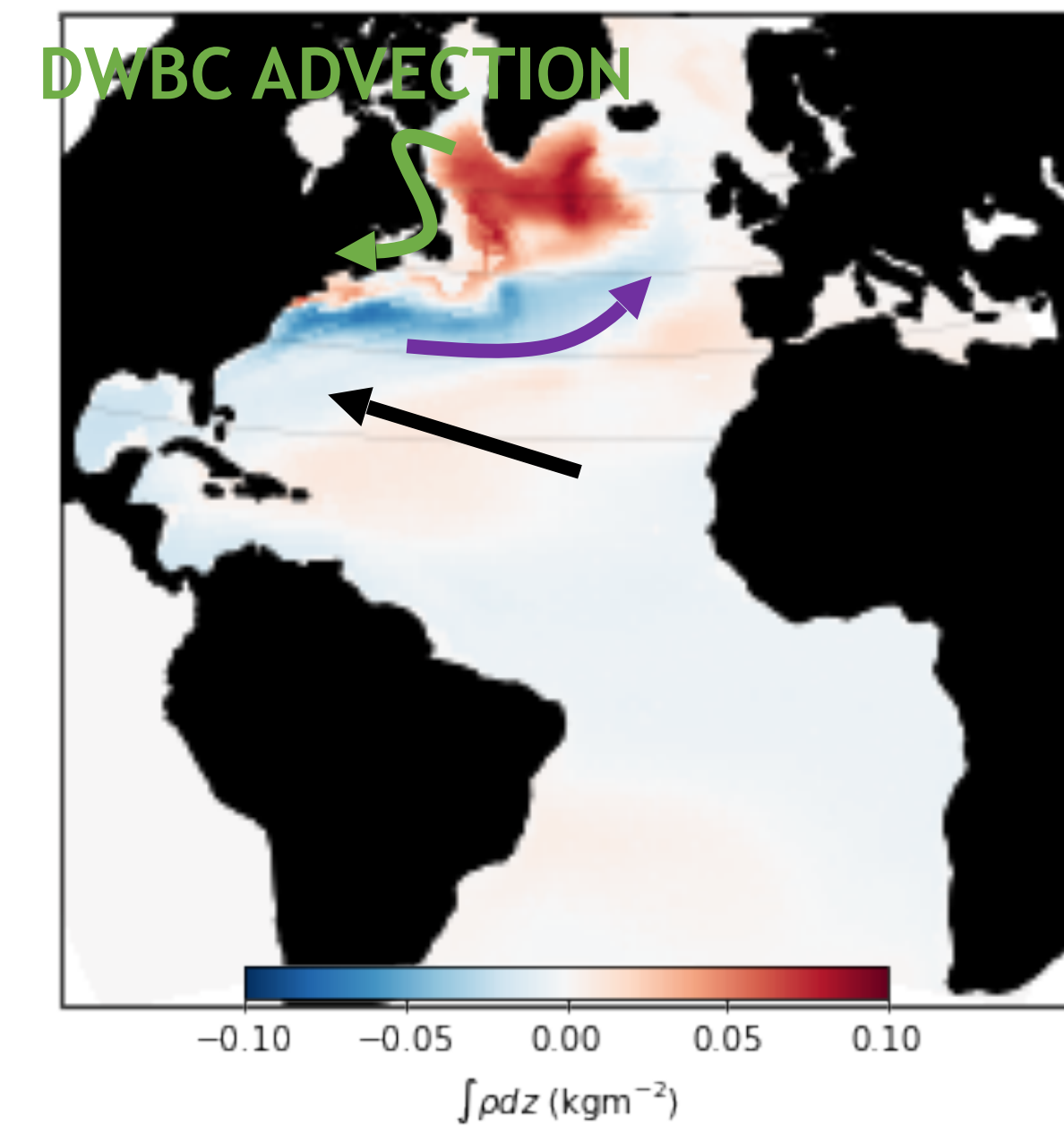
3 MONTHS



1 YEAR

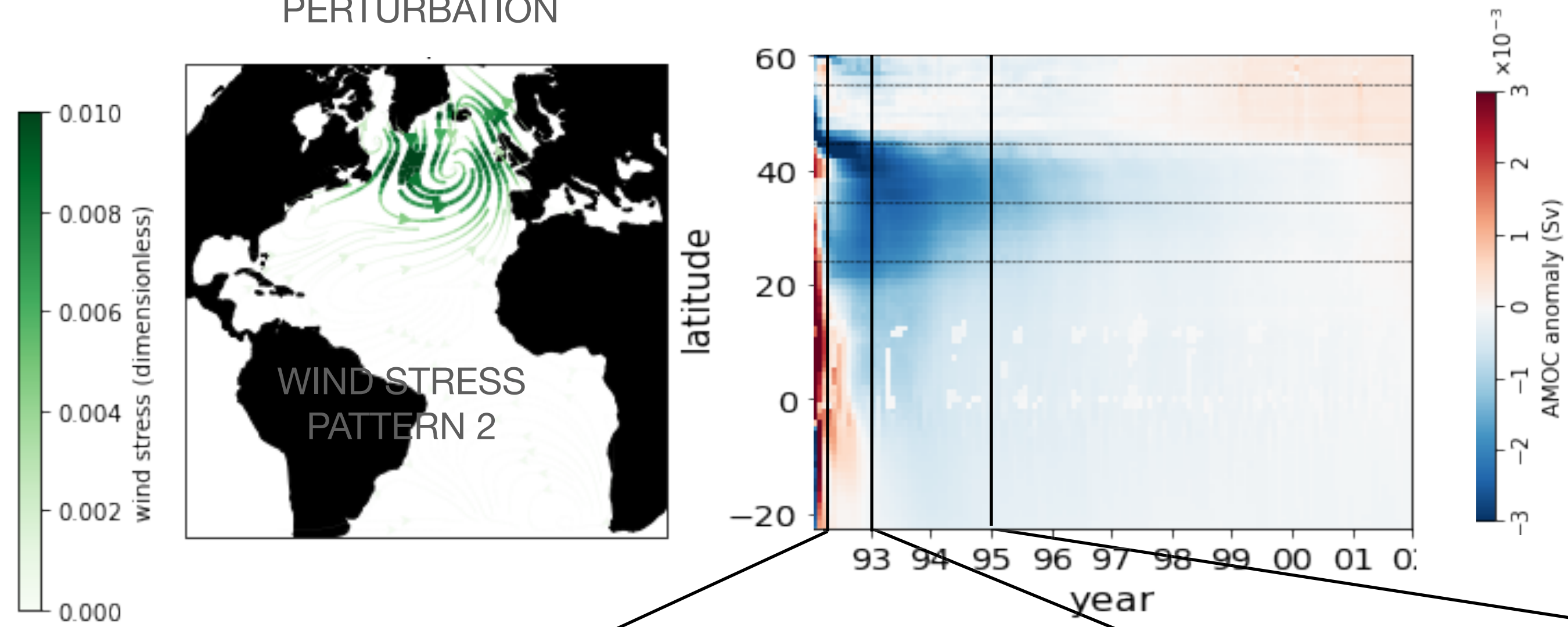


3 YEARS

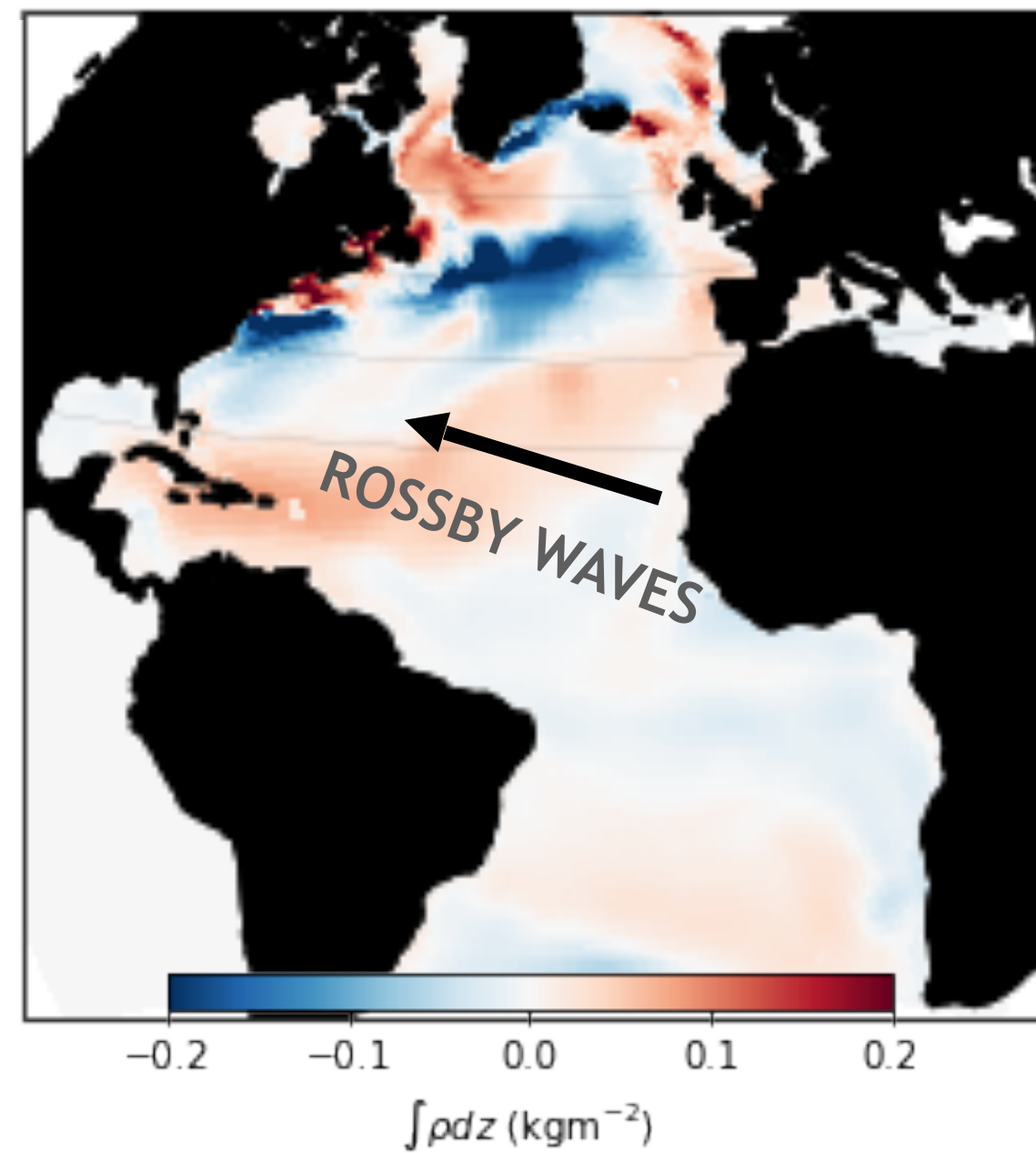


How are low-frequency AMOC anomalies established?

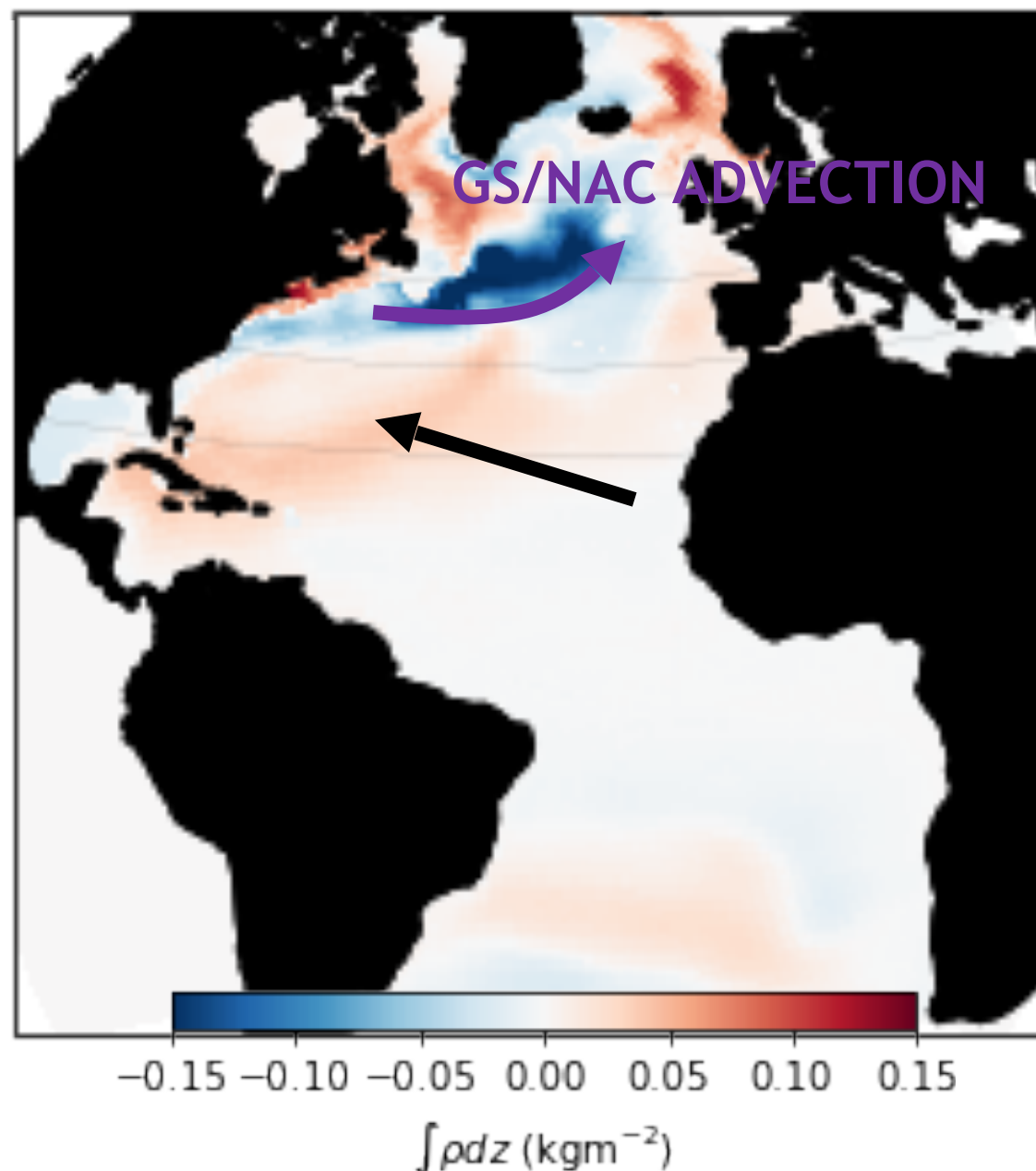
PERTURBATION



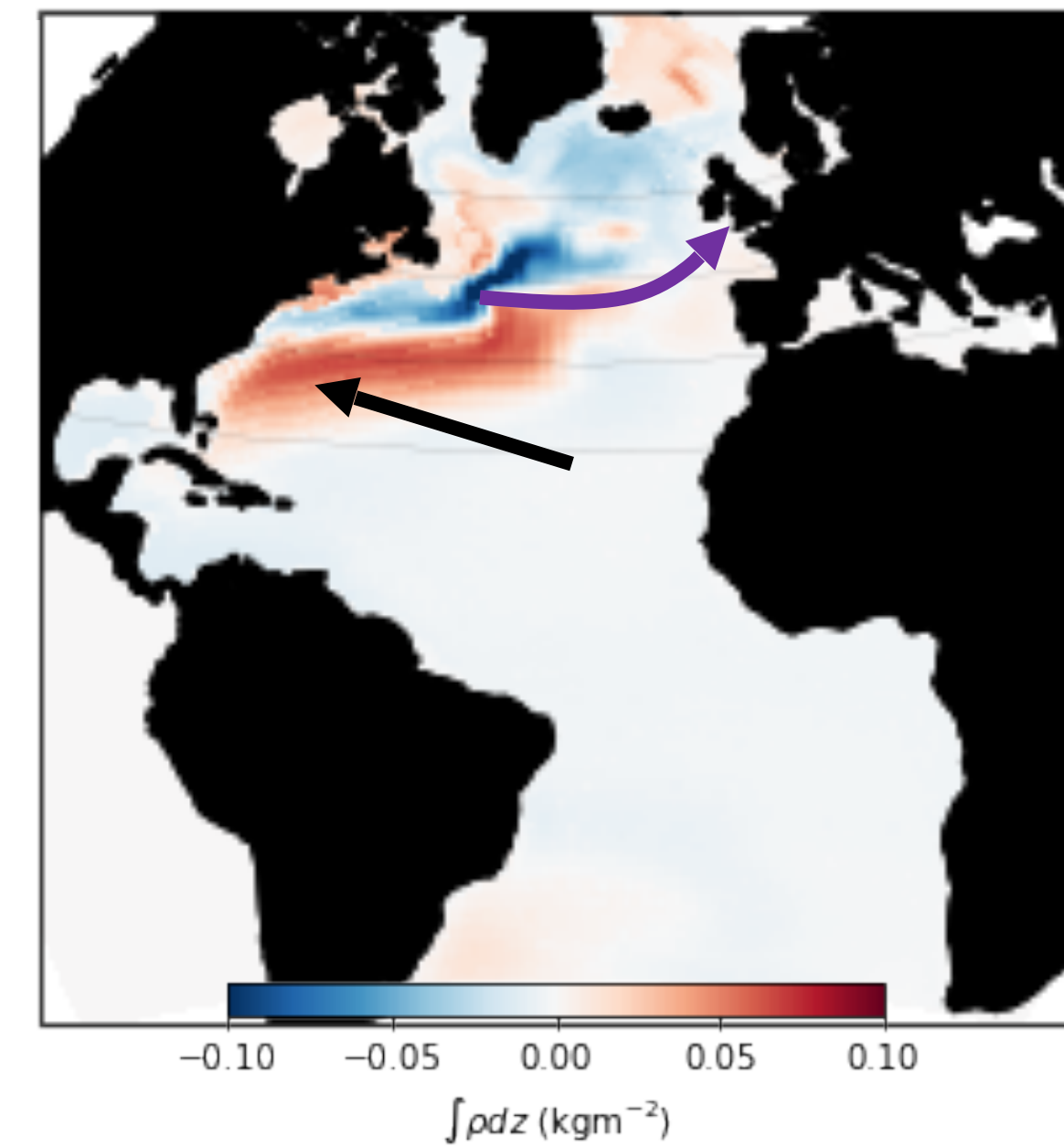
3 MONTHS



1 YEAR



3 YEARS



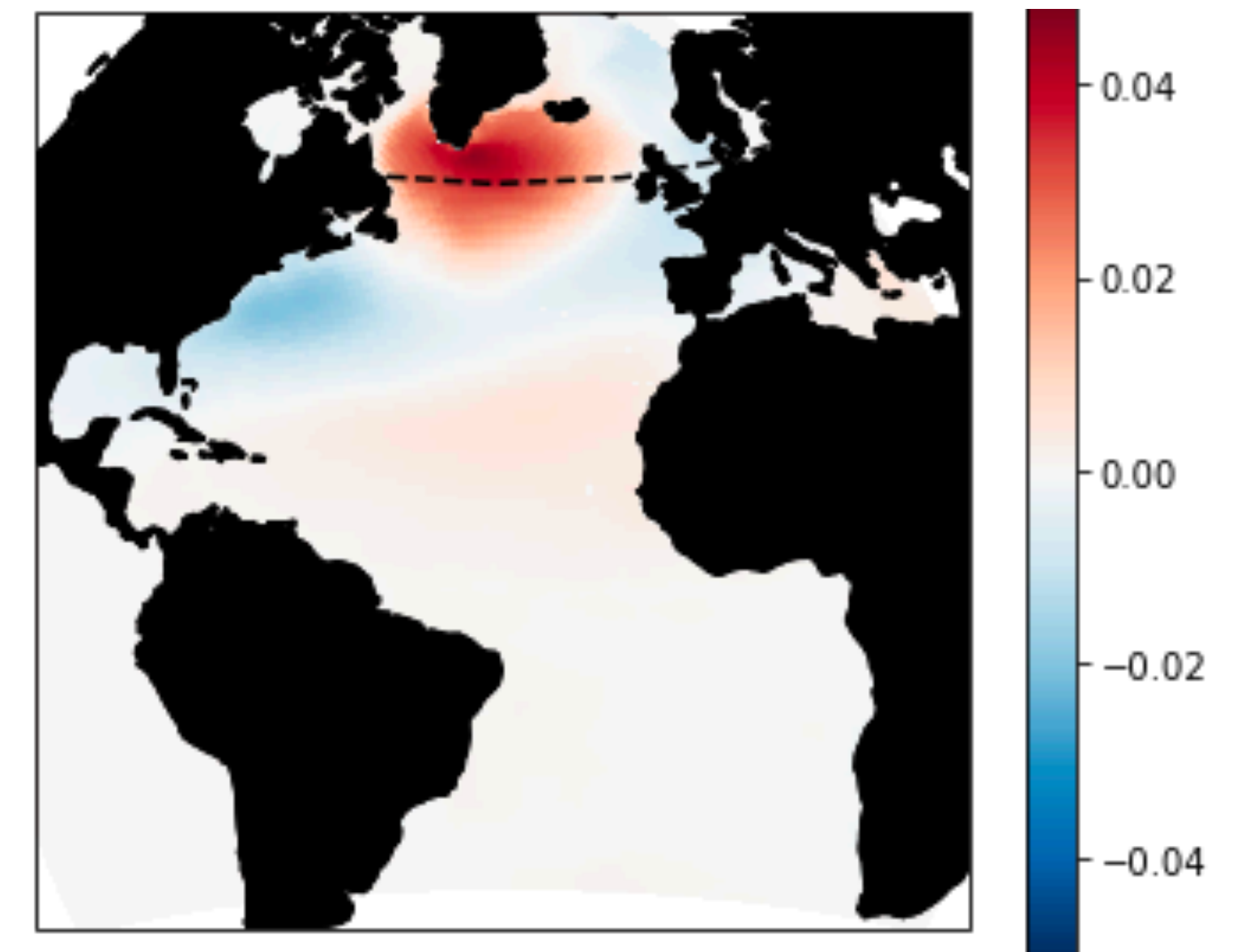
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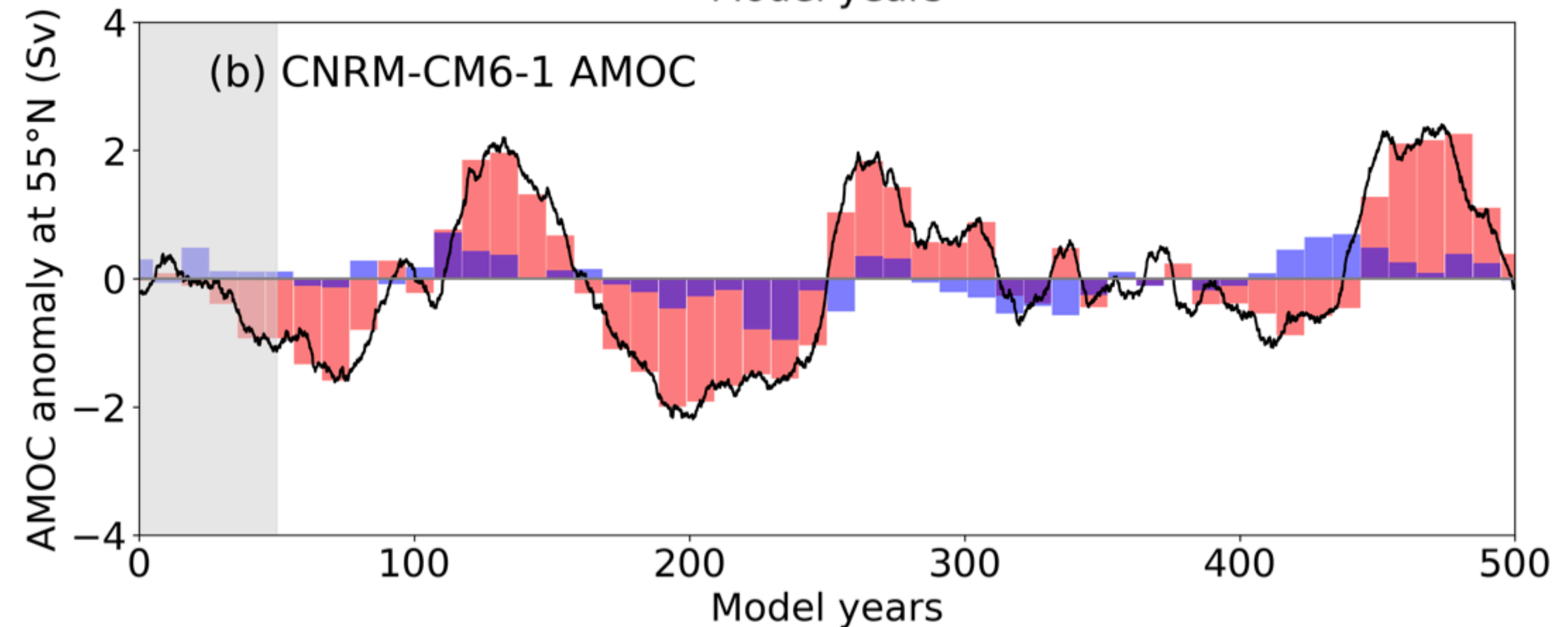
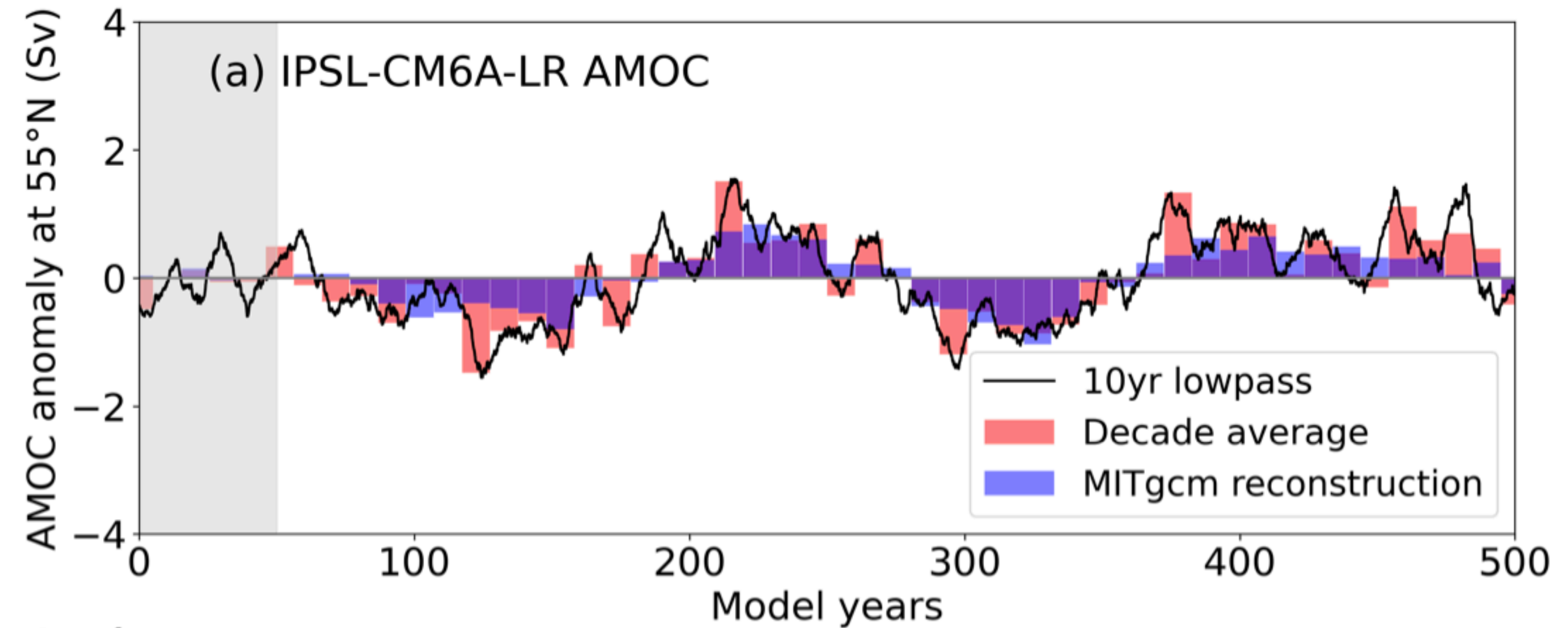


Papers:
Amrhein et al. (methods),
subm. J Clim

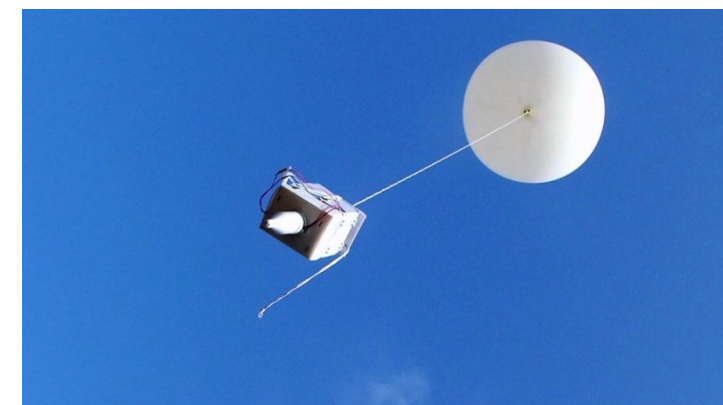
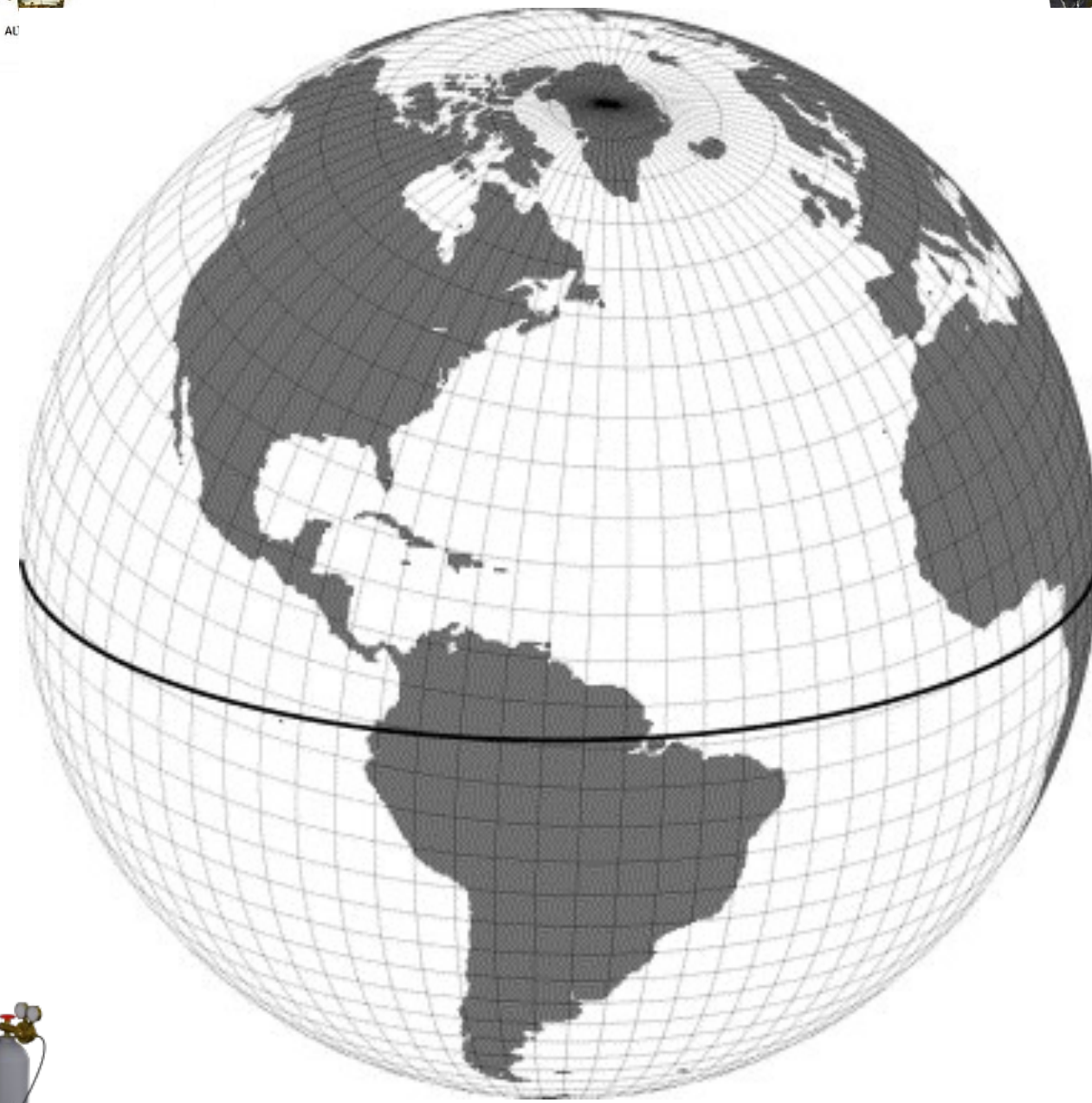
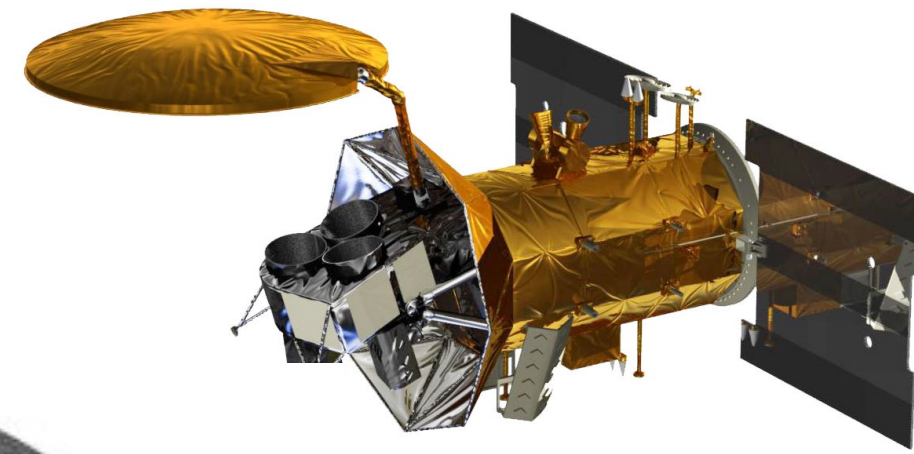
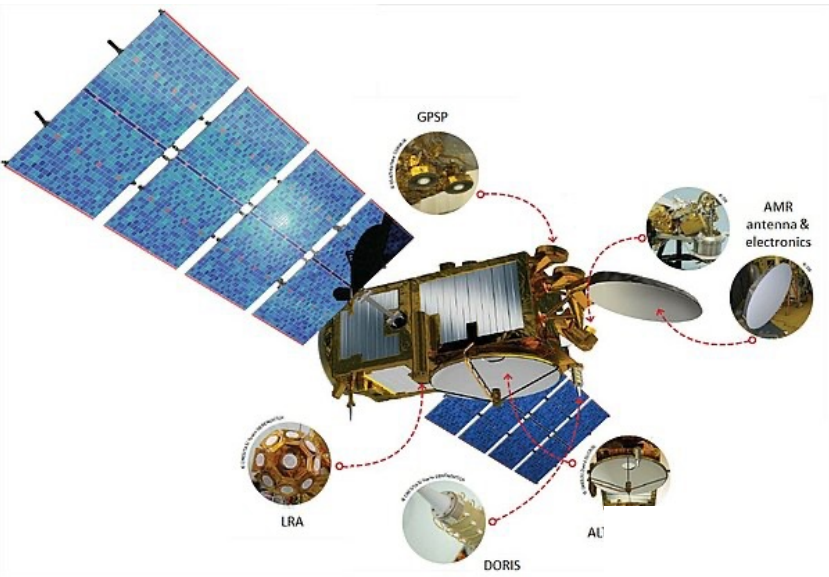
Stephenson et al.
(AMOC results), in prep

Conclusions and future work

Investigating MITgcm adjoint sensitivities in CMIP6 coupled models — a path for quantifying structural uncertainties?



Integration of CESM and the Data Assimilation Research Testbed



CAM6 reanalysis | Global 1°, 80 ensemble members, 2011-2020. Publicly available for forcing CESM.

High-resolution ocean DA | 80-member ocean reanalyses spanning 2011-2017 at 1° and 0.1°

DA and parameter estimation in CLM to improve carbon cycle, hydrologic, and atmospheric forecasting

DA tailored to “bounded” climate quantities (sea ice concentration, tracer concentrations, parameters...)

DA and parameter estimation in MOM6/MARBL

A workhorse DA compset for coupled climate data assimilation in CESM3 with DART

Dan Amrhein, Alper Altuntas, Helen Kershaw, Kevin Raeder, Jim Edwards

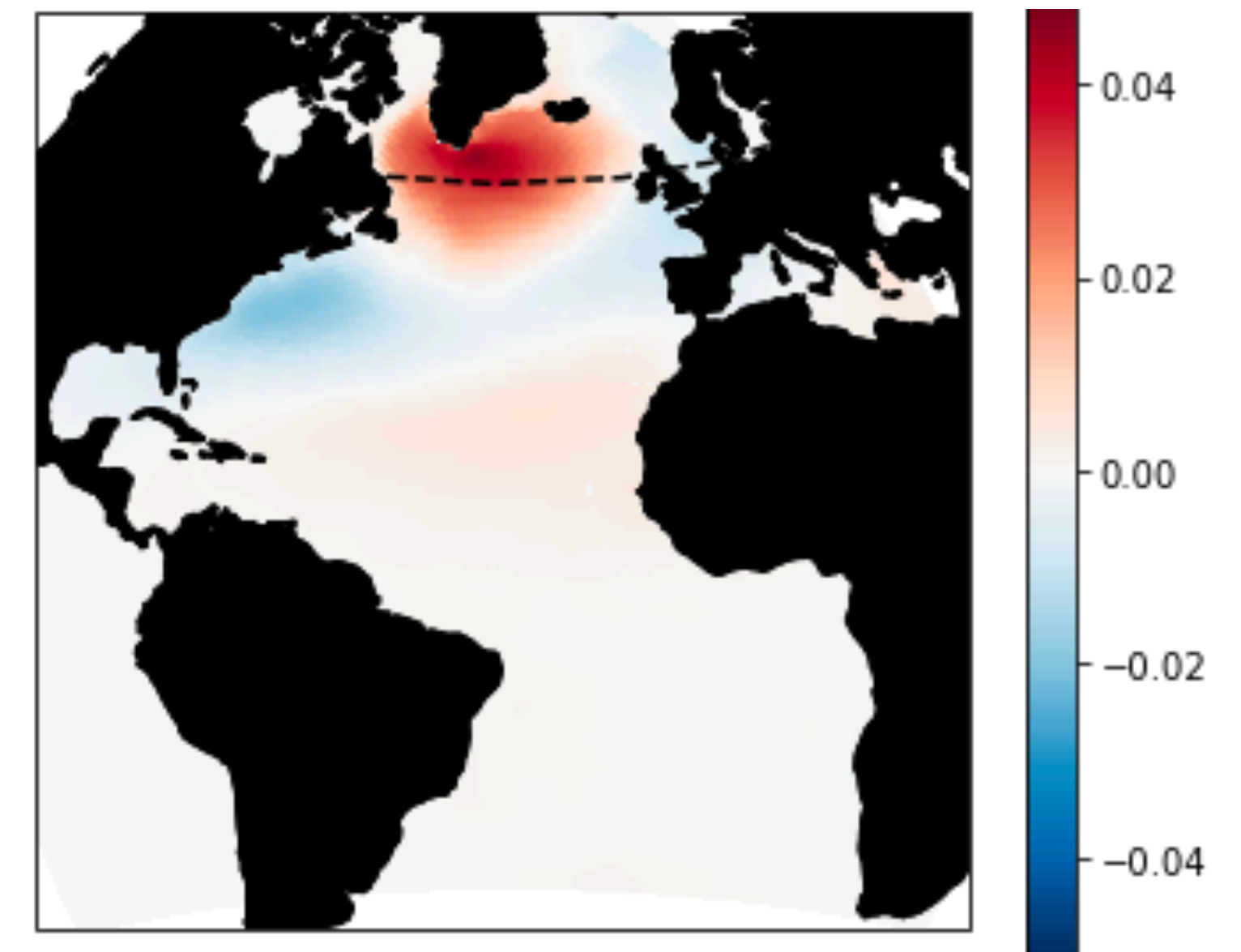
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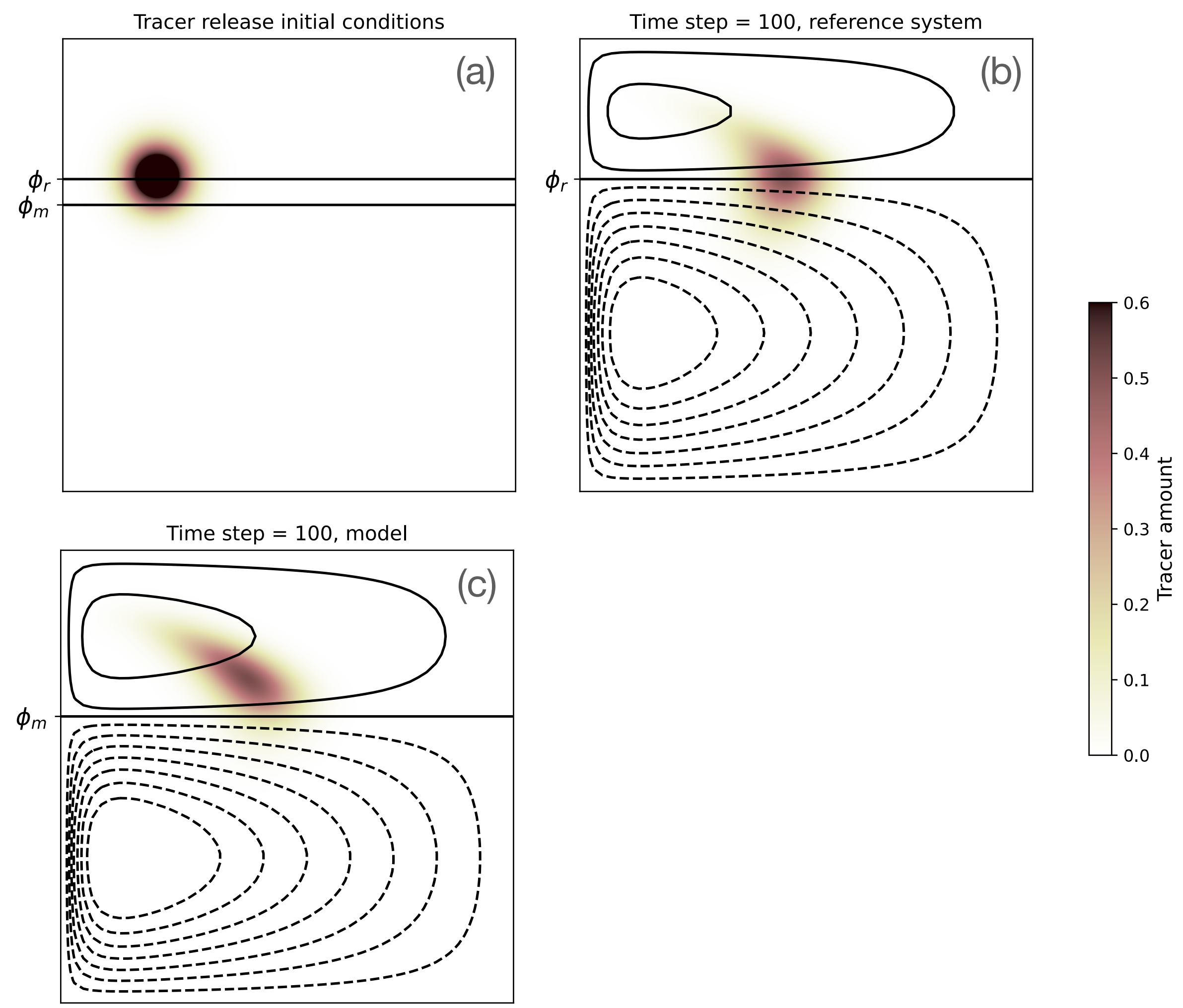
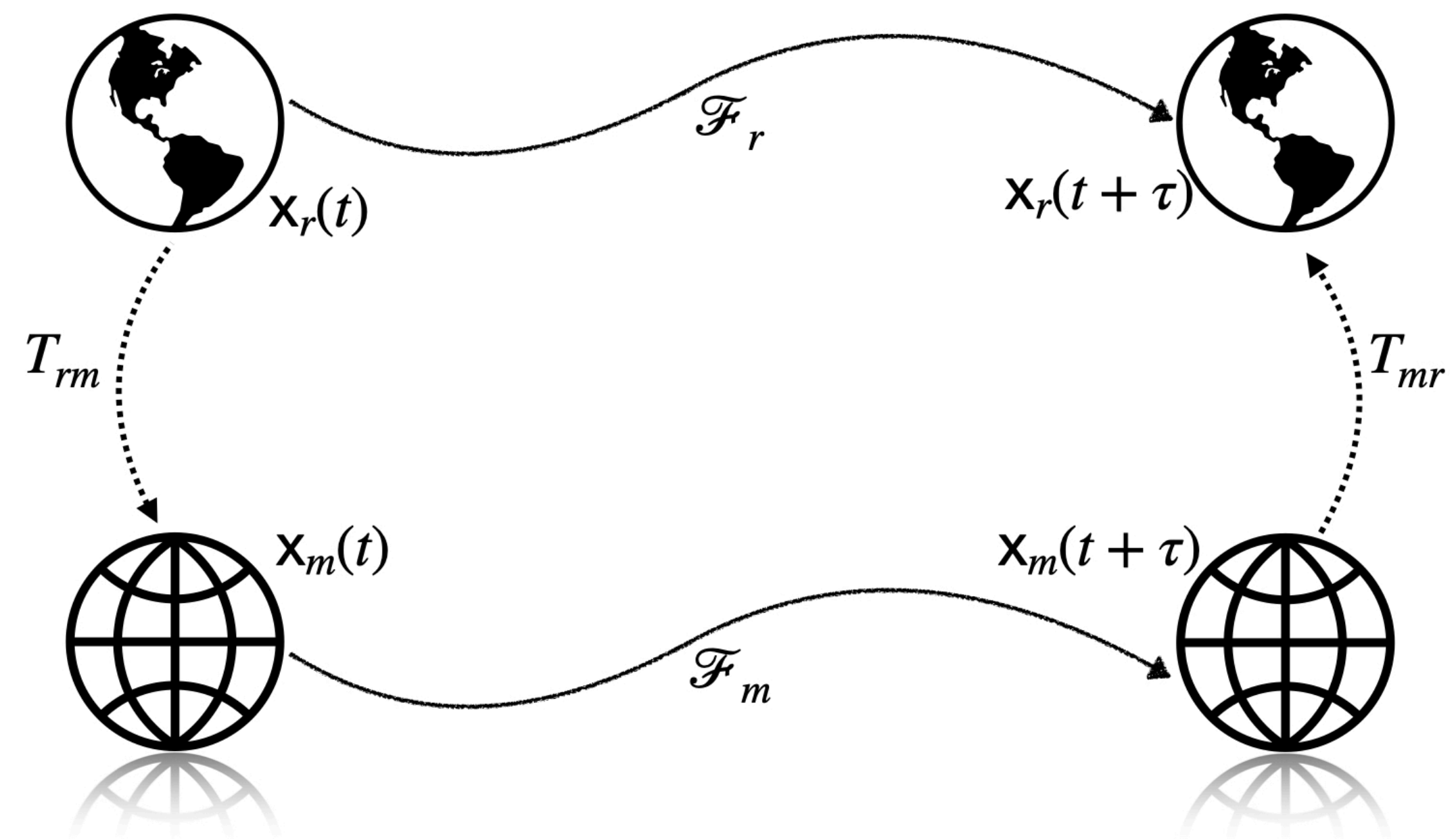
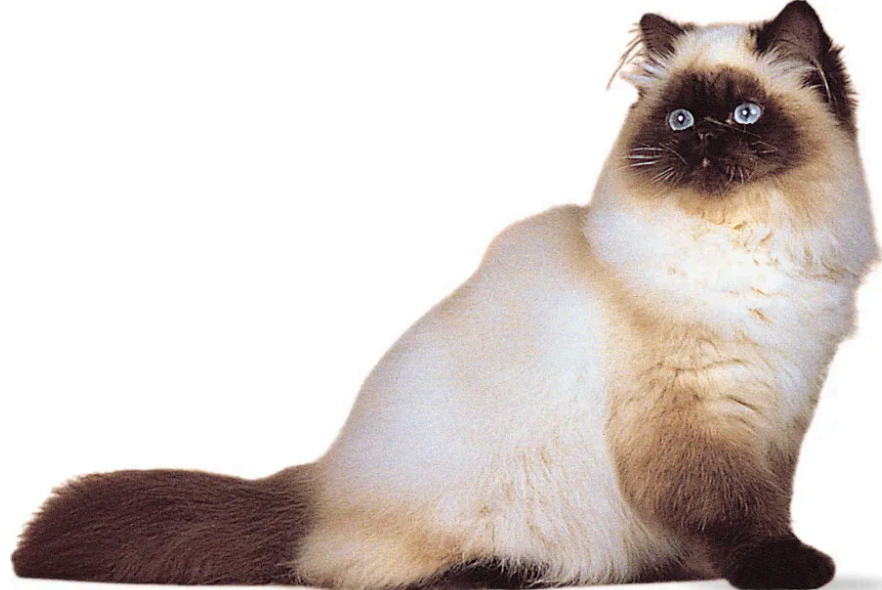
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(AMOC results), in prep

Other things / what's next:

Cross-attractor transforms (CATs)

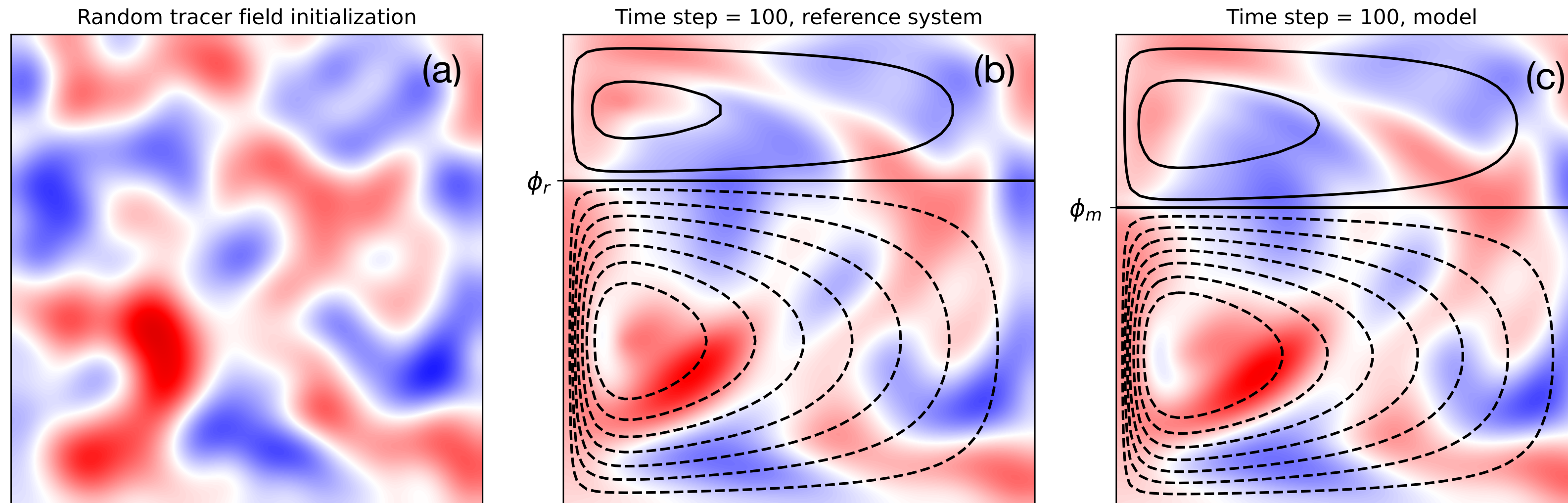
with Ian Grooms and Niraj Agarwal (CU)



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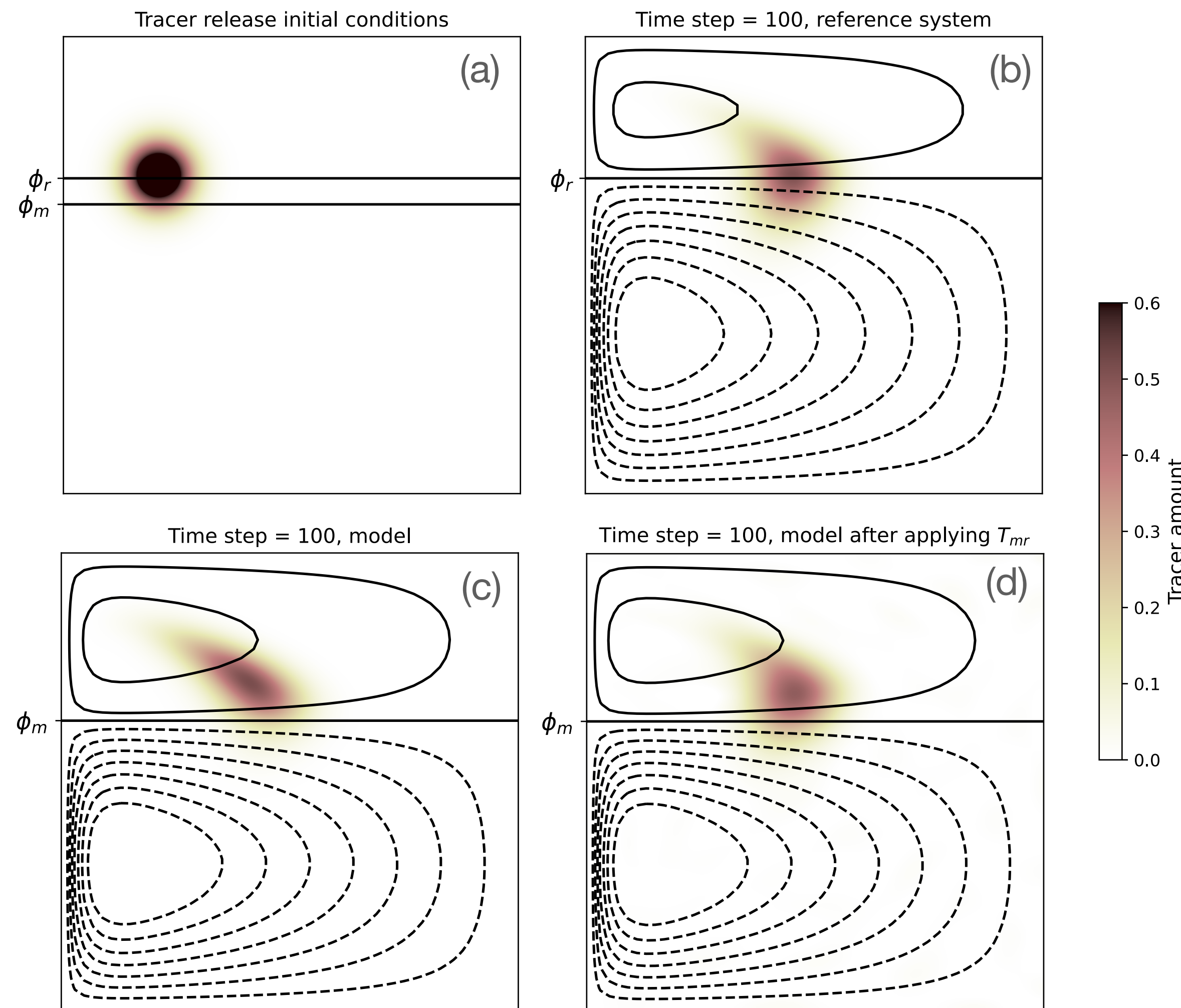
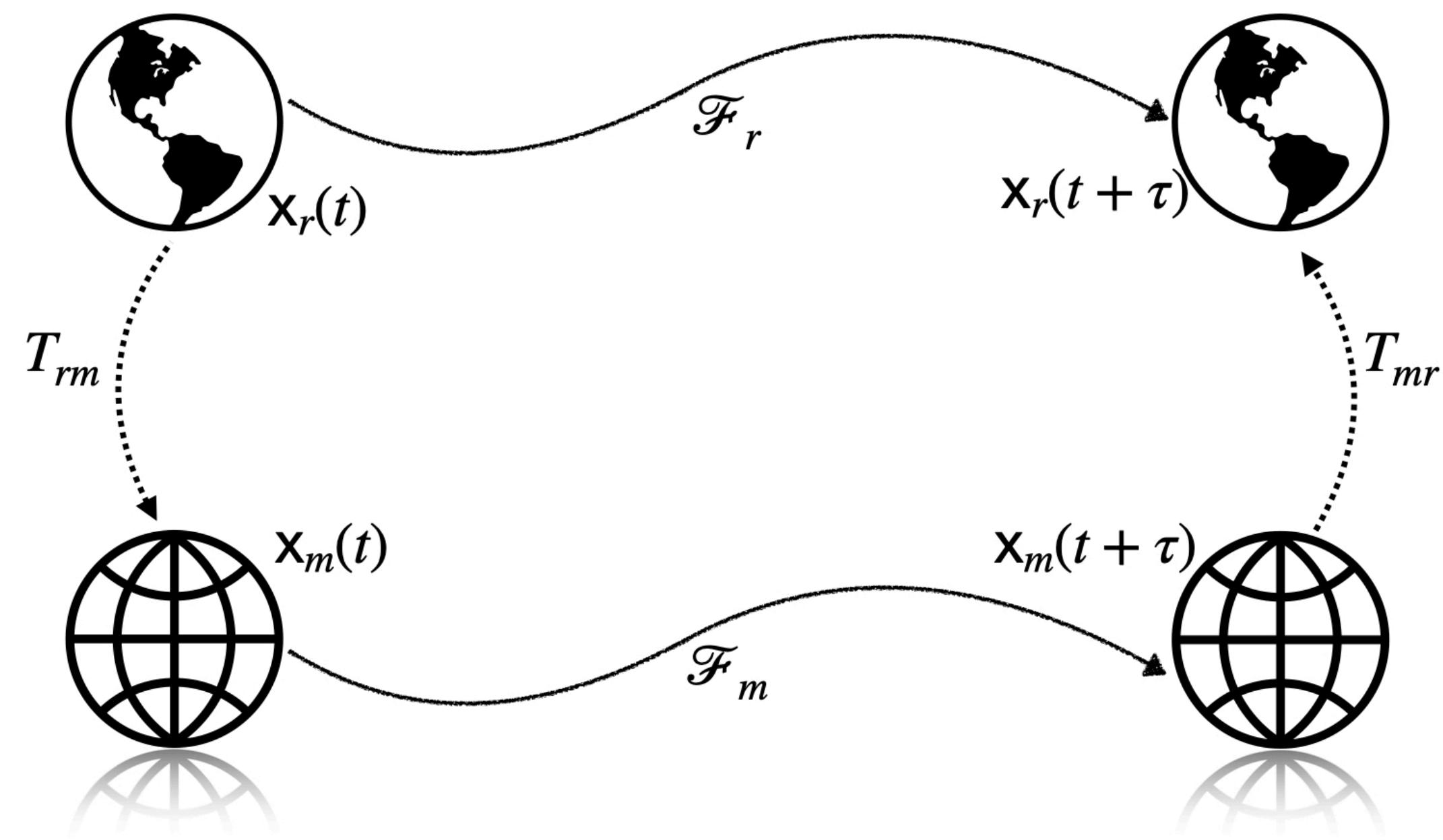
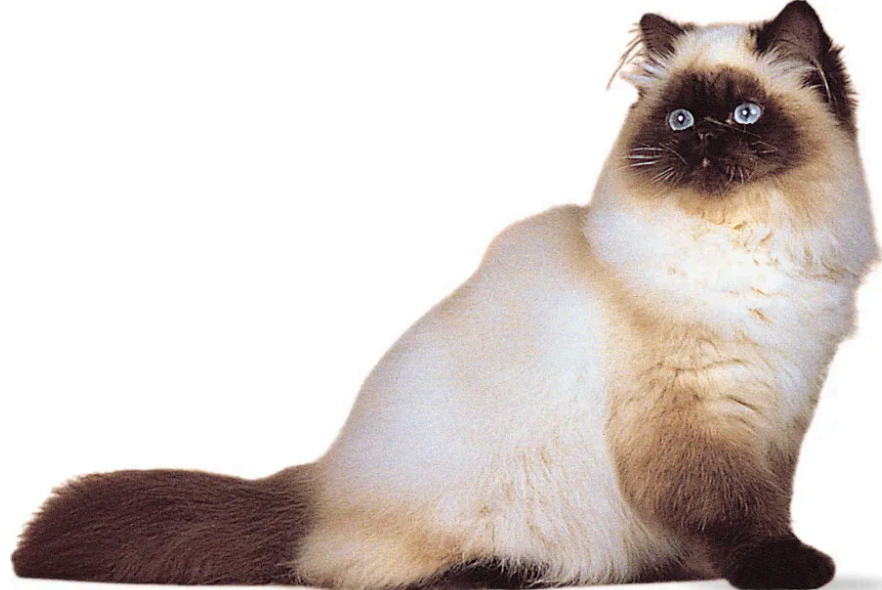


$$\tilde{\mathbf{T}}_{mr} = \mathbf{x}_m^+ \mathbf{x}_r^{+\top} \left(\mathbf{x}_m^+ \mathbf{x}_m^{+\top} \right)^{-1}$$

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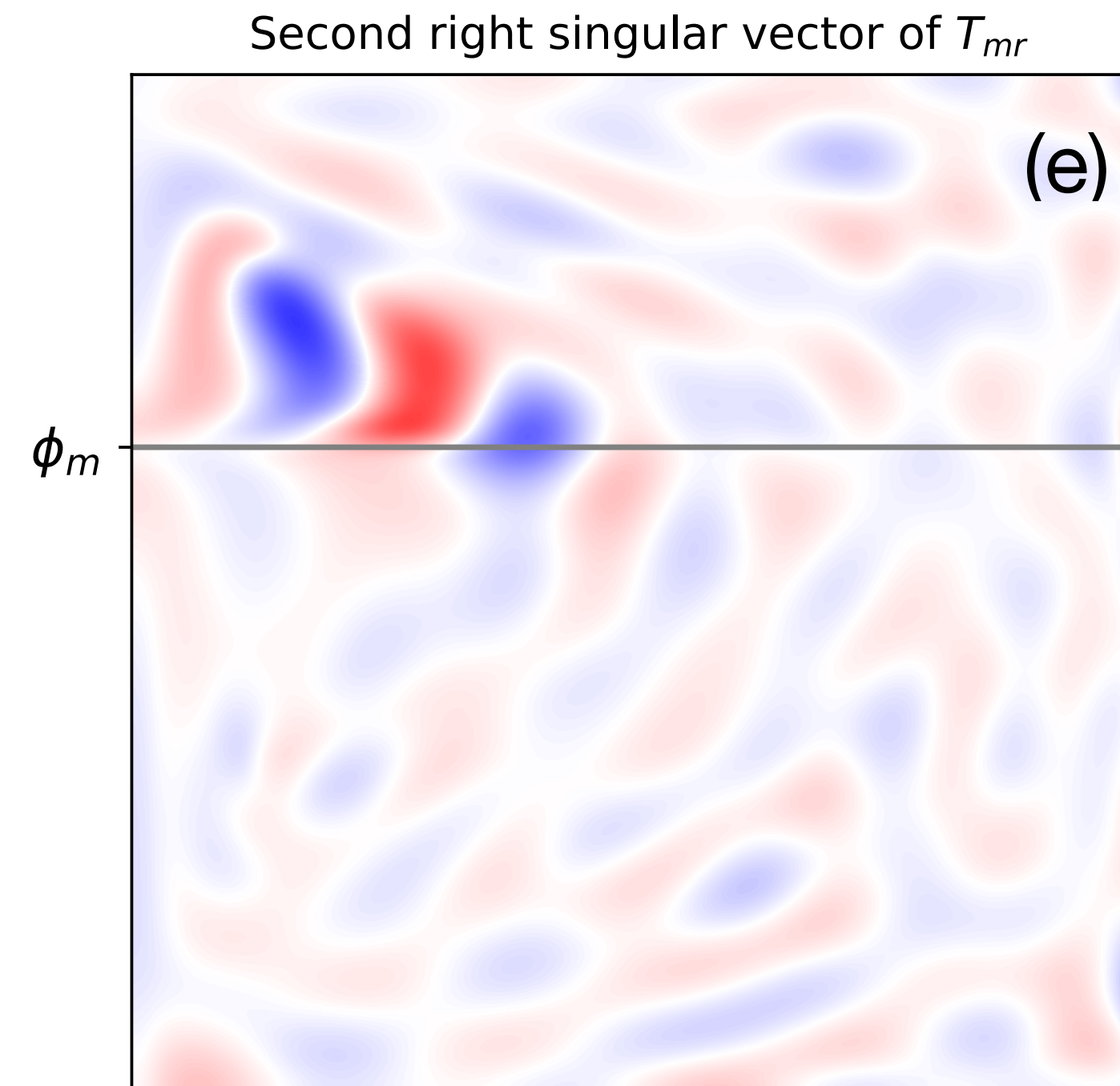
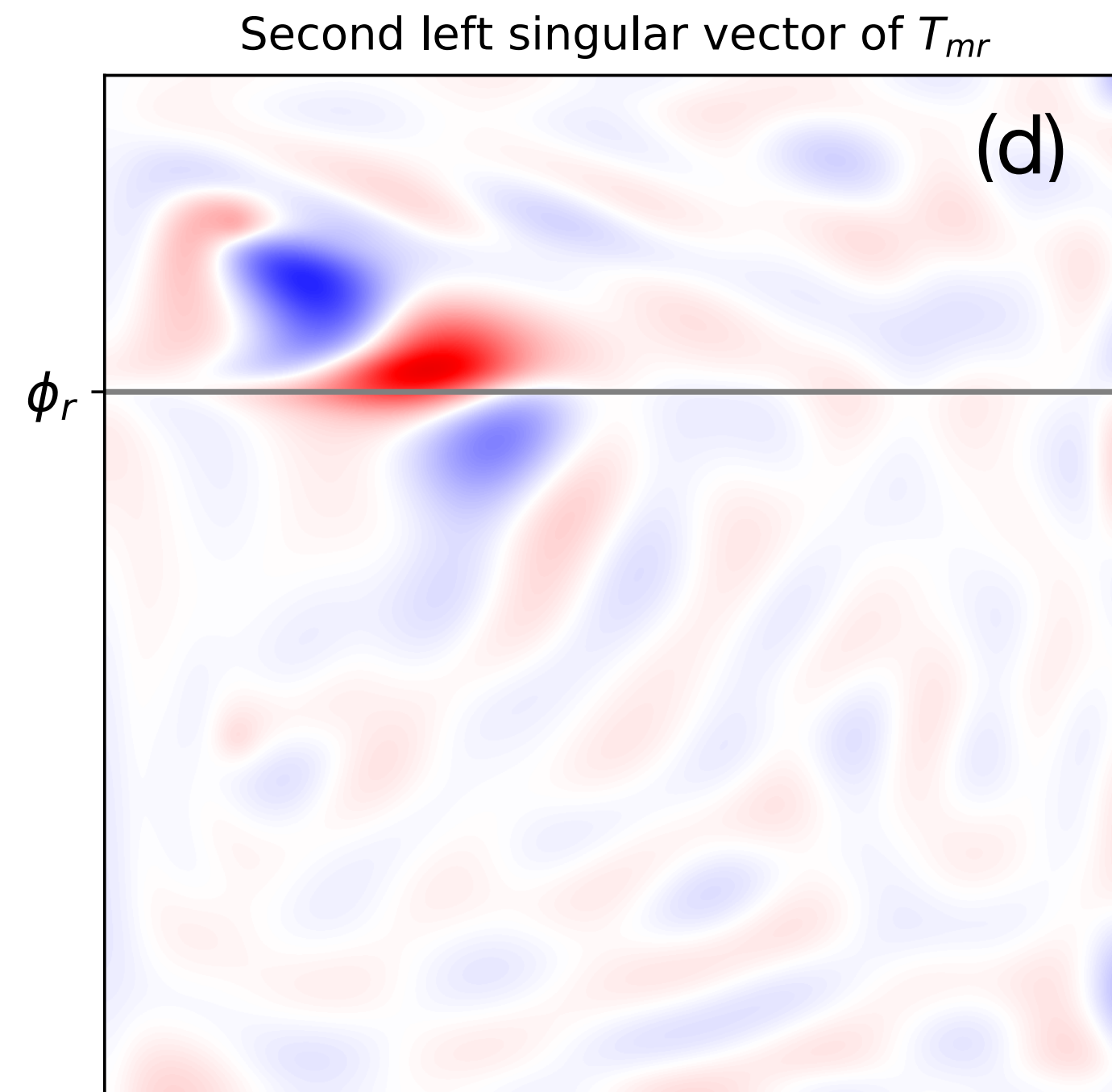
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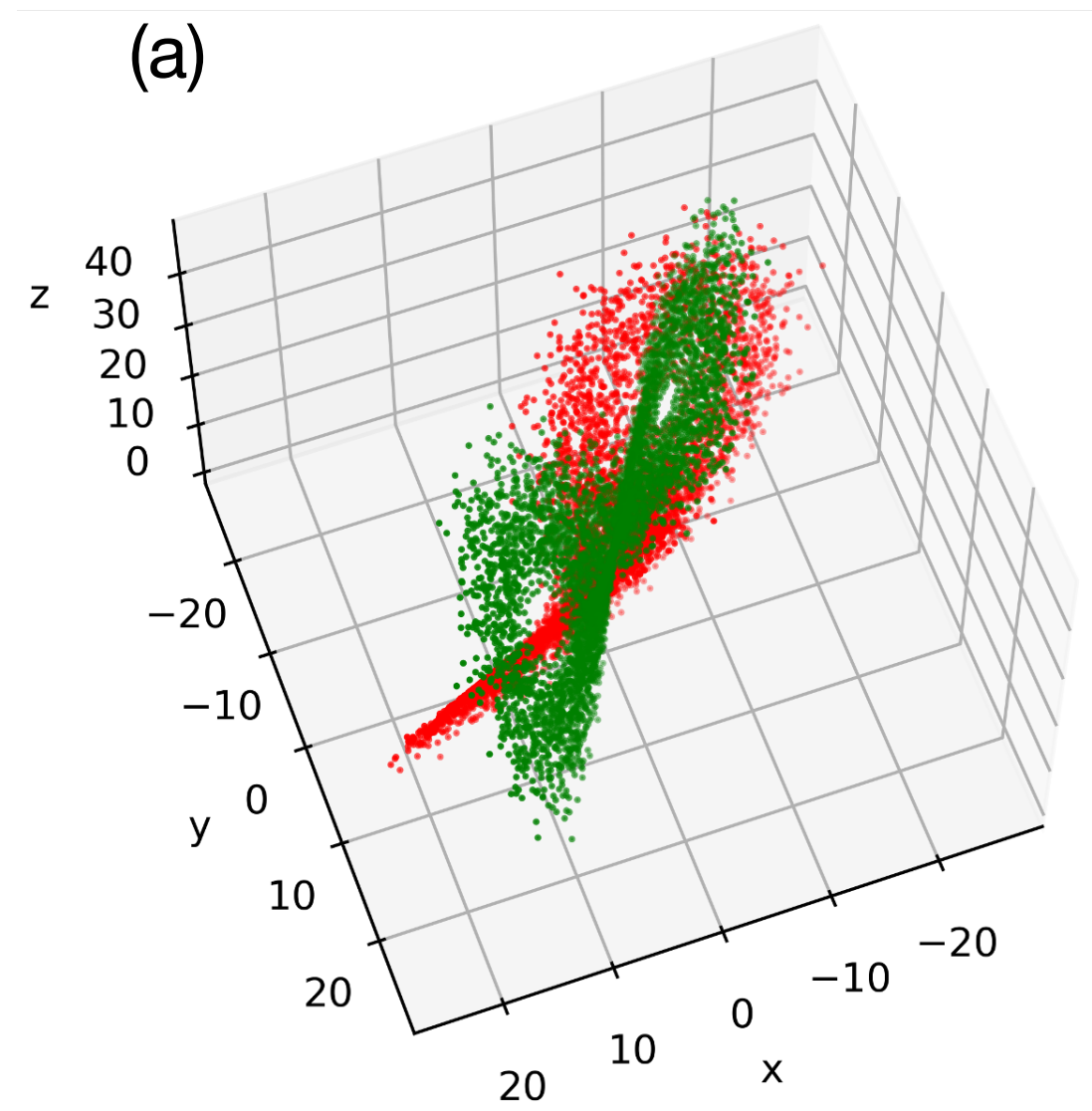
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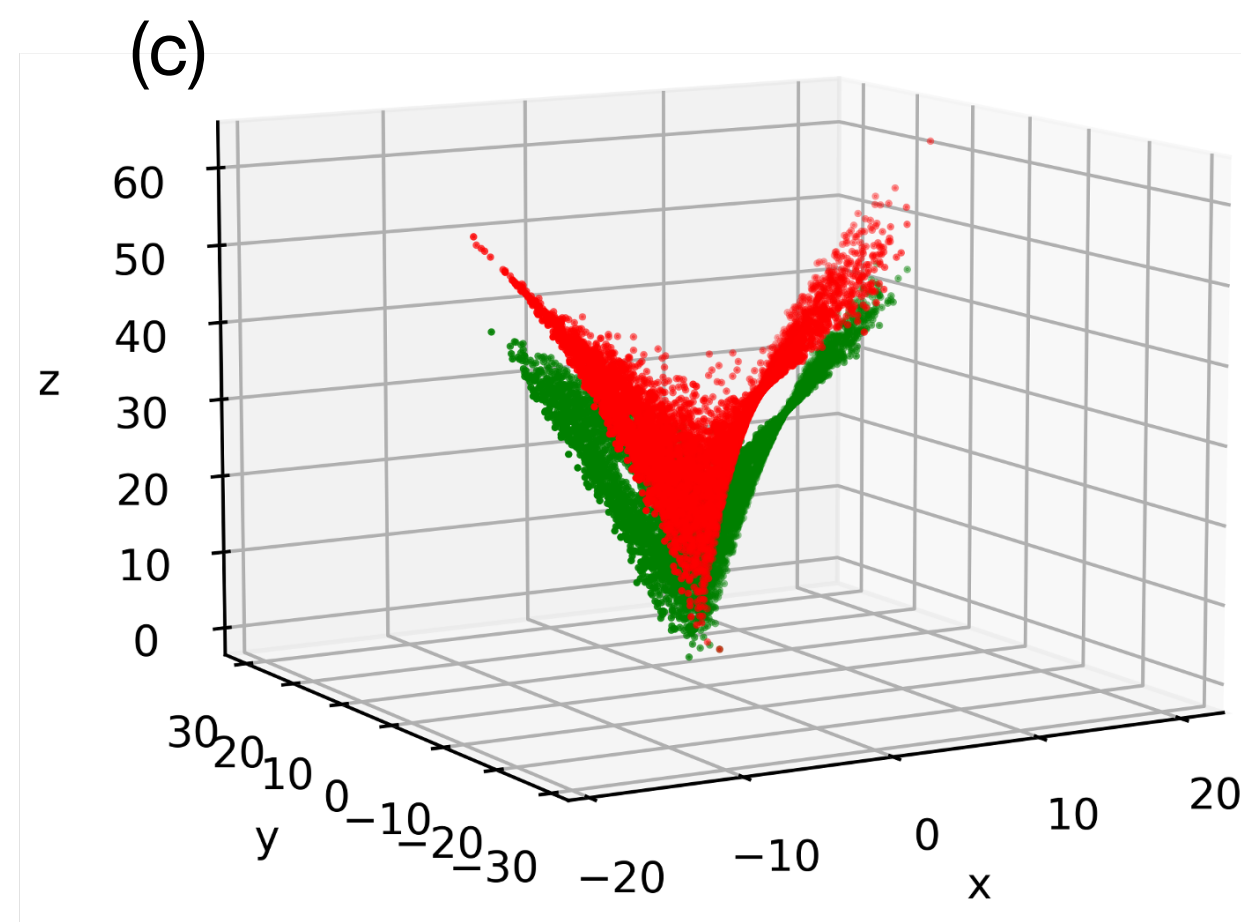
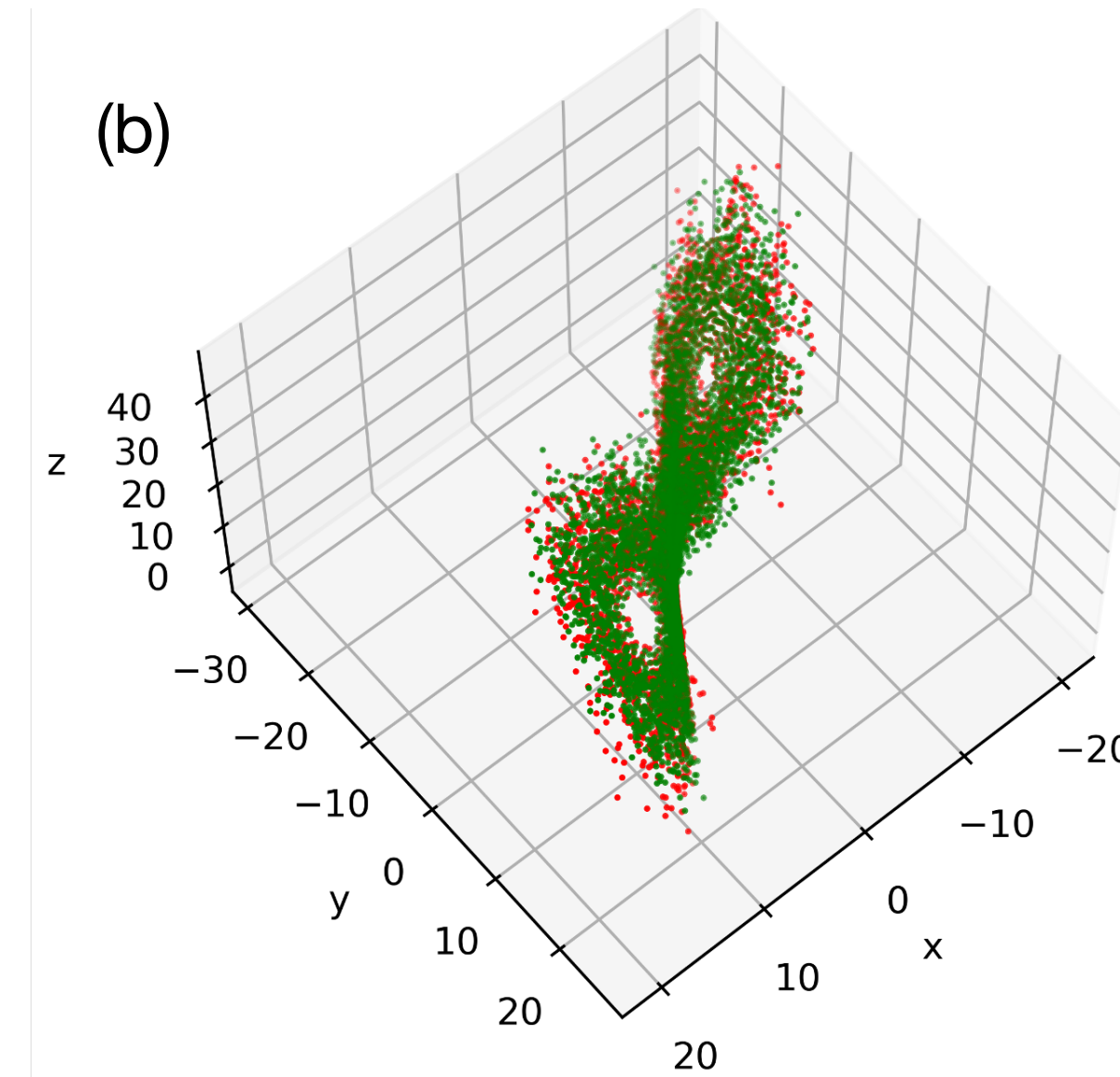
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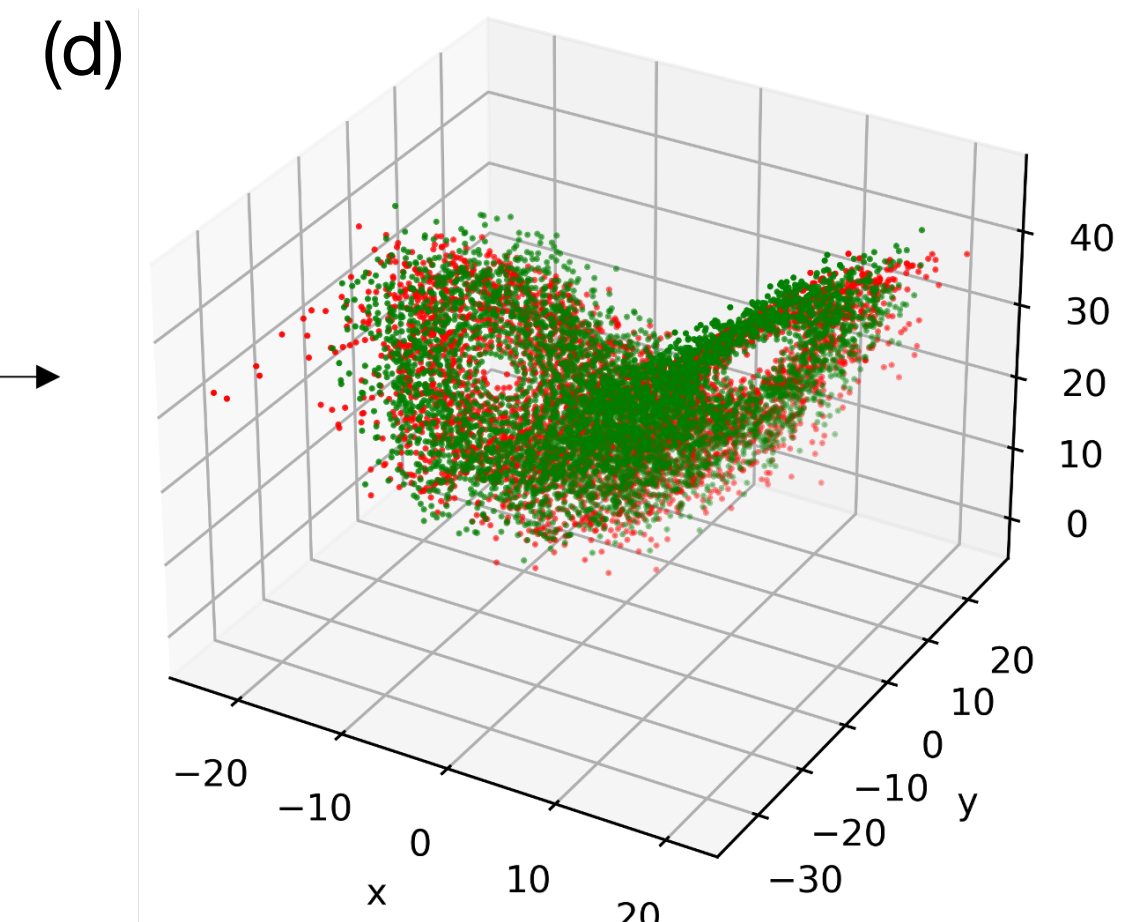
CATs

1 hidden layer with 3 neurons;
Linear activation;
MAE loss; Adam optimizer;
Standardized inputs; 50 Epochs; 32
batch size



CATs

NN configuration same as
in case 1



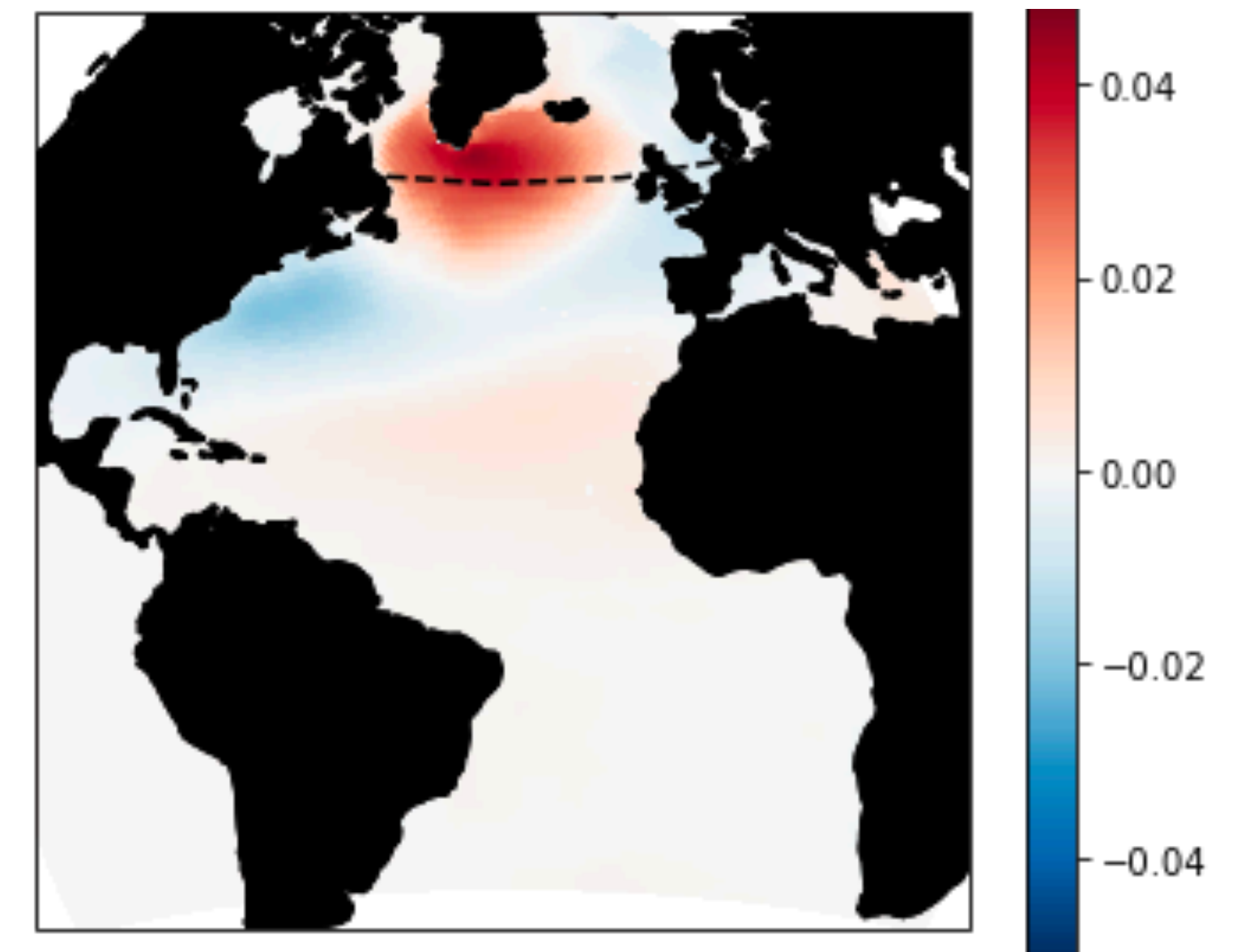
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Ocean surface temperature variability: Large model–data differences at decadal and longer periods

Thomas Laepple^{a,1} and Peter Huybers^b

